

Doctoral Dissertation

GEOCHEMICAL EVALUATION OF MATURITY OF POCKET BEACH SANDS IN SOUTHWEST JAPAN

(西南日本のポケットビーチの砂の成熟度の地球化学的検討)

BY

BAH MAMADOU LAMINE MALICK

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DOCTORAL COMMITTEE

Hiroaki ISHIGA Professor Dr. Sci. (Dept. of Geoscience) Toshiaki IRIZUKI Professor Dr. Sci. (Dept. of Geoscience) Yasushi SEIKE Professor Dr. Agr. (Dept. of Material Science) Tetsuya SAKAI Associate Professor Ph.D. (Dept. of Geoscience) Hiroki HAYASHI Associate Professor Ph.D. (Dept. of Geoscience) Andreas AUER Associate Professor Ph.D. (Dept. of Geoscience) Name: Bah Mamadou Lamine Malick
Date of Degree: March, 2017
Title of Study: Earth Science and Geoenvironmental Science
Major Field: Geoscience
Supervisor: Hiroaki ISHIGA Professor Dr. Sci. (Dept. of Geoscience)

ABSTRACT

The geochemical maturity of pocket beach sand was evaluated using major and trace element compositions of sample sets from nine Prefectures in South West Japan. These included northern Kyushu (n=30), Yamaguchi Prefecture (n=27), Shimane Prefecture (n=50), Tottori Prefecture (n=15), Tango Peninsula (n=38) and Noto Peninsula (n=30). The geochemical data obtained by X-ray fluorescence (XRF) analysis were compared to the content of beach sands from Kotobikihama and Kotogahama, which were assumed to be representative of matured sands. Data were also compared with the geochemical compositions of 15 local river sediments from the Geological Survey of Japan and National Institute of Advanced Industrial Science and Technology and with 15 near-shore marine sediments around Yamaguchi. Furthermore, comparison with average Upper Continental Crust of the Japanese Archipelago, and average Upper Continental Crust was performed. The relatively high concentration of quartz in the silica-rich sands from Kotogahama, Kotobikihama, Shimane and Yamaguchi was reflected in the geochemical analysis of those sands, the major and trace element compositions being characterized by high SiO_2 contents. The Tango Peninsula and Wakasa Bay were very similar, showing a moderate geochemical maturation. Beach sands from Tottori and Noto Peninsula had lower SiO_2 and Al_2O_3 values, which reflected the abundance of feldspar, suggesting geochemical immaturity, and relatively high K_2O and Na_2O associated with feldspars. CaO contents were generally low, although enrichments occurred in a few samples

due to presence of shell material. Following geochemical classification of the coastal beach sands from the six regions of South West Japan, a growing abundance of both quartz and feldspar was indicated by the sands being bracketed by arkose and subarkose, with a diminishing trend towards sub-litharenite. The relatively low-to-moderate values of weathering indices of Chemical Index of Alteration (CIA), Plagioclase Index of Alteration (PIA) and Chemical Index of Weathering (CIW), indicated that the beach sands from the sites in the source area have undergone low to moderate degree of chemical weathering. A-CN-K and A-CNK-FM plots, which suggest a granitic source composition, also confirmed that the sand samples from these sites have undergone a low to moderate degree of chemical weathering in consistent with CIA, PIA and CIW values. Investigated beach sands from the six coastal regions of South West Japan comprised variable mixtures of terrigenous detritus (represented by Al_2O_3 and SiO_2) and biogenous material (represented by CaO). The primary component of beach sands from Shimane was quartz, or silica (SiO_2) , Sands from Tottori were composed largely of weathered feldspar particles, while, in contrast, components of biogenic and quartz-rich sands from Yamaguchi were primarily shell fragments, quartz, and igneous rock. The sands of northern Kyushu might have been expected to exhibit a relatively high carbonate content, not least on account of the warm-water currents there. However, the water quality is poor (type B), which is likely to explain the low carbonate contents measured. Plentiful warm-water species and high-quality (type AA) water in the Yamaguchi area were reflected in the high to moderately low carbonate content of the beach sands from that location. Contents of local river and near-shore marine sediments differed distinctly from those at Yamaguchi, suggesting that inputs of existing river or marine sediment to the beach from currents or storm events are minimal.

Key words: Pocket beach, beach sand, foreshore, geochemistry, silica, biogenic, southwest Japan

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Chapter One

1. INTRODUCTION

Coastal region can be considered as a place with its own characteristic qualities and features, in which the inhabitants perform activities, both economic and social, that are specific only to that region. Numerous components and processes, including source composition, sorting, climate, relief, long shore drift, and winnowing by wave action, influence its sediments. Sandy beaches are regions of transition from the shore to the sea, and are subjected to significant influences from both ecosystems. These regions are open to significant variations in the length of exposure to the sun, immersion and submersion, the amount of rainfall and concentration of nutrients. Among other factors, beaches are also subject to local processes such as wave and tidal regimes, fluvial discharges, and wind transport (Carranza-Edwards et al., 2009). The geochemical composition of beach sands is influenced by many components and processes. These (components and processes) contain important information regarding the geochemical maturity, composition, weathering conditions and tectonic settings of both the provenance and associated depositional basins. The constituents of beach sands include various materials resistant to abrasion by waves, including silicates such as feldspar and quartz and shells and other derivatives of living organisms and lithic fragments. As described by Pettijohn et al. (1987), weathering, degradation and fragmentation of these materials leads to the particles that together form the sands. The geochemistry of clastic sediments can be effectively utilized in both

evaluating geochemical maturity and tectonic setting and determining provenance (Bhatia, 1983; Roser & Korsch, 1986; Roser & Korsch, 1988; Condie *et al.*, 1992). Pettijohn *et al.* (1972) first discussed the concept of maturity in sediment, suggesting that maturity should be assessed using a QFL diagram. A very mature sand, for example, may primarily comprise quartz. However, this method is limited by there being in excess of 50 sand and sandstone classification systems, laboratory analysis is required with point counting (typically of 500 points) and petrographic thin sections and carbonates are excluded. As an alternative, the X-ray fluorescence (XRF) analyses method was employed to this study.

The sands from six regions on the coasts of South West Japan were examined; these were the coasts of: Northern Kyushu, Shimane, the Noto Peninsula, the Tango Peninsula, Yamaguchi and Tottori. The chemical weathering intensity was investigated through the characterization of sands from each of these locations using three indices commonly used for describing weather profiles – the Chemical Index of Alteration (CIA), the Plagioclase Index of Alteration (PIA) and the Chemical Index of Weathering (CIW). For each sample, these indices combine the bulk of major element oxide chemistry into a single value. For further analysis, a variety of triangular and scatter plots were constructed from the geochemical data obtained. Additionally, the geochemical composition data was compared with that of four other sources. The composition of 15 local river sediments was obtained from the Geological Survey of Japan and the National Institute of Advanced Industrial Science and Technology (GSJ & AIST, 2013a). That of 15 near-shore marine sediments from around Yamaguchi was obtained from the same source (GSJ & AIST, 2013b). The remaining two sources were both upper crust data – the average upper continental crust (UCC) data (Rudnick & Gao, 2005) and the average upper crust of the Japanese Archipelago (UCJA) (Togashi *et al.*, 2000). By their very nature, local river and near-shore marine sediments broadly reflect the composition of the beach sands present in their drainage basins; studying the geochemical compositions of the sediments provides valuable baseline data for use in many geological and environmental fields. However, the chemical compositions of local river and near-shore marine sediments do not necessarily directly reflect those of their source rocks. Significant differences between the source and sediment compositions may result from factors such as the extent of source area weathering, sorting and average grain size, localized heavy mineral concentration, and alluvial storage or flushing of fine material.

The purpose of this study is to consider foreshore sampling and examine pocket beaches. This is important because the geochemical maturity of sands and biogenic production can be evaluated using silica, aluminium and calcium from major element X-ray fluorescence (XRF) analyses. The objective of this study is to conduct a systematic evaluation of geochemical maturity and weathering and to provide insight into the source area paleo-weathering conditions in beach sands of the six coastlines of interest of South West Japan. These factors are evaluated in this study using elemental abundances, weathering indices and principal ratios in comparison to the mean UCC and mean UCJA, as estimated from the representative surface rocks. In order to evaluate the biogenic productivity in Yamaguchi coast, the climatic conditions and water quality in Yamaguchi and Kyushu is observed. The aims of this study are twofold. The first aim is to present new data that is obtained by XRF. The second is to describe the characteristics of the four series of data previously mentioned – that is, the 15 local river sediments (GSJ & AIST, 2013a), the 15 near-short marine sediments (GSJ & AIST, 2013b), the average UCC (Rudnick & Gao, 2005) and the average UCJA (Togashi *et al.*, 2000) in terms of the general relationships between their geochemical composition and the abundance of elements contained within them.

1.1. Study area

The Japanese Islands have complex coastal landforms, where mountains and hills meet the sea in composite geometries. Japan has an extensive coastline of approximately 35,000 km in length, which contains a large number of beaches, among which are a significant number of pocket beaches concentrated along the coastline of South West of the country. From a tectonic point of view, South West Japan is generally subdivided into an Inner Zone (Asian Continent side) and an Outer Zone (Pacific Ocean side) that are separated by the Median Tectonic Line (MTL) (Figure 1). The Outer and Inner Zones have broadly similar geological structures, both comprising stacks of flat-lying tectonic layers that become progressively younger with increasing depth from the surface. In contrast to the Outer Zone, the Inner Zone contains Cretaceous to Palaeogene subduction-related, mainly granitic volcanic – plutonic complexes (Takagi, 2003; Nakajima *et al.*, 2004), locally associated with high-temperature, low-pressure "Ryoke" metamorphism (Nakajima 1997; Brown 1998; De-Jong *et al.*, 2008). Pocket

beaches are common along the coastline of South West Japan's Inner Zone. This is broadly coincident with the southwestern mountain arc.

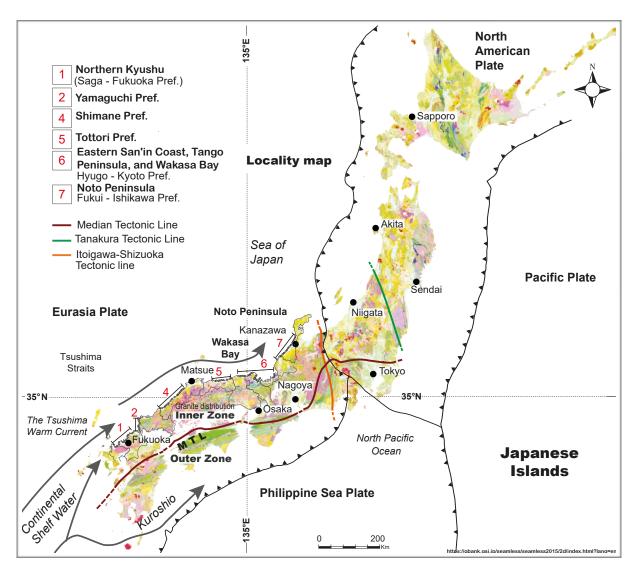


Figure 1. Geological map showing water circulation systems of the Sea of Japan, and geotectonic subdivision of the Japanese Island. Modified from the Geological Survey of Japan (GSJ) and National Institute of Advanced Industrial Science and Technology (AIST), 2016.

It also tends to follow the common pattern of the low- and highlands that curve around towards the Sea of Japan. This is in clear contrast to the beaches on the Outer Zone, which is characterized by an upper crust occupied with the Shimanto accretionary complex, gently dipping northward and with an extremely thin lower crust. The Shimanto and younger accretionary complexes intrude into the inner zone. The Sambagawa metamorphic rocks are distributed along the MTL, traced to a depth of about 20 km. Cretaceous and Paleogene granitoids are widely distributed in the San' in district of South West Japan. Locally, these granitoids are known as the Daito granodiorite, which consists of medium-to-coarse-grained hornblende-biotite granodiorite of the magnetite series (Ishihara, 1977). Tertiary sedimentary and volcanic complexes are distributed in sedimentary basins in the north of the district along the coast of the Japan Sea. Lower Miocene non-marine sediments unconformably overlie the basement granitoids; this may be indicative of Early Miocene paleo-weathering. The Miocene basins mainly developed along the Japan Sea coast during its opening, but small basins also occur sporadically in limited areas in the mountain regions.

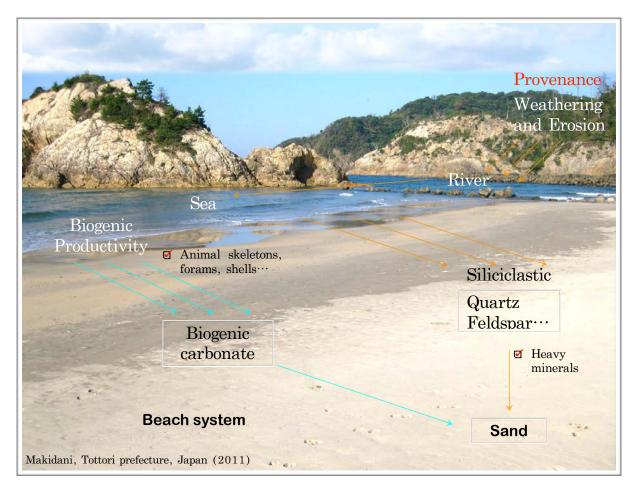


Figure 2. Genesis of beach sands. Siliciclastic rocks supply mostly quartz and feldspar, and rock fragments, biogenic carbonate mostly animal skeletons, forms, shells….

The Sea of Japan, one of the largest marginal seas of the Western Pacific Ocean, is located along the edge of the Eurasian continent and is partially separated from the unclosed ocean by Japan's islands (Gamo & Horibe, 1983; Danchenkov et al., 2006; Talley et al., 2006; Inoue et al., 2007). It is connected to the open Pacific Ocean through the Tsushima strait in the south and the Tsugaru, Soya and Mamiya Straits in the North. The Tsushima current, a warm current providing significant nutrients and heat as well as a means of transportation for marine organisms in the Sea of Japan, comprises three branch currents, one of which travels north-eastwards along the San' in coast (Inoue et al., 2007) towards Yamaguchi. Therefore, is reasonable to expect that not only did the Tsushima current play a significant role shaping the ecosystem, climate and environment of the Sea of Japan and its surrounding coastline within the Quaternary period (Kitamura et al., 1997), but also it continues to contribute to the flourishing biogenic productivity in the Yamauchi coastal region. The annual temperature range is 12-24° C (the average being 14° C), while the total rainfall ranges from 1,200 to 1,500 mm per year, with heaviest falls occurring in June and July.

1. 2. Pocket beach characteristics

Sand originates mainly from the weathering and erosion of land and is transported to the sea by river systems. The most common beach materials are quartz and feldspar (both siliciclastic rocks). Additionally, biogenic materials from the sea such as animal skeletons, Foraminifera and shells also produce sand.

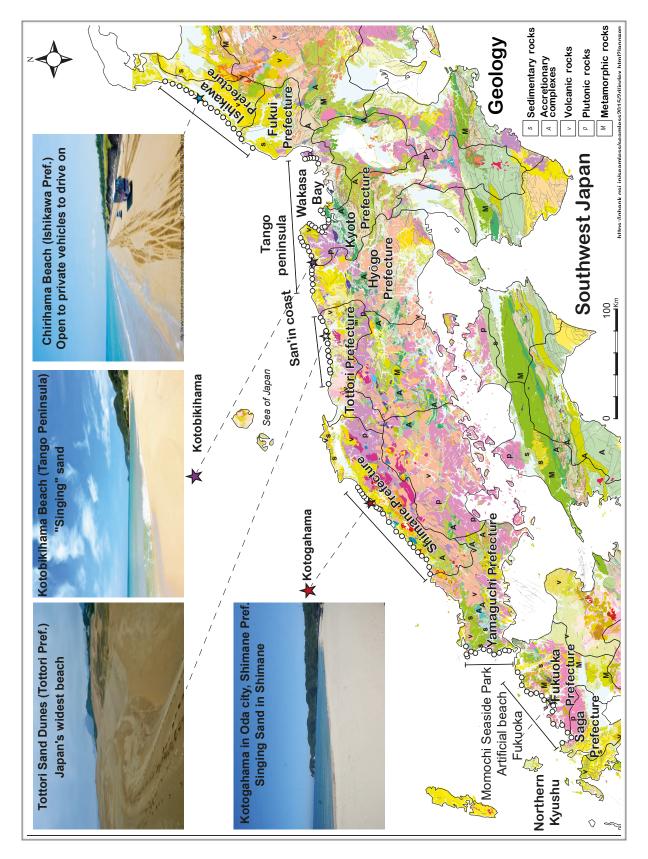


Figure 3. Detailed geological map of the coasts of South West Japan, showing the location of beaches sampled on the coastline of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, Wakasa Bay, and Noto Peninsula. Modified from, the Geological Survey of Japan (GSJ) and National Institute of Advanced Industrial Science and Technology (AIST), 2016.

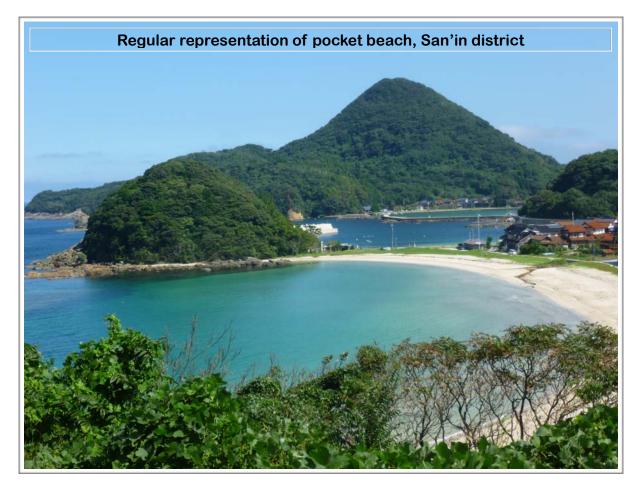


Figure 4. Regular representation of pocket beach, San' in district.

Biogenic carbonates of marine origin become important when the climatic conditions and the seawater qualities are good (Figure 2). It is worthwhile for someone wishing to investigate beach sands to discuss at least one of the locations identified in Figure 3, each of which has an unusual characteristic. For example, Momochi in Fukuoka is a man-made beach developed on reclaimed land. Chirihama beach, in Ishikawa Prefecture, has a driveway 8 km in length and 50 m wide, which is open to private vehicles, while the Tottori Sand Dunes within the San' in Kaigan National Park are a famous tourist attraction on account of being the country' s largest. Finally, three beaches are noted for the curious phenomenon of 'singing sand' ; the name of Kotobikihama Beach (on the Tango Peninsula) is derived from this, while there is a legend associated with it on Katogahama beach in Oda city.



Figure 5. Terms describing an ideal beach profile. Note the wave-formed ripples by foreshore tidal flat on the Nijo beach, Fukuoka Prefecture, Japan (2012).

Beaches are environments where loose sediments such as sand, gravel and cobbles are regulated by oceanic processes. Figure 4 is a typical representation of a pocket beach, a form of beach particularly common throughout Japan, especially in the southwestern part. Figure 5 shows a typical beach profile from foreshore to backshore; many of the investigated beaches have such profiles. Wave-formed ripples by foreshore tidal flat are observed on the Nijo beach. The first part is the foreshore, the part of the seashore that slopes from the low tide mark toward the crest of the berm. Next is the backshore bar, which protects beaches from erosion. Finally, sand dunes may form in the backshore environment through wind action. Based on field observations and measurements, the types of beaches present in the area reflect the processes from which the beaches formed and demonstrate characteristics of beach sands.

The beaches investigated can be classified into long beaches or pocket beaches according to the characteristics of their shapes using the length (L) and the arc length (l) (Figure 6). Pocket beaches are small (usually no more than a hundred meters) and are generally found between headlands. Small-pocket beaches are typically composed of sand and floating material such as algae, and provide isolated habitats for a variety of plants and animals. Beaches play an important role in protecting the coast; their loss negatively influences human activities as well as the environment.

The approximate arc lengths of the Makitani and Higashihama pocket beaches (Tottori Prefecture) are 1.48 km and 1.53 km. The Makitani pocket beach and the Hōjō long beach are illustrated in Figure 6. The average length of all the beaches investigated in Yamaguchi Prefecture is 1.13 km (Table 1), indicating they are relatively pocket beaches. The average length of beaches in the Shimane Prefecture is 1.73 km, three significant ones being Tinoza, Hashi 1 and Kuromatsu 3 with elongated seating areas of lengths 4.55 km, 3.80 km and 0.85 km respectively (Table 1). However, on the 82-km section of the Prefecture's coast between Masuda and Ohda, the beaches are essentially indistinguishable. At 12.5 km, the longest is Mochiishi, although this comprises

seven 'sub-beaches'. The coastline of Tottori Prefecture is some 129 km in length. The longest of the beaches is that at Houzyou, which stretches some 19.5 km, while the shortest is that at Ishiwaki (0.96 km). The average length of all the beaches investigated is 5.79 km. Among the ten beaches sampled, four are long (Houzyou, Hakuto, Karo and Sakyuuhigashi); three are of medium length (Tomari, Anedomaria, and Hamamura), and two (Makitani and Higashihama) are classical pocket beaches. The tenth beach was Ishiwaki; although this was undoubtedly very small, its characteristics were such that it could not be classified as a pocket beach.

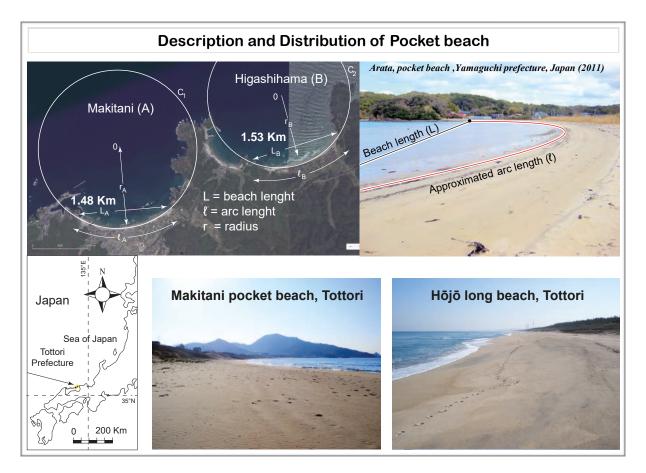


Figure 6. Shape of the Makitani and Higashihama pocket beaches in Tottori, Japan: (L) beach length, (l) arc length of the beach, and (r) radius of the approximated circle.

Sites	L (km) {	l (km)	Sites	L (km) l	(km)	Sites	L(km){	2 (km)
Yamaguchi prefecture	Э		Shimane prefecture			Tottori Prefecture		
Ayaragi + Yasuoka	2.40	2.50	Nakasu	0.15	0.18	Houzyou-1 + Houzyou-2	19.50	19.60
Fukue	2.20		Mochiishi	12.50	1.89	Tomari	2.91	2.98
Yoshimi	0.60		Araiso	0.69	0.71	Ishiwaki	0.96	0.98
Toyoura	2.20	3.00	Kitahama	0.28	0.30	Anedomari+Hamamura	5.80	6.13
Yoshimo	0.95		Tanoura	0.30	0.38	Hakuto + Karo	6.75	6.80
Doigahama	1.00	1.20	Orii	0.63	0.65	Sakyuuhigashi	7.50	7.58
Hinaka	0.25	0.35	Kuromatsu	0.85	0.90	Makitani	1.44 1.44	1.48
Akada	0.40	0.55	Nishihamada	1.43	1.78	Higashihama Average		1.53
Tunoshima	0.67	0.72	Shimokou	1.10	1.15	Average	5.79	5.89
Agawa	0.62	0.92	Hashi	3.80	3.98			
Average	1.13	1.32	Tunozu	4.55	4.80			
			Gohtsu	0.95	0.98			
			Asari	2.39	2.53			
			Iwamifukumitsu	0.74	0.78			
			Yusato	0.28	0.28			
			Kotogahama	1.25	1.45			
			Nima	0.50	0.53			
			Ohura	1.33	1.38			
			Isotake	1.25	1.30			
			Uozu	0.11	0.11			
			Shizuma	1.35	1.40			
			Average	1.73	1.31			

Table 1. Shape and characteristics of beaches on the coastline of Yamaguchi, Shimane and Tottori, South West Japan. (L= length of beach, and l = arc length of beach).

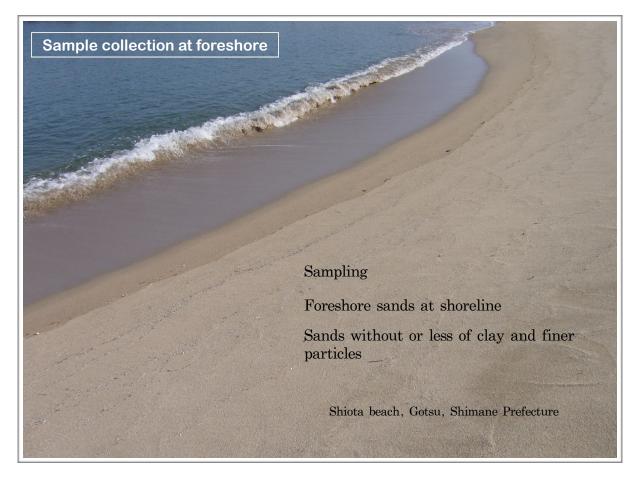


Figure 7. Foreshore sampling of beach sands at shoreline, Sands without or less of clay and finer particles.

Chapter Two

2. MATERIALS AND METHODS

2. 1. Sample collection at foreshore

Sampling sites were selected based on the accessibility and character of sites. Tidal information, obtained from the Japan Meteorological Agency, was used to ensure sampling was performed at low tide or moderate tide times. Beach sand samples were taken from the uppermost few centimetres of the beach; the location of sampling on each beach was chosen such that the clay content and fine particle content was as low as possible, and the coarse sands content was as high as possible, as illustrated in Figure 7. For each sample, approximately 200 grams of sand were collected using a stainless-steel scoop from the foreshore of selected sites along the coasts of South West Japan on the coastlines of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula and Noto Peninsula. In some sites, the extent of sea walls of artificial structures constructed for coastline protection by reducing the rate of erosion prevented sample collection. Samples were collected from the foreshore surface and the location from which the samples were obtained recorded. Beach widths were measured using a linen tape, and beach slope measured with an inclinometer. The samples were stored in their natural state (that is, wet from seawater) in the laboratory before further processing. The experimental workflow for the sample preparation and data analysis are illustrated in Figure 8.

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2. 2. Sample preparation and analysis

In the laboratory, approximately one-third of each sample was transferred to Pyrex beakers in their natural state, covered with aluminium foil to allow air circulation, and dried in an oven at 110° C for 24 hours. Once dried, subsamples of the sediments were crushed in an automatic agate mortar and pestle grinder to produce a powder suitable for analysis. Fused glass discs and pressed powder briquettes were prepared from the crushed samples for major oxide and trace element analysis, respectively. In order to determine the Loss on Ignition (LOI), 5.000 \pm 0.001 g of the dried powder sample were transferred to porcelain crucibles. The samples were ignited for at least two hours in a muffle furnace at 1050° C and the weight differential reported as a percentage loss.

The ignited material was then manually disaggregated and re-crushed in an agate pestle and mortar, and returned to a 110° C oven for at least 24 hours. The fused glass discs were prepared in an NT-2000 automatic bead sampler using the ignited material in addition to an alkali flux comprising 80% lithium tetraborate and 20% lithium metaborate, with a sample:flux ratio of 1:2. Analytical methods, instrumental conditions and calibration followed those described by Kimura and Yamada (1996). The pressed powder briquettes were prepared by pressing about 5 g of powdered sample into 40 mm diameter plastic rings, using a force of 200 kN for about 60 s in an automatic pellet press (E-30 T.M Maekawa) following the Ogasawara (1987) method. Average errors for all elements were less than $\pm 10\%$ relative. Analytical results for GSJ standard JSl-1 were acceptable compared to the proposed values of Imai *et al.* (1996).

Major elements expressed as oxides (SiO₂, TiO₂, Al₂O₃, Fe₂O₃*, MnO, MgO, CaO, Na₂O, K₂O, and P₂O₅) and 18 trace elements (As, Pb, Zn, Cu, TS, Ni, Cr, V, Sr, Y, Nb, Zr, Th, Sc, F, Br, I, and Cl) were obtained using an automated RIX 2000 system (Rigaku Denki Co. Ltd.) at Shimane University. The composition of the sand in terms of particle size was investigated using the Shimadzu SALD-3000S Laser Diffraction Particle Size Analyzer; microscopic observation of the sand was performed on sand holders.

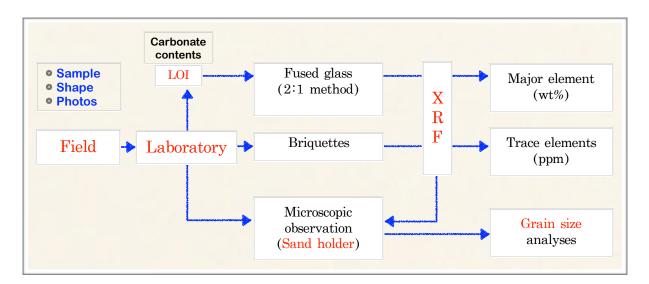


Figure 8: Experimental workflow for the sample preparation and data analysis.

Chapter Three

3. RESULTS

3. 1. Major and Trace Elements Geochemistry

Beach sand collected along the coasts of Southwest Japan, on the coastlines of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula and Noto Peninsula comprise siliciclastic (quartz and feldspar) and biogenic deposits. The sands are representative of the characteristics of their constituent materials. These materials originate from a variety of sources and processes some are transported to the beaches by marine currents from external sources or by rivers from nearby land, some are produced by erosion of the shoreline in the vicinity of the beach, some are created in situ by living organisms while others are related to The abundance of the major elements determined from the human activity. beach sand samples from the six coastal sites under investigation are summarized in Table 8. The corresponding LOI values are included in each Table for each location. The average geochemical composition of 15 local river sediments from AIST & GSJ (2013a), 15 near-shore marine sediments along the San' in district from AIST & GSJ (2013b), UCC (Rudnick & Gao, 2005) and UCJA (Togashi & Imai, 2000) estimated from the representative surface rocks and the means of major elements and trace elements are included in Table 8 for comparison.

3. 1. 1. Northern Kyushu

XRF major and trace element analyses of the beach sands from Northern Kyushu, Japan, are listed in Table 2 and Table 8. The beach sands had moderate to high SiO₂ contents, with abundances ranging from 54.43wt% to 91.23wt% (mean 77.24wt%); this is well above the 66.62 wt% present in the average Upper Continental Crust (UCC) reported by Rudnick and Gao (1995). The higher values in the beach sands reflected their quartz content. The next most abundant element, Al₂O₃, ranges from 4.71wt% to 18.35wt%, averaging (10.62wt%), less than in UCC (15.40wt%). In most samples, CaO contents are low (< 5wt%) and less than UCC (4.76wt%), reflecting low shell contents. Samples from Ashia and Munakata-1 are exceptions, with higher CaO contents of 13.47wt% and 28.43wt%, respectively. Among the remaining major elements K₂O (average 2.97wt%, range 1.63 - 4.64wt%), Na₂O (2.17wt%, range 0.87 - 3.63wt%) and Fe₂O₃* (1.35wt%, range 0.31 - 3.97wt%) are the next most abundant. Other major elements (MgO, TiO₂, MnO, and P₂O₅) are less abundant, and average values for all are less than in UCC.

Loss on ignition (LOI) data are presented in Table 2 to indicate variations in organic matter and calcium carbonate content of the beach sand sediment. Average LOI was low, averaging only 3.45wt%. However, the very high CaO contents at Ashiya and Munakata-1 identified in the XRF analysis was reflected in two the LOI readings for these two sites, which were significantly higher that this average figure, at 10.87wt% and 18.78wt% respectively. Table 2 also shows the concentration of trace elements in the beach sands. The two elements with the highest contents were chlorine (average concentration 2920ppm (range 44 to 10660ppm)) and sulphur (average concentration 871ppm (range 405 to 2260ppm)). The strontium content was significant, averaging 382ppm (range 76ppm to 989ppm), whereas iodine content varied from 3ppm to 3630ppm, averaging 145ppm. Fluorine content ranged from 11ppm to 340ppm, and zirconium from 8ppm to 67ppm. With average concentrations of 27ppm and 20ppm respectively, the vanadium and chromium contents were considerably lower than those seen in the UCC (97ppm and 92ppm respectively). Concentrations of other trace elements such as As, Pb, Zn, Cu, Ni, Y, Th, Sc, and Br were less than 20ppm on average, and below the abundances in UCC.

3. 1. 2. Yamaguchi Prefecture

Yamaguchi coastal beach sand samples had low to high SiO₂ contents, with abundances ranging from 4.72wt% to 92.16wt%, and averaging 61.20wt%, the high values reflecting their quartz content (Table 3 and Table 8). CaO was the next most abundant, the average value being 28.18wt% with a range of 0.77wt% to 87.37wt% (low values corresponding to high SiO₂ content and vice versa). The wide range of these values represented two situations – those samples with higher values were indicative of a significant biogenic CaCO₃ presence, but a low shell content was suggested by those with lower CaO values. Al₂O₃ was the next abundant element with (4.63wt%) average and ranges from 0.65wt% to 9.50wt%. Among the remaining major elements, K₂O (average 1.76wt%, range 0.34 - 4.40wt%), MgO (average 1.46wt%, range 0.10 -4.93wt%), Na₂O (average 1.31wt%, range 0.30-1.99wt%), and Fe₂O₃*

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(1.18wt%, range 0.01-2.66wt%) were the next most abundant. Other major elements (TiO₂, P_2O_5 , and MnO) were less abundant.

The Yamaguchi coastal beach sand samples demonstrated relatively low to high LOI contents with CaO contents (averaging 28.18wt%, ranging from 0.77 to 87.37wt%). Among the analysed trace elements, Cl had the highest content as a result of contaminations from seawater, averaging 12602ppm, and ranging from 40ppm to 54231ppm, followed by total sulphur (TS) averaging 2263ppm, with a range from 30ppm to 5557ppm. Sr contents were significant, averaging 695ppm and ranging from 44ppm to 1358ppm, whereas F contents varied from 11ppm to 274ppm, averaging 11ppm. Zr contents ranged from 41ppm to 144ppm and V from 3ppm to 43ppm. The average contents of I and Sc were 28ppm and 20ppm, respectively. Concentrations of other trace elements were less than 20ppm on average. The Yamaguchi beach sands contained small amounts of SiO_2 , Al_2O_3 , and $Fe_2O_3^*$, whereas CaO and LOI accounted for over 65% and 35% respectively and were relatively rich in organic matter. Of particular note were the samples from three beaches - Doigahama, Hinaka and Akada. These beaches' samples comprise carbonate or biogenic sands, composed primarily of foraminifera, ostracod, shells and sea urchins, the primary constituent of which is $CaCO_3$. The high LOI results were demonstrated by XRF analysis, confirming the high $CaCO_3$ content of the sands.

3.1.3. Shimane Prefecture

As expected, SiO_2 was the most abundant from the Shimane coastal beach sand samples, averaging 82.27wt%, with a range of 54.38-89.50wt%, followed by Al₂O₃ (average 8.77wt%, range 4.54-16.46wt%) (Table 4 and Table 8). Among the remainder, CaO (2.70wt%, range 0.45-35.35wt%), K₂O (2.16wt%, range 0.88-4.44wt%), Fe₂O₃* (1.66wt%, range 0.42-3.14wt%) and Na₂O (1.65wt%, range 0.87-2.97wt%), were the next most abundant on average. MgO (average 0.49wt%) and TiO₂ (average 0.24wt%) were present in small amounts, whereas MnO and P₂O₅ (both averaging 0.48 and 0.03wt%) were present only in trace amounts. In these samples, overall, LOI contents ranged from 0.42-22.59wt%, averaging 2.18wt%. Cl was the most abundant trace element (again a result of seawater contamination), with an average value of 2995ppm, and a maximum of 9959ppm. The TS values were significant, ranging from 275ppm to 3398ppm, with a mean value of 621ppm. Sr was the next most abundant, with a maximum of 1126ppm, a minimum of 77ppm, and an average value of 209ppm. Among the remaining trace elements, only F, Zr, Zn, Cr, V, As, and I contents were present in moderate concentrations, other trace elements, Pb, Cu, Ni, Y, Nb, Th, Sc and Br showing very low concentrations.

3. 1. 4. Tottori Prefecture

The samples from the Tottori coastal beach sands generally comprised silicate materials (for example, feldspar and quartz), the high quartz content being reflected in their high SiO_2 content (mean value 72.05wt%, range 66.30-82.23wt%). (See Tables 5 and 8.) There were also relatively high levels of Al_2O_3 present, albeit only in the order of one-fifth the level of SiO_2 (average 14.71wt%, range 10.05-17.35wt%). A relatively wide range of values about the average of 3.86wt% was seen for CaO content (0.84-7.49wt%), with a

significant shell material content being indicated by similar values from the LOI analysis (a range of 0.05-5.50 wt%, average 1.81 wt%). Na₂O and K₂O, also likely to be contained within feldspar, was less abundant, averaging 2.91 wt% and 2.62 wt%, respectively. Three major elements (Fe₂O₃*, MgO and TiO₂), were present only in minor amounts (averages 2.48 wt%, 1.02 wt%, and 0.26 wt% respectively), and P₂O₅ and MnO were present only in trace amounts (both averaging 0.05 wt%). Iodine (I) was the most abundant trace element averaging 3698 ppm, with a maximum of 7394 ppm. It was followed by total Chlorine (Cl), which averaged 578 ppm (range 342-1007 ppm), and Sr (average 384 ppm, range 131-598 ppm). Average concentrations of all other trace elements except TS (152 ppm) were less than 100 ppm, reflecting the high SiO₂ content and marked quartz dilution in this suite of sediments.

3. 1. 5. Tango Peninsula

Results showed that SiO₂, dominated the analysed sand samples averaging 78.02wt% (Eastern San' in coast sands), 81.02wt% (Tango Peninsula sands) and 84.83wt% (Wakasa Bay sands) (Tables 6 and 8). The other sites' samples showed relatively high SiO₂ content, with three exceptions. These three sites were Kirihama (average 69.06wt%), Shibayama (66.82wt%) and Takeno (41.89wt%) and, as would be expected given their high SiO₂ content, they exhibited high CaO contents of 13.77wt%, 11.21wt% and 43.35wt% respectively, as shown in Table 6. The high SiO₂ concentrations resulted in low contents of other elements, with Al₂O₃ contents of between 5.30 and 11.86wt% in the Eastern San' in coast sands, 7.60 to 12.44wt% in the Tango Peninsula

and 5.55 to 10.27wt% Wakasa Bay sands (Table 6). Average Al_2O_3 contents were 8.45wt% in the Eastern San' in coast sands, 9.92wt% for the Tango Peninsula sands and 7.95wt% in the Wakasa Bay sands (Table 8). Seven other elements were present in significant quantities; these were TiO₂, Fe₂O₃*, MnO, MgO, Na₂O, K₂O and P₂O₅. As shown in Table 6, their concentrations were, in the most part, less than 5wt% and, in some cases, less than 1wt%. Among the trace elements, the contents of ferromagnesian elements (Ni, Cr, V and Sc) and large cations (Y, Nb, Zr, Th and Sr) tended to be less abundant than they were in the UCC and the JAUC (Table 8).

3. 1. 6. Noto peninsula

Generally, beach sands from the Noto Peninsula were characterized by moderate contents of SiO₂ (75.44-83.43wt%, average 79.00wt%) and Al₂O₃ (8.13-13.02wt%, 11.39%). Furthermore, the Fe₂O₃* content was low (1.79-5.27wt%), as were those both of MgO (0.37-1.85wt%) and of TiO₂ (0.20-0.66wt%, 0.31wt%). These can be accounted for by the presence of a high level of quartz and a low level of mafic components (see Tables 7 and 8). The low CaO contents (0.57-1.83wt%, 1.27wt%) indicated that all the beach sand samples had very low carbonate components. The contents of Cr, Ni and Sc showed a wide range from 3 to 45ppm, detection limits to 9ppm, and 1 to 9ppm respectively, suggesting a contribution from more mafic components. Contents of the high field strength elements Th, Y, Nb, showed similarly wide ranges from 6 to 9ppm, 15 to 22ppm, and 5 to 8ppm respectively.

Table 2. XRF ma Alteration (CIA), F	Table 2. XRF major (wt%) and trace (ppm) element analyses (Alteration (CIA), Plagioclase Index of Alteration (PIA) and Chemic	e (ppm) Alterati) elen ion (P	nent a 'IA) an	analyse id Cher	is of b∈ nical In	of beach sands from norther cal Index of Weathering (CIW)	sands from northern of Weathering (CIW).	m nc ring (CIW).		Kyushu, Japan.	Japa	an. LC	LOI, 0	oven-dried loss	ied I	oss c	on igr	ignition;		and indices	es of		Chemical	Index	(of
SAMPLE	SiO ₂ TiO ₂ Al ₂ O ₃ I	Fe ₂ O ₃ *	MnO	MgO	0 CaO	Na ₂ O	K ₂ O P ₂ (D ₅ LOI	Ö	A PIA	CIW,	As F	Pb Z	Zn Cl	Ž	ç	>	Sr ≺	q N	Zr	ТЪ	Sc	ы М	ğ	-	Ö	
Fukuoka Prefectu	Prefecture (n=26))																							
Ashiya	67.12 0.29 9.64	3.11	0.08	3 1.45	5 13.47	2.09	2.70 0.06	06 10.8	37 24	1 19	26	23	14	26	ლ	÷	41	556 1	6 3	37	-	30 1		101 1		6197	97
Namitsu 1	88.25 0.10 6.21	0.99	0.02	0.36	3 1.50	0.94	1.63 0.0	01 5.58	8 51	1 51	59	20	13	5		.	36	423 1	4 2	46	N	21-1	1535 4	10 12	23	10660	900
Namitsu 2	75.17 0.18 9.98	2.21	0.04	0.91	6.49	¢.	2.77 0.04	04 1.87	7 35	31	39	15	12	14		53	14	181 1	12	46	N	9		35 5		1007	07
Munakata 1	54.43 0.19 8.28	1.73	0.04	1.84	1 28.43	1.96	3.00 0.08	08 18.7	78 13	8	1 3	÷	÷	18 6		2	17	989 1	4.		N	46 2				5232	32
Munakata 2	85.39 0.09 5.76	0.92	0.01	0.32	2 4.62	1.03	1.81 0.03	33 3.58	8 32		36	17	12	13		2	0	337	9	46	\sim	13 7	724 2	\sim			
Katsuurahama	89.25 0.10 5.81	0.72	0.01	0.25	0.61	1.16	2.07 0.01	0.98	8 53	3 54	99	œ	12	13		21		115 1	0 3	53	ო		742 1	11 3(• •	\sim	
Tsuyazaki	80.87 0.13 9.34	1.13	0.02	0.39	3.53	-	2.61 0.02	02 2.58		3 40	49	12	13	16	2		16	365 1	3 2	54	ო	11	665 2	227 4	24	•	
Miyaji	77.56 0.16 9.25	1.46	0.03	0.65	6.13	2.01	2.70 0.04	04 5.02			39	13	12	19 6		7		433 1	3 2	47	ო	-					39
Koga	84.03 0.13 9.11	0.98	0.02	0	1.02		2.67 0.02	0.93 0.93		58	67	10	15		10	8		241 1	54	55	N						
Singu 1	81.41 0.16 9.67	1.18	0.02	0.52	2 1.59	1.82	3.61 0.02	02 1.22	2 50		62	10	16	17	ლ	25	15	247 1	15 3	54	N	2	530 6	61 5			
Singu 2	82.32 0.09 10.98	0.72	0.01		0.96		2.79 0.01	11 1.25			70	÷	16			25	17	300 1	16 2	45	N						62
Wajiro	91.23 0.09 4.71	0.50	0.01	0.26	0.29		2.05 0.00	0			71	ω	12		2	21		62	о Э	46	-						21
Ikinomatsubara	86.74 0.03 7.40	0.41	0.01	0.11	0.68		3.44 0.0	01 0.55		2 54	70	С	18	12 (N			13 2	56	-	۲					
Imajuku 1	82.40 0.04 10.67	0.31	0.01	0.13	3 0.60		3.86 0.01	01 0.58			71	ო	19			7		231 1	16 2	53	-						84
Imajuku 2	76.67 0.09 13.38	0.76	0.02	0.31	1.15		4.64 0.02				99	с С	21		N	N	•		7 3	60	-						95
Nagahama 1	68.12 0.17 13.65	1.33	0.03	0.00	8.99		3.64 0.05		2 35		39	9		18	∞	13	19	654 1	14 1	25	N	22 9		143 6		44	4
Nagahama 2	74.48 0.16 12.92	1.30	0.03	0.91	4.04		3.26 0.04	04 2.45			52	9	13 13		6 15	38	26	464 1	14 2	39	-						8
Itoshima 1	72.93 0.15 10.19	1.33	0.03	0.76	9.04		3.39 0.04		9 30) 25	34	~			9	19	19	557 1	4 2	38	\sim						17
Itoshima 2	79.66 0.11 10.56	0.85	0.02	0.44	1 3.70		2.85 0.02	02 3.92			52	o	15		2	33	23	434 1	16 3	45	\sim						77
Itoshima 3	77.84 0.10 10.10	0.81	0.02	0.48	3 4.79		3.97 0.03	33 3.70			46	0	16	13	\\	6	2	410 1	15 2	45	с		805				97
Nijo 1	75.72 0.27 12.47	2.29	0.05	1.14	1 2.90	2.65	2.46 0.04	04 1.13			57	4	13	23	7 15			316 1	14 5	58	N			340 9			12
Nijo 2	71.23 0.29 15.39	2.34	0.04	1.19	3.36	3.19	2.94 0.02	02 1.28	8 52	2 52	58	4	15	24	9 14		62	386 1	4	48	N	.					õ
Nijo 3	81.01 0.11 11.05	0.91	0.02		1.47	2.26	2.67 0.02	02 0.84		1 56	64	4	17	15	10	14	10	223 1	э Э	55	N		652	Ť.			05
Nijo 4	75.72 0.16 13.21	1.28	0.02		2.92	2.55	3.49 0.04	04 1.34	4 50) 50	58	ω	15	16 6	ю Э	20	1 3	354 1	4 3	49	-		503	ى			
Nijo 5	79.20 0.09 11.58	0.73	0.01	0.32	2.64	2.23	3.17 0.02	02 2.26	6 49	9 49	58	ω	17	15	4	19	10	365 1	63	41	N		662	ى		693	33
Shikaka	76.50 0.17 12.54	1.42	0.03	0.74	1 2.73	2.69	3.16 0.02	02 1.29	9 50) 49	57	4	15	20	7 9	26	31	359 1	15 4	44	\sim	8	568 3	31 6			õ
Saga Prefecture (n=4)	(n=4)																										
Karatsu 1	64.77 0.52 18.35	3.97	0.06	1.63	3 4.03	3.54	3.09 0.03	03 2.73		3 53	58	ŝ			_			419 1	15 6	54	N		788 5	59 8		1 2737	37
Karatsu 2	73.96 0.23 14.31	1.70	0.04	0.69	9 2.71	3.08	3.25 0.02	02 1.40		2 52	59	ω		22	6 3		38	369 1	13 4	49	N	о С		103 6	26		
Karatsu West 1	74.69 0.14 14.10	1.14	0.02	0.46	3 2.51	3.68	3.24 0.03	03 1.50	0 50		57	ω	15			17			13 3	56	N		665		36		85
Karatsu West 2	68.46 0.22 8.00	2.02	0.03	1.30	15.91	1.70	Ö.	07 11.7	7 19	9 15	20	13				20		816 1	2	ω	ო		2081	∞	т	3674	74

Table 3. XRF maji	Table 3. XRF major (wt%) and trace (ppm) element analyses of beach sands from Yamaguchi Prefecture, Japan. LOI, oven-dried loss on ignition; and indices of CIA, PIA and CIW) (mqq)	eleme	ent ana	Ilyses o	f beac	sh sanc	ds from	ı Yam	aguch	i Pref	fectur	e, Ja	pan.	LOI,	oven	-drie	d lose	s on i	gnitio	n; an	id ind	lices o	of CIA	, PIA	and	CIW.	
SAMPLE	SiO ₂ TiO ₂ Al ₂ O ₃ Fe ₂ O ₃ * MnO	Fe ₂ O ₃	* MnC	O MgO	O CaO	Z	$0 K_2 O$	a20 K20 P205 L01	ō	CIA PIA		CIW As	Рр	Zn	C	ï	C /	V Sr	≻ 、	qN	Zr	₽	S	ЦS	F Br		Ö	
Toyoura 1	84.52 0.05 6.84	0.52	0.01	1 0.21	1 3.04	1.51	3.27	0.02	2.80	37 2	29 46	6 7	17	12	N	9	10	1	55 18	-	55	С	-	745	-	11 32	2 3480	8
Toyoura 2	84.00 0.03 7.79	0.36	0.01	0.13	3 2.27	1.11	4.29	0.02	2.15	42	35 57	7 7	18	2	-	2	10	÷	110 24	0	50	С		499	235	5 30		
Kawatana	73.73 0.09 9.50	0.77	0.02	0.46	6 9.0Z	1.97	4.40	0.03	7.18	28	19 33	3 5	18	16	က	-	7	е Е	386 23	3	73	2	8	1113		1 21	4804	4
Kogushi	84.42 0.09 7.14	0.01	0.20	20 0.20	3.14	1.39	3.39	0.01	2.75	38	30 47	7 4	17	15	-	Ю	18	÷	150 20	-	83	2	с С	735		1 26	3 10049	149
Doigahama	4.72 0.06 0.65	0.45	0.03	3 4.93	3 87.37	1.31	0.35	0.12	41.12			-	7	-	4		2	. 13	1348 7	•			39 5	5557	62 1	15 .	18114	14
Arata	57.40 0.23 5.73	1.64	0.03	3 1.47	7 29.85	1.77	1.81	0.09 1	19.35			7	÷	23	Ю		10	96	954 12			-	24 2	2517	78 1	14 .	11784	84
Kanda	31.39 0.14 2.73	0.99	0.02	02 2.76	5 59.35 5	1.40	1.05	0.16	23.66			4	ω	ω	С		10	. 13	1354 8	•	•	•	37 2	4106	145 1	14 .	13222	22
Tunoshima 1	46.62 0.11 2.01	0.67	0.02	1.23	3 47.69	0.86	0.65	0.12	26.44			0	ω	Ð	4		ω	12	1254 5			-	33 9	3817	24	6	8941	41
Tunoshima 2	37.30 0.24 3.03	1.35	0.03	3 1.46	5 54.86	0.97	0.63	0.13	28.82			0	~	8	2		13	. 13	1316 5		•	-	35 3	3645	181	6	6715	15
Agawa	48.29 0.19 5.01	1.77	0.03	3 1.83	3 39.92	1.33	1.51	0.12	24.08			б	6	19	9		6	÷.	1143 1-		•	\sim	30 2	2789		10 .	5094	94
Hinaka	18.90 0.18 3.43	1.79	0.05	11 4.11	1 68.69	1.89	0.83	0.14	35.49			2	7	16	2		ω	. 13	1336 10		•	-	35 4	4023	89 1	13 .	54231	31
Akada	21.36 0.11 2.62	0.98	0.03	3 3.50	0 68.71	1.69	0.85	0.14	35.01			4	ω	9	4	·	16	. 13	1358 8	•	·		38 4	4693	96 1	15 .	180	0
Ushirohama A	84.68 0.04 7.25	0.38	0.01	1 0.17	7 2.43	1.24	3.79	0.01	2.68	41	33 53	3 5	17	∞	-	Ð	18	÷.	20 21	-	53	Ю	-	. 086	115 1	14 30	6568	80
Ushirohama B	89.28 0.03 5.63	0.33	0.01	11 0.11	1 0.77	0.94	2.88	0.01	1.56	48	46 66	6 5	14	7	-	£	15	∞.	85 19	1	51	С		422	114	4 31		
Narabimatsu	87.89 0.05 5.94	0.36	0.01	01 0.10	0 1.74	0.83	3.07	0.01	1.30	43	37 57	7 3	16	8	13	ω	1 3	4	44 18	თ	41	4		744	116 1	12 35	40	0
Yasuoka	81.64 0.22 4.85	1.25	0.02	0.60	0 8.16	1.44	1.78	0.05	6.52	20	15 22	2 7	14	. 32	2	2	32	ස ස	335 12		77	\sim	101	1414		9 22	26640	40
Ayaragi	88.90 0.11 4.28	0.79	0.02	0.29	9 2.95	06.0	1.73	0.02	2.93	33	26 39	9 6	12	18	\sim	6	19	- -	142 1-	-	55	\sim	4	1010		14 36	3 7541	41
Yoshimi A	87.24 0.32 3.41	1.42	0.03	0.44	4 5.62	0.86	0.62	0.04	4.67	22	19 23	34	12	27	4	ω	40	3 25	255 7	2	79	\sim	9	1061	76 1	12 27	6361	5
Yoshimi B	87.07 0.31 3.72	1.39	0.03	3 0.45	5 5.38	0.97	0.64	0.04	4.59	24 2	21 25	5 5	÷	26	N	6	37	3 25	258 8	N	78	\sim	8	1031	137 1	11 28	\$ 5297	97
Yoshimi C	80.54 0.90 4.49	2.66	0.04	0.66	<u></u> 3 8.86	0.94	0.85	0.05	6.79			·		34	Ð	9	59 4	43 38	383 9	ß	144	с	13.1	1212	124	8 16	3748	48
Yoshimi D	76.95 0.76 4.90	2.51	0.04	0.83	3 11.87	1.14	0.93	0.06	8.96			·	12	34	4	\sim	50 2	28 477	77 10	4	103	\sim	14	1577	11	11 19	6211	÷
Yoshimo	42.39 0.20 3.83	1.40	0.04	04 2.31	1 47.17	1.51	1.01	0.13	26.23			·	6	20	က	9	16	12	1213 9			-	33 3	3652	55	∞	8016	16
Arata A	32.29 0.18 3.90	1.66	0.04	04 3.14	4 55.41	1.94	1.33	0.13	29.15			·	10	20	9		20	÷.	1197 13	თ	•	\sim	37 2	4735 2	274 1	12 .	9163	33
Arata B	43.88 0.18 4.39	1.53	0.04	04 2.47	7 44.25	1.69	1.45	0.11	26.21			·	0	20	5		12	10	1083 12		•	\sim	33 3	3917	-	18 .	17912	12
Arata C	28.29 0.16 3.71	1.69	0.04	04 3.59	9 59.23	1.99	1.16	0.14	32.35			·	0	20	9		2	. 12	1238 12		•	\sim	37 4	4518	89	20 .	46472	.72
Arata D	56.58 0.19 5.94	1.68	0.03	3 1.55	5 30.47	1.63	1.83	0.10	19.39			·	10	24	4		16	. 97	970 12		•	\sim	26	30	14	16 .	21873	73
Fukue	92.16 0.38 2.27	1.61	0.03	3 0.29	9 2.60	0.30	0.34	0.02	1.86	29	27 31	4	÷	15	ი	Q	32	14 9	97 6		73	\sim	-	561	128	5 34		
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Shimane Prefecture,	ō	0.64	1.24	1.46	1.01	1.35	0.75	1.08		2.35	1.12	0.86	0.80	0.64	0.51		0.97	0.73	0.42		1.0.1		107	1 44	2.03	9	1.43	2.45	5.05	2.59			145	1.72	1.25	1.36	1.62	1.45	1.64	1.00	2.60
sands from	205 I	05	36	.05	.05	.02	-01	90.	.05	.04	.02	.02	.03	.02	.02	02	6	10.0	2 C.	05				2; F	05	.05	.03	.04	.03	10	5.0		80	02	.05	.05	.04	.02	80. 0	30	02
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ses o	CaO	2.43	1.70	3.05	2.99	1.77	0.74	1.16	2.21	1.95	1.01	1.02	1.37	1.10	0.95	0.86	0.45	0.56	0.69	0.77	0.77	10.1	- С С С С С С С С С С С С С С С С С С С	14.36	3.65	6.23	1.73	3.02	4.61	35.35	00.1	000	0.00	1.20	0.96	1.06	1.09	1.21	0.87	0.98	1.75
nalys	MgO	0.90	.47	.52	.77	0.16	0.10	0.44	.39	.64	.34	0.32	.52	.39	.26	.21	0.16	0.26	5.0	0.24		0.10	40.0	240	28	.97	.75	.83		1.57			040	0.29	0.40	0.40	0.34	0.14	0.39	0.49	.27
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þm)	Fe ₂ O ₃ *	3.10 1 05	1.30	1.59	2.68	0.70	0.42	1.91	1.62	2.13	1.26	1.24	2.93	1.80	1.31	0.92	0.58	1.25	2.64	0.85	0.84	- 14 0 14	0 9	28.0	3.14	2.16	3.06	2.57	1.16	1.41	1 7 00	1 82	89.0	1.21	1.82	1.79	1.50	0.64	1.64		1.20
d) əc	AI ₂ O ₃ F	29	75	42	16	N	4	22	74	46	84	6	-	4	<u>-</u>	-	E		о С	20	4 0	N T	- 4 7 0	2 0	2	2	56	88	4	000	9.40 8.60	να	2 9	2	96	33	90		с С	03 03	45
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) and	TiO ₂	0.35		0.17	0.30	0.07	<u>0</u> .0	0.21	0	0.26	0.17	0.16	0.53	0.31	0.20	0.12	0.08	0.27	\neg	0.12		0.09	0.10	0.40	0.41	0.25	0.64	0.39	0.21	0.15		0.00	0 19	0.11	Ö	0.19	0.15	0.06	00	0.41	0
(wt%	SiO ₂	77.71 76.05	0.90 81.51	75.56	76.63	33.49	89.50	34.23	31.05	1.07	30.02	82.63	80.46	34.08	33.58	32.70	37.11	1 80 1 80 1 80 1 80 1 80 1 80 1 80 1 80	87.74 2001	36.87	00.00 10.00	02.4%	7 50	70.66	80.17 80.17	6.49	36.55	34.96	37.29	54.38	22.U/	20.20	501	85.37	85.06	35.55	86.30	87.45	87.05	85.07	85.91
lajor			- 00	~							ω	ω	ω	ω	ω	ω	ω	ω	0	ω	υc	00		- ^	- 00		ω	ω	ω	<u>ц)</u> (υα	σ	σ	νœ	ω	ω	ω	ω	ωο	υα	ω
Table 4. XRF major (wt%) and trace (ppm) element analyses							ma 2			wamifukumitsu	-	2							g					ά	5						ר ס ס		- 0	1 က	4	5	9				
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able	SAMPL	Shizuma	sotake	Ohura	Nima	Kotogahama	Kotogahama	Kotogahama	Yusato	vam	Kuromatsu	Kuromatsu	Kuromatsu	Asari 1	Asari 2	Asari	Gohtsu	Iunozu	Ukinohama	Hashi 1	Hashi 2	Hashi 3	Chimokou	Nishihamada	Orii 1	Orii 2	lanoura	Tanoura 2	Kitahama	Araiso	warnilsuda	Mochilehi	Mochiishi	Mochiishi	Mochiish	Mochiishi	Mochiish	Mochiisl	Nakasu	Nakasu Nakasu	Kohama
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Table 5. XRF m	Table 5. XRF major (wt%) and trace (ppm) element analyses of be	(mqq) əc) eleme	nt anal	yses of	^t beach	ach sands from Tottori Prefecture, Japan. LOI, oven-dried loss on ignition; and CIA, PIA and	om Tott	ori Pr	efectur	e, Jap	an. L	ol, o	ven-c	dried I	o sso	n ignit	ion; ¿	and C	ΊA, Ρ	IA an	d CIW.	>		
SAMPLE	SiO2 TiO2 Al2O3 Fe2O3* MnO MgO CaO Na2O K2O	3 Fe ₂ O ₃	* MnO	MgO	CaO	Na ₂ O	K ₂ O P ₂ O ₅	LO	CIA P	PIA CIW	V As	Pb	Zn (Cu	Ni Cr	>	Sr Y	dN ×	o Zr	님	Sc	TS	БB	_ _	Ö
Houzyou 1	66.20 0.38 17.31	1 3.77		0.08 1.98	5.32	3.28	1.60 0.07	06.0	51 5	51 54	~	15	15	4	7 24		239 1	19 2	73	со С	4	198	9 9	31 293	3 447
Houzyou 2	82.23 0.11 10.05	5 0.89	0.02	2 0.24	0.84	1.83	3.77 0.02	0.50	54 5	57 69	9	13	6		4 23		131 2	20 1	59	m		162	5 2	25 .	355
Tomari	73.57 0.20 14.62	2 1.87	. 0.04	t 0.66	2.78	3.07	3.15 0.04	0.90	52	53 59	14	13	22	ю	7 19	9 5	371 1	17 2	79	ო	7	62	10 2	26 4064	34 621
Ishiwaki	68.52 0.36 16.69	9 3.21		0.06 1.15	3.68	3.39	2.89 0.06	1.20	52	53 58	14	14	32	Ð	8 23	3 32	455	17 3	80	4	10	350	0 7	22 3404	04 613
Anedomari	68.61 0.23 17.35	5 2.28		0.04 0.87 4.30		3.50 2	2.77 0.05	1.80	51 5	52 56	16	14	30	4	8 22	6 2	506 1	14 2	73	Ю	7	157	10 2	20 2753	53 580
Anedomari	66.20 0.38 17.31	1 3.77		0.08 1.98	5.32	3.28	1.60 0.07	06.0	51 5	51 54	÷	11	40	ص	12 24	4 43	598	12 3	61	ო	15	115	с Т	15 .	355
Hamamura	78.03 0.14 12.39	9 1.46		0.03 0.42 2.22		2.61	2.68 0.03	1.30	53	53 60	17	13	19	4	6 18		287 1	15 1	75	N	N	156	10 2	28 4505	5 679
Hamamura	76.56 0.16 13.43	3 1.70		0.03 0.58	2.52	2.58	2.39 0.03	0.80	54	55 61	16	12	22	N	7 24	4 0	342 1	14 2	78	N	4	131	0 8	26 .	342
Hakuto	77.43 0.28 11.12	2 2.84		0.06 1.26 2.52		2.34	2.11 0.03	1.30	51 5	51 57	15	:	29	D	12 34	4 24	277	15 3	77	ლ	2	200	10 2	23 3766	36 598
Karo	70.72 0.48 14.46	6 3.91		0.08 1.67 3.59		2.90	2.15 0.05	1.50	52	52 56	15	12	41	4	14 36	5 49	396	16 4	. 89	ო	10		9	10 2951	51 576
Sakyuuhigashi	77.17 0.25 12.36	6 2.30		0.05 0.86 1.84		2.44	2.69 0.04	1.40	55	57 63	16	14	29	ന	13 36	6 18	227	17 3	78	4	9		10 2	24 2651	51 549
Makitani	68.58 0.22 14.79	9 2.09	0.04		0.86 7.49	2.91	2.96 0.05	5.50	41	39 45	16	13	30	4	9 23	о е	526 1	17 2	69	ო	÷	170	12 1	17 5198	966 86
Makitani	69.88 0.25 16.43	3 2.45		0.05 1.00 4.40		3.04	2.45 0.05	1.90	51 5	52 56	17	13	36	O	11 31	1 22	486	16 3	73	e	÷	102	3	22 .	432
Higashihama	68.16 0.25 16.03	3 2.40	0.04	t 0.90	5.71	3.31	3.14 0.06	4.10	46 4	45 51	21	14	37	9	8 24	4 17	480	16 3	76	4	10	140	14	18 7394	94 1007
Higashihama	68.89 0.24 16.26	6 2.29	0.04	t 0.80	5.34	3.12	2.96 0.05	3.20	48 4	47 52	21	14	37	4	9 22	6	480 1	16 2	75	с С	7	37	4	18	517

Table 6. XRF maj ignition; and indic	Table 6. XRF major (wt%) and trace (ppm) element analyses gnition; and indices of Chemical Index of Alteration (CIA), Plagi	(ppm) x of Alte	elemer ∍ration	nt ana (CIA),	lyses (Plagic	of beach sar ioclase Index		ids from the of Alteration		Eastern S (PIA) and	'n Sai Ind Ch	San'in coast, Tango Peninsula I Chemical Index of Weathering	coast, ical Inc	Tango Peninsula dex of Weathering	o Pe Wea	ninsu		and Wakasa (CIW).	ikase	ı Bay,		Japan. I	LOI, oven-dried	-uen-	dried	loss	U
SAMPLE	SiO ₂ TiO ₂ Al ₂ O ₃ Fe ₂ O ₃ * MnO MgO CaO	e2O3*	MnO N	MgO (CaO N	Ja2O K	K ₂ O P ₂ C	D ₅ LOI	CIA	PIA C	CIW As	s Pb	Zu	CU	ī	ر ک	~	يد ۲	Ž	o Zr	₽	S	З	ш	Ъ	_	
East San'in coasts (n=17)		1																									
Kasumi Shihavama	84.67 0.24 7.42 66 82 0 34 11 86	1./0 3 31	0.04	0.54 1 27 1	2.08	1.42 1.	25600	л 48.	00 48 35 30	4/7/	22 33 33	3 15 13 15	ວα ວດ	n U	· ົ	2 2	ດ ເ	1/2 1	1 1 2 2 2 2	4Ω α			530 338	6 8 0 8 0	- 0	23 1t 26	520
Sazu	6.8.3	0.75			0.33		67 0 01		29 55						1 17							r Ç	567				1435
Yasugi	10.31	1.02			1.70		0	48.														2.	390				8.
Kirihama	8.76	1.99			13.77										ω		21 2			63		ß	513				
Takeno	5.30	2.45	0.06		43.35		0								N								514				
Kumihama 1	9.74	2.22			7.45		0	4 51.73														<u>φ</u>	936				4118
Kumihama 2	9.78	3.89			7.34		0	55							17							15	881				17
Kumihama 3	7.64	1.18	0.02		2.77		o o	49							<u>,</u>							4	765				- L
Kumihama 4	8.50	1.72 0.05			3.53		<u> </u>								~ *							1 0	731				152
Kotohikibama	0.30 7 + 7	0.85		0.0 1 0 1 0	4.79		- c	00	31 33 62 40						4							- V	090				787
Sunakata	83 80 0 19 7 27	0.00 1 73	0.03		- 2- 6		\dot{c}	02 53 9							•				 ວິດ	76		9 5 7 2	1551				4356
Taiza	7.92	0.43			0.61		0	53.												71			528)
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Table 1. XFF major (wfs) and trace (ppm) element analyses of bach stards (row (b)) and trace (ppm) element analyses of bach stards (row (b)) and trace (ppm) element analyses of bach stards (row (b)) and trace (ppm) element analyses of bach stards (row (b)) and trace (ppm) element analyses of bach stards (row (b)) and (row (b))		Ū	3727								5078	1476	510	4227	2047	3093	1414	5392	2823	6019	3110	39	5187	5808	6273	5998	6730	3734	5081	6044	5078	3994
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	ach	K ₂ O	2.51	2.45	2.93	2.60	2.37	2.64	2.59	1.80	2.07	1.66	1.98	2.22	2.27	2.34	2.53	2.35	2.55	2.49	2.68	3.05	2.78	2.65	2.41	2.95	2.68	2.52	2.67	2.49	2.48	2.75
Table 7. XFF major (wt%) and frace (ppm) element analyses SMPLE SiO MIO MIO <th< td=""><td></td><th></th><td>0</td><td>.12</td><td>88.</td><td><u> 60</u></td><td>5</td><td>0</td><td>.84</td><td>.38</td><td>.86</td><td>.58</td><td>77.</td><td>.07</td><td><u>.</u>0</td><td>.02</td><td>С</td><td></td><td>9</td><td></td><td>\sim</td><td>G</td><td>.38</td><td>29</td><td>9</td><td></td><td>35</td><td>.24</td><td>с.</td><td>22</td><td>28</td><td><!-- -->N</td></th<>			0	.12	88.	<u> 60</u>	5	0	.84	.38	.86	.58	77.	.07	<u>.</u> 0	.02	С		9		\sim	G	.38	29	9		35	.24	с.	22	28	N
Table 7. XFF major (wt%) and trace (ppm) element analy SAMPLE SIO TIO Mod Mod Ca SAMPLE SIO TIO Allo Mod Mod Mod Ca Samrhama-B 75.44 0.49 12.52 3.84 0.07 1.06 1.37 0.14 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.33 0.05 0.10 1.0 0.05 0.33 0.05 0.11 1.0 0.05 0.33 0.05 1.0 0.05 0.11 1.0 0.05 0.11 1.0 1.12	ses								-	-		-																				
Table 7. XRF major (wt%) and trace (ppm) elements SAMPLE SiO2 TiO2 Al>O3 Fe2O3* MnO MgC Santhama-A TiO2 N10 TiO2 N10 TiO3 N10 N10 Santhama-A TiO3 TiO2 N10 TiO3 S03 0.07 1.00 0.03 <th0.03< th=""> 0.04</th0.03<>	analy														-															-		
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Table T. XFr Finder (wt%) and frace (ppm) SAMPLE SiO2 TiO2 Al2O3 E>O3* Samrihamada 76.42 0.49 12.50 3.63 Awara-B 80.36 0.26 11.11 2.30 Awara-B 80.36 0.26 11.11 2.30 Awara-B 80.36 0.22 9.43 2.11 Awara-B 81.52 0.24 10.22 2.11 Kaga-B 81.52 0.24 10.49 2.04 Kaga-B 81.52 0.24 10.49 2.01 Kaga-B 81.52 0.23 1.179 2.04 Komatsu-B 81.52 0.24 10.49 2.04 Komatsu-B 82.08 0.27 10.16 2.04 Komatsu-B 82.43 0.20 1.179 2.06 Komatsu-B 78.17 0.66 8.78 2.01 Komatsu-B 78.17 0.66 8.13 2.06 Komatsu-B 78.17 0.66 8.13 2.06 Mikawa-B 79.14) elem	MnO	0.07	0.07	0.05	0.06	0.05	0.05	0.05	0.07	0.03	0.11	0.07	0.03	0.03	0.04	0.04	0.04	0.04	0.04	0.04	0.05	0.04	0.04	0.04	0.04	0.04	0.04	0.04	0.04	0.04	0.04
Table 7. XRFmajor (wt%) and traceSAMPLESiO2TIO2Al $2O3$ FeSanrihama-A75.440.4912.593Sanrihama-B76.120.4712.523Awara-B80.360.2611.112Awara-B81.790.2410.223Awara-C83.000.229.433Awara-B81.520.2410.223Awara-C83.000.229.163Kaga-B82.080.2210.163Kaga-B82.780.308.133Komatsu-B83.430.209.163Komatsu-B80.300.2711.133Komatsu-B80.300.2711.133Komatsu-B80.300.2711.423Komatsu-B80.300.2711.423Komatsu-B80.300.2711.423Komatsu-B79.140.3011.423Mikawa-B79.170.3011.423Mikawa-B77.950.3312.613Hakusan-B77.950.3312.613Mikawa-B77.950.3312.613Hakusan-B77.550.2912.773Shiroo-B77.550.3012.763Shiroo-B77.550.3212.173Senrihama-A77.550.3212.163Senrihama-A77.610.32<	mdd)		3.84	3.63	2.30	2.16	11	11	2.04	3.19	.79	5.27	3.41	2.06	60.9	2.51	2.15	2.39	2.54	09.5	2.59	2.74	2.34	2.46	2.45	2.34	2.50	09.5	2.63	2.80	2.81	2.87
Table 7. XRF major (wt%) and SAMPLE SiO2 TiO2 Als/ Sanrihama-A 75.44 0.49 12.5 Awara-B 80.36 0.26 11. Awara-B 80.36 0.24 10.2 Awara-B 80.36 0.24 10.2 Awara-C 83.00 0.22 9.4 Awara-C 81.52 0.24 10.2 Awara-C 83.00 0.22 9.1 Kaga-B 81.52 0.20 9.1 Kaga-B 82.08 0.20 9.1 Komatsu-C 80.36 0.22 10.2 Komatsu-B 80.36 0.20 11.1 Komatsu-C 80.30 0.21 11.2 Komatsu-B 79.14 0.30 11.2 Mikawa-A 79.58 0.30 11.2 Mikawa-B 77.90 0.31 11.2 Mikawa-B 77.95 0.33 12.1 Mikawa-B 77.95 0.33 12.6 Mikawa-B 77.71 0.22 12.6	trace	D ₃ Fe																														
Table 7. XRA rajor (wf%) SAMPLE SIO2 TIO2 Samrihama-A 75.44 0.49 Sanrihama-B 76.12 0.24 Awara-B 80.36 0.22 Awara-B 81.79 0.24 Awara-B 81.79 0.24 Awara-B 80.36 0.22 Awara-B 81.52 0.23 Awara-C 81.52 0.23 Kaga-A 81.52 0.23 Kaga-B 82.43 0.20 Kaga-B 82.43 0.20 Komatsu-C 80.30 0.27 Komatsu-B 80.44 0.30 Komatsu-C 80.30 0.27 Komatsu-C 80.31 0.20 Komatsu-B 77.10 0.27 Mikawa-B 77.29 0.33 Mikawa-B 77.50 0.28 Mikawa-B 77.50 0.28 Mikawa-B 77.50 0.20 Mikawa-B 77.50 0.28 Mikawa-B 77.50 0.28 Mikawa-B	and	AI ₂ O	12.5	12.5	<u>-</u>	10.2		10.4	10.1					10.9			11.4	11.4	11.0	12.1	12.6	13.0		12.2	11.7	12.9	12.5	12.1	12.6	12.1		12.7
Table 7. XRF major (v SAMPLE SiO2 Sanrihama-A 75.44 Sanrihama-B 76.12 Awara-C 80.36 Awara-C 83.00 Kaga-A 81.52 Awara-C 83.00 Kaga-A 81.52 Kaga-A 81.52 Kaga-B 82.08 Komatsu-B 80.30 Komatsu-B 82.43 Komatsu-B 80.30 Komatsu-B 80.30 Komatsu-B 80.30 Komatsu-B 79.14 Komatsu-B 79.17 Komatsu-B 70.17 Komatsu-B 70.17 Komatsu-B 70.55 Komatsu-B 70.55 Mikawa-B 70.17 Hakusan-B 77.29 Mikawa-B 77.05 Shiroo-B 77.55 Shiroo-B 77.55 Shiroo-B 77.55 Shiroo-B 77.55 Shiroo-B 77.55 Mikawa-A 77.55 Senrihama-C	vt%)	TiO ₂	0.49	0.47	0.26	0.24	0.22	0.24	0.22	0.30	0.20	0.66	0.39	0.27	0.27	0.32	0.29	0.30	0.32	0.33	0.33	0.33	0.27	0.29	0.28	0.28	0.29	0.29	0.30	0.32	0.32	0.32
Table 7. XRFmaiSAMPLESSanrihama-A75Sanrihama-A81Awara-B81Awara-B81Awara-B82Awara-B82Awara-C83Awara-C83Awara-C82Awara-C82Awara-C80Awara-C80Kaga-B82Kaga-B82Kaga-B82Komatsu-C80Komatsu-C73Mikawa-B77Shiroo-B77Shiroo-B77Shiroo-B77Shiroo-B77Shiroo-B77Senrihama-A77Senrihama-A77Senrihama-A77Senrihama-A77Senrihama-A77Senrihama-A77Senrihama-C77 <td>or (v</td> <th></th> <td>.44 (</td> <td>.12 (</td> <td>.36</td> <td>.79 (</td> <td>00</td> <td>.52 (</td> <td>.08</td> <td>.78 (</td> <td>.43 (</td> <td>.17 (</td> <td>.57 (</td> <td>.48 (</td> <td>30</td> <td>.14 (</td> <td>.58</td> <td>.17 (</td> <td>40</td> <td>.95 (</td> <td>.29</td> <td>.37 (</td> <td>.71 (</td> <td>00.</td> <td>.85 (</td> <td>.05 (</td> <td>.55</td> <td>66.</td> <td>.21</td> <td>.62 (</td> <td>.41 (</td> <td>.64</td>	or (v		.44 (.12 (.36	.79 (00	.52 (.08	.78 (.43 (.17 (.57 (.48 (30	.14 (.58	.17 (40	.95 (.29	.37 (.71 (00.	.85 (.05 (.55	66.	.21	.62 (.41 (.64
Table 7. XRF SAMPLE Sanrihama-f Sanrihama-f Awara-B Awara-B Awara-C Kaga-A Komatsu-B Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Komatsu-C Hakusan-A Mikawa-B Kanazawa Hakusan-A Shiroo-B Shiroo-A Takamatsu-C Takamatsu-C Senrihama-C Senrihama-C Senrihama-C Senrihama-C	maj	ŝ	_		80	81	83	8	82		83	78	80	80	80	79	79	79	78	77	77	76	77	78				77				76
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	Table	SAM	Sanri	Sanri	Awar	Awar	Awar	Kaga	Kaga	Katay	Kom	Kom	Kom	Komâ	Komá	Kom	Mika	Mika	Haku	Haku	Kana	Haho	Shiro	Shiro	Takar	Takaı	Takar	Imah	Senri	Senri	Senri	Haku

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3. 2. UCJA and UCC – Normalized compositions

As has been seen, there were many differences between the geochemical compositions of the coastal sand samples taken from the six coastal regions. These variations in composition are the result of a number of factors, including differential weathering, differences in the relative proportions of the quartz, feldspar and biogenic content of the sands in the locations and differences in the origins of the constituent substances. To compare the compositions of the beach sand in the individual sites, average values were normalized against the average UCJA, according to (Togashi et al. 2000), and against the average UCC (Rudnick & Gao, 2005). Both the UCJA-normalized and the UCC-normalized patterns for five of the six study regions' beach sands compositions had very similar shapes. These regions were: Northern Kyushu, Shimane, Tottori, Tango Peninsula, and Noto Peninsula. Normalization revealed that, with the exception of SiO_2 , elements were depleted relative both to UCJA and UCC, as shown in Figure 9. The depletion was particularly marked for the ferromagnesian elements (MgO, Fe₂O₃*, TiO₂, Ni, Cr, V) which are typically strongly depleted in felsic volcanic rocks such as granites and for mobile elements CaO, Na₂O and Sr, all of which are liable to loss during weathering (Nesbitt & Young, 1984). In contrast, the UCJA-normalized as well the UCC-normalized diagrams for the Yamaguchi beach sand samples, for most of major and trace elements were less than 1 (Figure 9), exceptions being CaO, Sr, and Th.

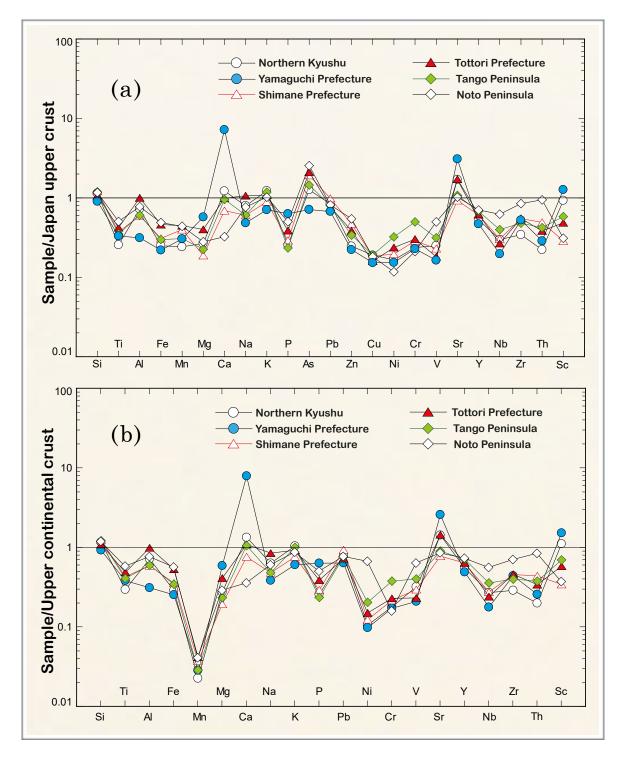


Figure 9. Average major and trace elements diagrams of beach sands collected along the coasts of South West Japan, on the coastline of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, and Noto Peninsula, normalised to average Upper Crust of the Japanese Archipelago (UCJA), according to (Togashi *et al.* 2000), and average Upper Continental Crust (UCC) (Rudnick & Gao, 2005).

3. 3. Inter-Element Relationship

Figure 10 illustrates Harker variation diagrams for selected elements in the beach sand samples, in each case the major element content being plotted against the SiO_2 content. The increase of SiO_2 reflects a greater geochemical maturity. All of the major elements plotted show general trends towards decreased abundance with increased silica content. The significant SiO_2 , Al_2O_3 and Na_2O contents of beach sands form northern Kyushu, silica rich sands form Yamaguchi, Shimane, Tottori, Noto Peninsula and Tango Peninsula, indicate that quartz and feldspar are the main constituents. The overall depletion of CaO and MgO suggests that the carbonate content of the beach sand sediments is generally low, in the majority of samples. Silicate or siliciclastic sands – those sands comprising grains originating from clasts, or fragments of silicate rocks – generally consist of feldspars, micas, quartz and other silicate substances. While the majority of the components are quartz and plagioclase feldspar, as indicated by the high SiO_2 content and significant CaO, Na_2O and Al_2O_3 contents, relatively high K_2O contents suggested significant K-feldspar and K-bearing micas contents were also present. CaO showed a well-defined decrease with increasing SiO_2 , except for higher values in a small group of Yamaguchi beach sands with lower SiO₂ contents. These samples also had higher LOI values, and hence were likely to contain a biogenic $CaCO_3$ component, such as shell material.

To better understand the major element composition, the samples were broadly subdivided as follows: silica, aluminium, and calcium (Figures 11 and 12). The geochemical compositions of these three subgroups (represented by SiO_2 , Al_2O_3 , and CaO) for the six study areas' coastlines were displayed on a modified map

of the Geological Survey of Japan (GSJ) and National Institute of Advanced Industrial Science and Technology (AIST). As expected, SiO_2 contents were greater in samples from the silica-rich sands from the Yamaguchi and Shimane pocket beaches due to their abundance of quartz; these account over 80% of silica. The maximum SiO_2 concentration was observed in the silica-rich sands from Yamaguchi, Shimane, Kotogahama and Kotobikihama, which may be due to the input of sediments supplied through major river systems from the Chūgoku Mountains to the Japan seacoast (the Takasu, Gono Hii, Hino, Tenjin and Sendai rivers). The Al₂O₃ contents varied from approximately 5wt% to over 10wt% in some samples; this variation can be attributed to the contribution of fine grains and feldspars. The maximum Al_2O_3 contents were seen in the northern Kyushu and Tottori beach sand. Beach sand samples from Tottori typically consisted of silicate minerals with little quartz but a significant quantity of feldspar. Furthermore, the presence of considerable CaO in the Yamaguchi beach sand samples indicated a significant biogenic $CaCO_3$ component; in this case, the likely constituents were Foraminifera, shells or animal skeletons, which - like sea urchin and ostracod – primarily comprise the compound (Figure 13). Biogenic carbonate sands found in the coastal zone and shallow water consist of porous or hollow particles with a rough texture.

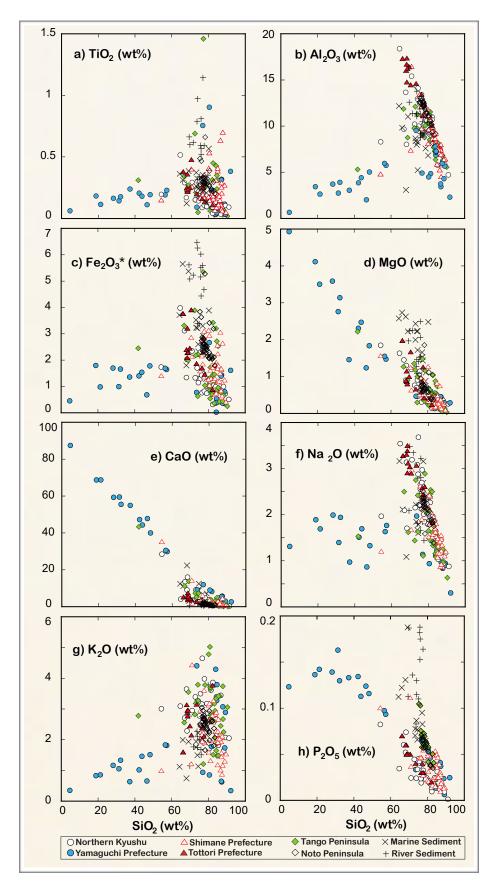


Figure 10. Harker diagrams for selected major elements of beach sands collected along the coasts of South West Japan, on the coastline of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, and Noto Peninsula.

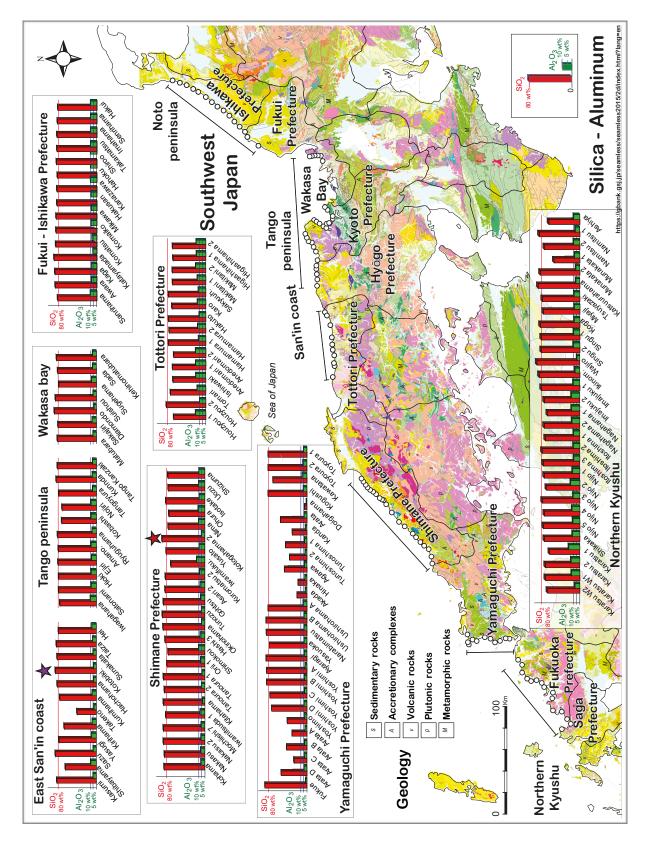


Figure 11. Geological map of South West Japan, showing the geochemical composition of SiO_2 and Al_2O_3 elements, and location of beaches sampled along the coastline of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, Wakasa Bay, and Noto Peninsula. Modified from the Geological Survey of Japan (GSJ) and National Institute of Advanced Industrial Science and Technology (AIST), 2016.

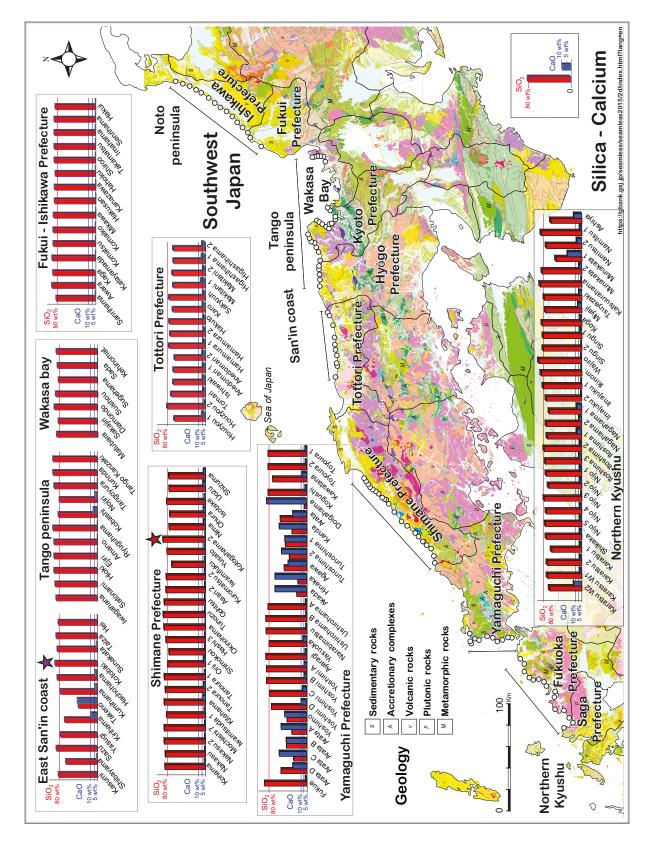


Figure 12. Geological map of South West Japan, showing the geochemical composition of SiO_2 and CaO elements, and location of beaches sampled along the coastline of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, Wakasa Bay, and Noto Peninsula. Modified from the Geological Survey of Japan (GSJ) and National Institute of Advanced Industrial Science and Technology (AIST), 2016.

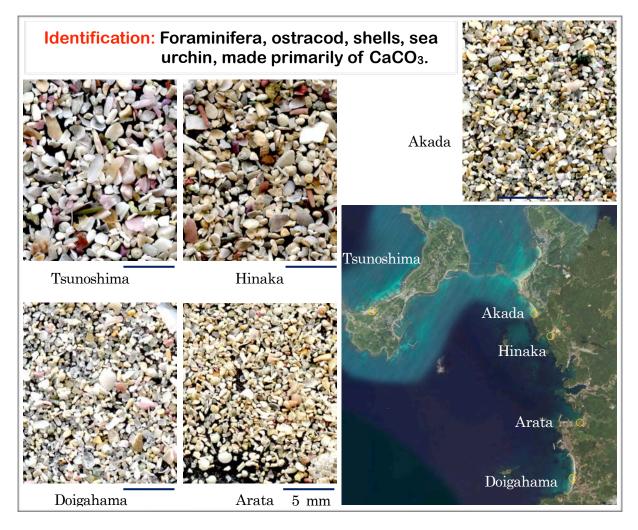


Figure 13. Selected microscopic photographs biogenic carbonate sands from the Tsunoshima, Akada, Hinaka, Arata, and Doigahama Beaches, Yamaguchi Prefecture, South West Japan. The identified species are Foraminifera, ostracod, shells, and sea urchin, made primarily of $CaCO_3$.

The variations of ecosystem and geography of marine condition in Hinaka and Arata beach are shown in figure 14. Typical small-scale pocket beach, and inlet of beach shape characteristic. Rocky points of both sides of the Hinaka beach, *Zostera marina* and other sea weed are exposed on the shore of the Hinaka beach suggesting the existence of under-water sea plant field. On the left side of the Hinaka beach, *Zostera marina* are found on the shore (Figures 15a, 15b and 15c. Doigahama beach is more strait than Arata beach (Figure 16), but sea plant habit is suggestive. Under-water structure may make a variety of geography for ecosystem.



Figure 14. The variations of ecosystem and geography of marina condition in Hinaka and Arata beaches, Yamaguchi Prefecture, South West Japan. Typical small-scale pocket beach, and inlet of beach shape characteristic.



Figure 15a. Rocky points of both sides of the Hinaka beach, *Zostera marina* and other sea weed are exposed on the shore of the Hinaka beach.



Figure 15b. Rocky points of both sides of the Hinaka beach, *Zostera marina* and other sea weed are exposed on the shore of the Hinaka beach.



Figure 15c. Rocky points of both sides of the Hinaka beach, *Zostera marina* and other sea weed are exposed on the shore of the Hinaka beach. On the left side of the Hinaka beach, *Zostera marina* are found on the shore.



Figure 16. The variations of ecosystem and geography of marine condition in Doigahama beach, Yamaguchi Prefecture, South West Japan.

Chapter Four

4. DISCUSSION

4. 1. Evaluation of biogenic productivity

To evaluate the biogenic productivity in the Yamaguchi coast, the climatic conditions and water quality were observed.

Temperature is one of the most important parameters controlling the environment. The Tsushima Warm Current (TWC) is the main surface current flowing into the Japan Sea (JS) through the Tsushima Strait and emerging through the northern Tsugaru Strait and Soya Strait. As a branch of the Kuroshio Current, the TWC is characterized by high temperature. Figure 17 shows the mean surface water mass temperature at 50 m for two periods -11^{th} August to 20^{th} August, 2016, and 11^{th} May to 20^{th} May, 2016 May, early summer Japan Meteorological Agency (2016). The current transports a large quantity of heat and warm-water marine organisms into the Sea of Japan; it is also a major source of nutrients in the JS, contributing 55% of the phosphorous and 67% of the nitrogen in the upper 200 m of the JS (Yanagi, 2002). Primary production in the JS is relatively high among marginal seas (Yamada *et al.*, 2005). This suggests that the TWC through Tsushima strait, may be responsible for high biogenic productivity in Northern Kyushu and Yamaguchi coast, which ultimately influences contents of beach sand geochemistry. The occurrence of tidal flats is another factor in higher carbonate productivity. At Arata beach, Yamaguchi, the tidal difference exceeds 1 m (Figure 18).

Water quality types for beaches are assessed by Prefectural Government of Fukuoka in northern Kyushu and Yamaguchi for the Ministry of Environment of Japan (2016). Beach classification is based on two factors, the chemical oxygen demand (COD) and the faecal coliform content. Classification is into one of three categories – in decreasing order of quality, Type AA and Type A beaches are considered to be 'good' quality, while Type B beaches are considered to be 'good' quality, while Type B beaches are considered to be 'good' quality, Type AA and Type A beaches are considered to be 'good' quality, while Type B beaches are considered to be of 'satisfactory quality. The specifications for each of the factors, as shown in Figure 19, are as follows: Type AA – COD < 2 g/l, faecal content not detected; Type A – COD < 2 mg/l, faecal content < 100 / 100 ml; Type B – COD < 5 mg/l, faecal content < 400 / 100 ml. The water quality, as designated by the Type, could affect the biogenic CaCO₃ productivity. The Yamaguchi beaches are classified as Type A.

4. 2. Geochemical Maturity

Pettijohn *et al.* (1972) first discussed the concept of geochemical maturity in sediment, believing that maturity should be assessed using the QFL diagram, but there are more than 50 sand and sandstone classification systems. It is necessary to use laboratory analysis is required with point counting (usually considering 500 points) and petrographic thin sections; carbonates are excluded. Therefore, the QFL methods, the beach shape and grain size are not relevant in the context of this study.

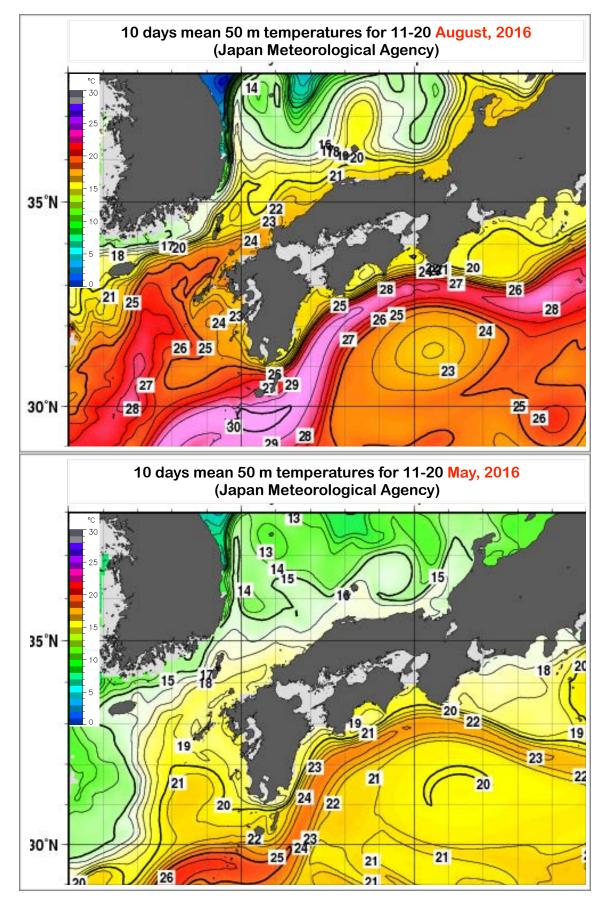


Figure 17. Surface water mass temperature; mean 50 m in western Japan for 11th-20th August 2016. In addition, for 11th-20th May, 2016 (Japan Meteorological Agency).

Geochemical maturity is a compositional state of a clastic sedimentary body in which quartz is dominant, while less-resistant particles, such as feldspars, detrital carbonates or lithic fragments, are either absent or present in much smaller quantities (Blatt *et al.*, 1972; Pettijohn *et al.*, 1972; Daniel, 2004). Geochemically mature sandstones are classified as quartz arenites or orthoquartzites if they comprise at least 95% quartz. Sedimentary petrologists and sedimentologists (Schwab, 1975; Potter, 1978; Suttner *et al.*, 1981; Franzinelli & Potter, 1983; Suttner & Dutta, 1986; Johnsson *et al.*, 1988; Potter 1994; Nesbitt & Young, 1996; Nesbitt *et al.*, 1996, 1997; Potter *et al.*, 2001; Daniel, 2004) have extensively used geochemical maturity as a key indicator of sediment provenance, transport history, and weathering history.

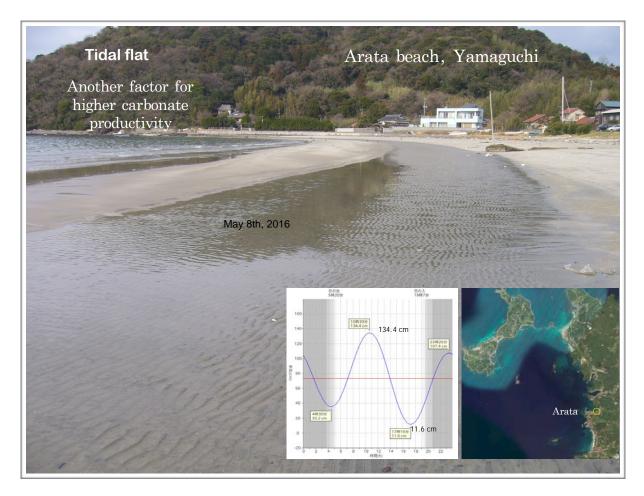


Figure 18. Tidal-flat deposits on the foreshore of Arata beach, Yamaguchi, South West Japan.

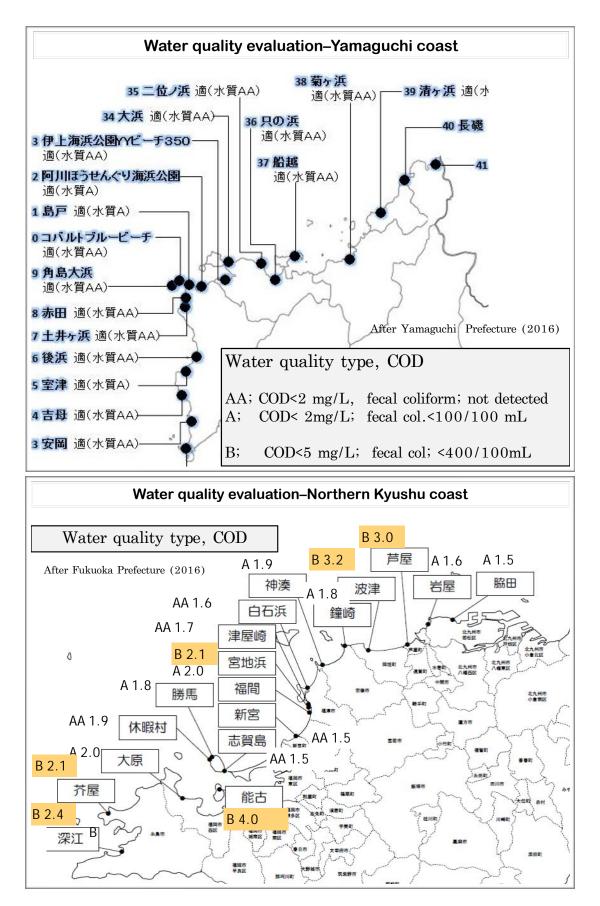


Figure 19. Classification based on the chemical oxygen demand (COD) and faecal coliform criteria of levels of water quality in Yamaguchi in contrast to level water quality in northern Kyushu, South West Japan.

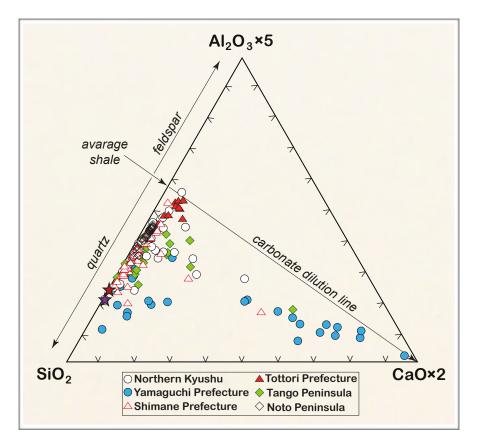


Figure 20. Ternary diagram of relative proportions of $Al_2O_3 \times 5$, SiO_2 , and $CaO \times 2$ (Brumsack, 1989), of beach sand samples from Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, Wakasa Bay, and Noto Peninsula, southwest Japan. An arbitrary multiplier of 5 and 2 are used respectively for Al_2O_3 and CaO in order to better distribute the data points within the graph.

In a very simplistic way, the investigated beach sand from the coasts of Southwest Japan, on the coastline of Northern Kyushu, Yamaguchi, Shimane, Tottori, the Tango Peninsula and the Noto Peninsula comprise variable mixtures of terrigenous detritus (represented by Al_2O_3 and SiO_2) and biogenous material (represented by CaO). To compare the relative proportions of the major components, the relative proportions of CaO (mostly calcium carbonate), SiO_2 (quartz and alumosilicates), and Al_2O_3 (alumosilicates, feldspar) were plotted in a triangle diagram (Figure 20) (Brumsack, 1989). The diagram showed that the sediments from Yamaguchi trend from the CaO to the SiO_2 poles showing a distinct distribution pattern along the carbonate dilution line, indicating that they represented a simple background sedimentation diluted by biogenic carbonates. By contrast, most northern Kyushu beach sand and the sands rich in silica from Kotogahama, Kotobikihama, Shimane, Yamaguchi and the Tango Peninsula plotted on a $SiO_2-Al_2O_3$ mixing line. This was consistent with these samples having a broadly lower content of carbonate and more variable distribution. The shift towards the SiO_2 edges indicated excess silica contents; this reflected an abundance of coarse-grained particles and was particularly high in those samples exhibiting a low carbonate content. It is probable that tides and rivers deposited these sands. The majority of beach sand from Noto Peninsula and Tottori plotted on a straight mixing line between quartz and feldspars components richer in Al_2O_3 and Average Shale. This suggested that more intensely weathered clays (possibly with a higher kaolinite proportion) were characteristic for these sediments.

4. 3. Geochemical classification

Maturity is reflected best in quartz, rock fragments, feldspars and grain size. As the percentage of quartz increases, the mineralogical maturity also increases. The SiO_2/Al_2O_3 ratios of clastic rocks are sensitive to sediment recycling and weathering processes and can be used as indicators of sediment maturity. With increasing sediment maturity, quartz survives preferentially to feldspars, mafic minerals and lithics (Roser & Korsch, 1986; Roser *et al.*, 1996). Average $SiO_2/$ Al_2O_3 ratios in unaltered igneous rocks ranging from approximately 3.0 (that is, basic rocks) to approximately 5.0 (acidic rocks). Values of the SiO_2/Al_2O_3 ratio greater than 5.0 in sandstones are an indication of progressive maturity (Roser *et* al., 1996). SiO₂/Al₂O₃ and Na₂O/K₂O ratios may vary depending on the maturity of the sediments. Low values of SiO₂/Al₂O₃ ratios and high values of Na₂O/K₂O indicate mineralogically immature sediments. The SiO₂/Al₂O₃ ratio was higher in the silica rich sands from Yamaguchi, Shimane and the Noto Peninsula than in northern Kyushu, the Tango Peninsula and Tottori sands (Figure 21a). Similarly, the Na₂O/K₂O ratio followed a similar inverse trend; it was high in Tottori, Tango Peninsula and northern Kyushu sands and lower in the Noto Peninsula, Shimane and the silica-rich sands from Yamaguchi (Figure 21b). The lower Na₂O/K₂O ratios in the studied sands were attributed to the enrichment of K-feldspar compared to plagioclase.

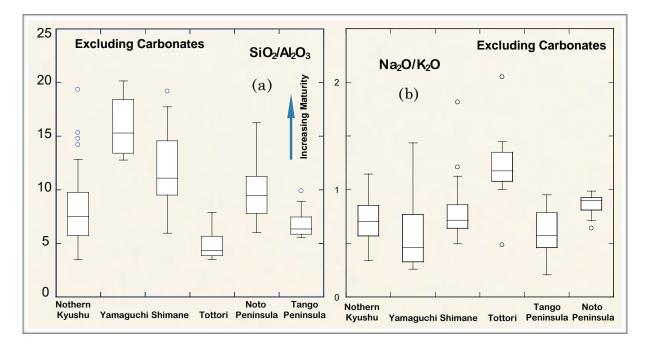


Figure 21. Box plots showing the SiO_2/Al_2O_3 and Na_2O/K_2O ratios in the investigated beach sands from the coasts of South West Japan, on the coastline of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, and Noto Peninsula.

The most used classification parameters are the SiO_2/Al_2O_3 ratio, primarily reflecting the abundance of quartz, clay and feldspar, the Na_2O/K_2O ratio, which defines an index of chemical maturity and the Fe_2O_3*/K_2O ratio that defines an index of mineral stability. In Figure 25, the log ratios of Na₂O/K₂O are plotted against the log ratios of SiO₂/Al₂O₃, and the log ratios of Fe₂O₃*/K₂O are plotted against the log ratios of SiO₂/Al₂O₃ (Pettijohn *et al.*, 1972; Herron, 1988).

Figures 19a and 19b show the geochemical classification diagrams of Pettijohn (1972) and Herron (1988) for the beach sand of the same coastal regions as were investigated in this study. As can be seen, they demonstrated the same petrographic results. While the sands from the Noto Peninsula can be seen in the Na_2O/K_2O vs. SiO_2/Al_2O_3 diagram of Figure 22a to have plotted in the arkose field, most of those in the Tango Peninsula, Shimane, northern Kyushu and Yamaguchi plotted in the arkose and subarkose fields. Tottori beach sands are classified as litharenites and arkoses. Figure 22b shows that the investigated beach sand from the coasts of all six investigated areas of Southwest Japan were bracketed by arkose and subarkose, with a diminishing trend towards sublitharenite, reflecting the increasing abundances of quartz and feldspar. The diagram may be used to ascertain not only the dominant mineral components, as previously considered, but also to identify trends in soil maturity and aging. That is, beach sands rich in feldspars are usually classified as young, while those in which quartz is dominant are considered more mature, the higher age indicating greater transport. By observation of sands holder of the Hashi beach, Shimane Prefecture, beach sands composed primarily of well-sorted quartz, a durable mineral that is hard and does not weather easily in Figure 23. In addition, grain size distribution of beach sand collected at the shorelines of selected beaches on the western San' in coast is shown in Figure 24.

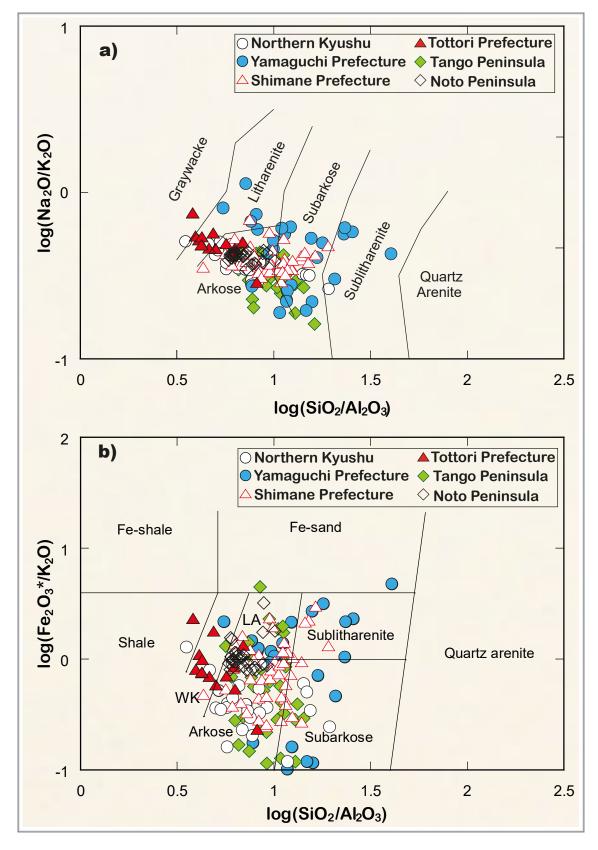


Figure 22. Geochemical classification schemes of beach sands along the coastline of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, Wakasa bay, and Noto Peninsula. Based on: **a**) $\log(SiO_2/Al_2O_3)$ versus $\log(Na_2O/K_2O)$ diagram of Pettijohn *et al.* (1972), and **b**) the log (SiO_2/Al_2O_3) versus log $(Fe_2O_3^*/K_2O)$ diagram of Herron (1988). LA. Litharenite and WK. Wacke.

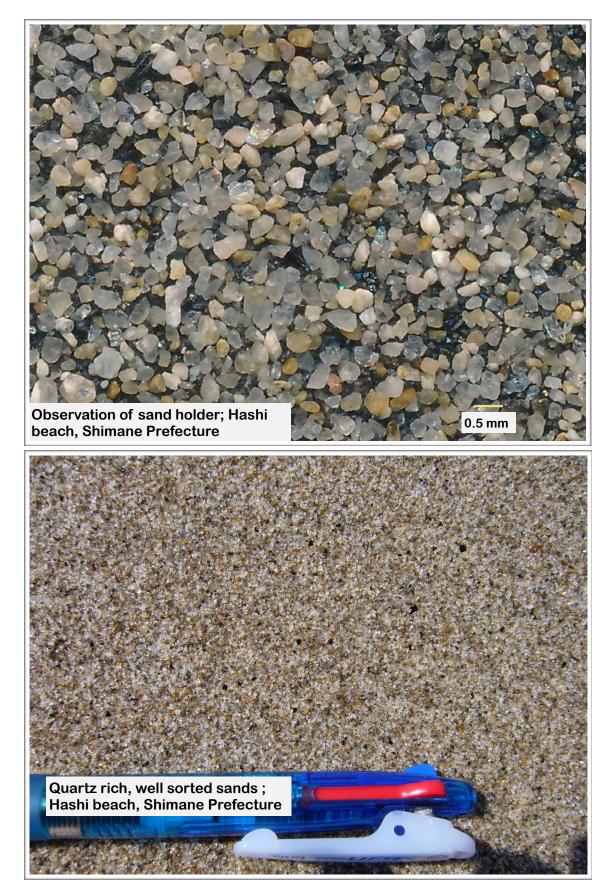


Figure 23. Observation of sands holder of the Hashi beach, Shimane Prefecture. Beach sands composed primarily of well-sorted quartz, a durable mineral that is hard and does not weather easily.

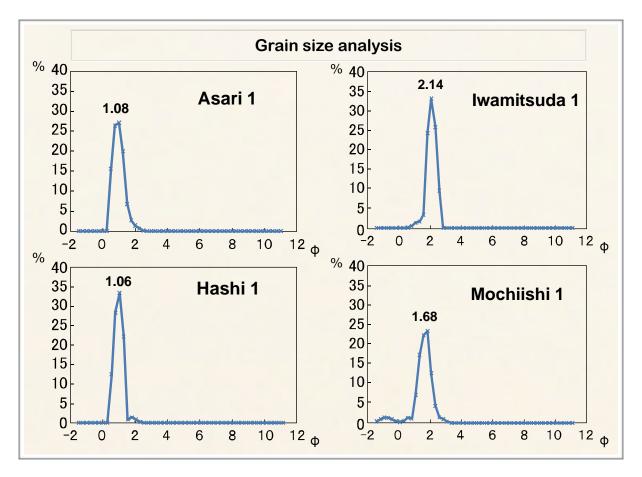


Figure 24. Grain size distributions of beach sands collected at the shorelines of selected beaches on the western San' in coast. Ishiga *et al.* (2010).

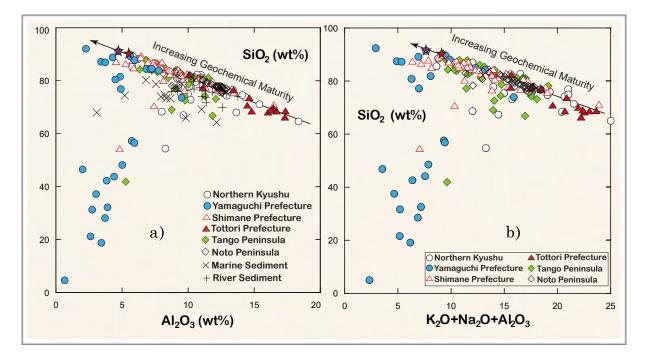


Figure 25. a) Bivariate plot of SiO_2 against Al_2O_3 , and **b)** plot of SiO_2 (reflective of quartz content) versus $K_2O+Na_2O+Al_2O_3$ (reflective of feldspar content) of beach sand samples from Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, Wakasa bay, and Noto Peninsula, South West Japan.

A bivariate plot of SiO₂ against Al_2O_3 is shown in Figure 25a; Figure 25b is a plot of SiO₂ (reflective of quartz content) against $Na_2O+K_2O+Al_2O_3$ (reflective of feldspar content). The latter plot is representative of quartz content against feldspar content and overall reflects chemical maturity as a function of climate, according to Suttner and Dutta (1986). The plotted samples revealed semi-arid to semi-humid climatic conditions in the area from various sources tending towards increasing chemical maturity. Beach sand samples from Shimane and quartz-rich sands from Yamaguchi showed high degrees of maturity, which probably indicate rich chemical weathering in the respective source areas. Overall, the investigated beach sands from the northern Kyushu, Tottori, Tango Peninsula and Noto Peninsula beaches showed variable degrees of chemical maturity, ranging from low to intermediate levels. The beach sands plotted in the semi-arid region may have experienced little or no chemical weathering and are far less mature than those plotted in the semi-humid area.

The composition of beach sand is highly variable according to the local rock sources and conditions. While the primary component of beach sand from silicarich sands from Yamaguchi, Shimane, Kotogahama, Kotobikihama and the Tango Peninsula is quartz, or silica (SiO₂), the distributions of sand sediment along these beach areas are substantially regulated by the influence of sand the sediments derived from the Chūgoku Mountains. large quantity of sediment originates from the basin areas located upstream, as most of the river basins lie above granite, which is easily weathered (Somura *et al.*, 2012). Meanwhile, Tottori sands were largely composed of weathered feldspar particles. A number of influences on these sediments, including the fluvial systems and erosion of the older materials that underlie the inner shelf. Furthermore, it is likely that further erosion is the result of longshore currents and coastal mechanisms, including reworking of the beach and winnowing. Admitting that sediments from Tottori sand dunes are formed from sediments derived from the nearby Chugoku Mountains and transported by the Sendai River to the ocean. In contrast, the biogenic carbonate sands from Yamaguchi are primarily composed of shell fragments, which might be expected, given the influence of the warm Tsushima Strait in increasing biogenic productivity in the region.

4. 4. Weathering process

Wronkiewicz & Condie (1987) stated that climate and the rate of tectonic uplift determine the extent of chemical weathering. Reduced tectonic activity and a shift towards warmer, more humid conditions are correlated with an increase in chemical weathering (Jacobson *et al.*, 2003). Therefore, the weathering indices of sedimentary rocks can provide useful information regarding the tectonic activity and climatic conditions in the source area. Various researchers have proposed different weathering indices (Nesbitt & Young, 1982; Harnois, 1988) commonly used by many researchers across the globe. Chemical weathering strongly affects major element geochemistry of siliciclastic sediments (Nesbitt & Young, 1982; Johnsson *et al.*, 1988; McLennan, 1983).

The Chemical Index of Alteration (CIA), a quantitative measure proposed by Nesbitt and Young (1982, 1984), is a potentially useful means of evaluating the degree of chemical weathering and was used to evaluate the extent of

weathering in this study. This index measures the extent to which feldspar has been converted to aluminous weathering products. CIA ratios in feldspar and fresh source rocks are typically ~ 50 , whereas those in residual weathering products such as kaolinite and gibbsite can reach 100. This index can be calculated using molecular proportions, from the formula: $CIA = [Al_2O_3]/$ $(Al_2O_3+CaO^*+Na_2O+K_2O)$ × 100. CaO* was corrected using the subsequent methodology proposed by McLennan et al., (1993), in which CaO values are accepted only if $CaO < Na_2O$; when $CaO > Na_2O$, it is assumed that the concentration of CaO equals that of Na_2O . High CIA values reflect the removal of mobile or unstable cations (Ca, Na, K) relative to highly immobile or stable residual constituents (Al, Ti) during weathering (Nesbitt & Young, 1982). Conversely, low CIA values indicate the near absence of chemical alteration and consequently may reflect cold and/or arid conditions (Nesbitt & Young, 1982, 1989). As defined by Nesbitt and Young (1982), a CIA of between 50 and 60 represents incipient weathering, a value between 60 and 80 represents intermediate weathering, while a value greater than 80 represents extreme weathering.

The CIA values are shown in Figure 26. As can be seen, the average CIA values were: Tottori beach sand 51 (range 41 - 55), the Tango Peninsula 53 (range 35 - 62) and the Noto Peninsula 58 (range 54 - 60). The slightly elevated CIA of beach sands from Shimane, Tottori, Tango Peninsula, and Noto Peninsula relative to average granodiorite reflected very weak alterations due to chemical weathering, suggesting that progressive weathering had occurred.

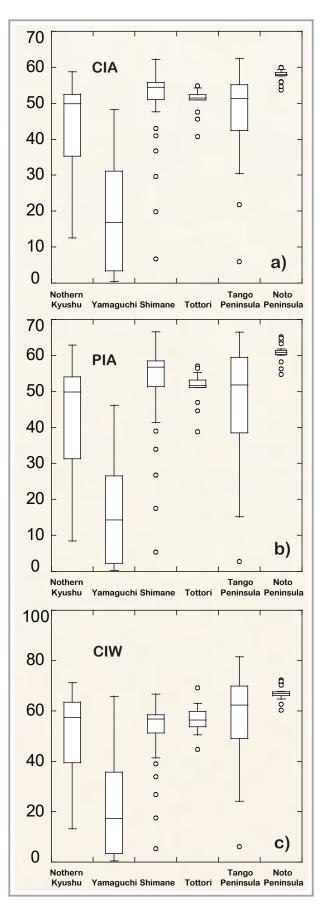


Figure 26. Box-plot diagrams of geochemical weathering indices. a) Chemical Index of Alteration (CIA), b) Plagioclase Index of Alteration (PIA), c) and Chemical Index of Weathering (CIW).

The overall range of the CIA values was similar to or slightly greater than the CIA values of the UCC. These CIA values indicate that the sediments were slightly weathered.

The degree of the chemical weathering can be estimated using the Plagioclase Index of Alteration (PIA) modified from the CIA equation to monitor plagioclase (Fedo *et al.*, 1995). The plagioclase index of alteration is calculated according to the following equation in molecular proportions: PIA = $(Al_2O_3-K_2O)$ / $(Al_2O_3+CaO^*+Na_2O-K_2O) \times 100$. High PIA values (>84) indicate intense chemical weathering while lower values (~50) are characteristic of unweathered, or fresh, rock samples. Post-Archean Australian Shales (PAAS) have a PIA value of 79. The PIA values of beach sands from the northern Kyushu coast ranged from 44 to 63. The mean PIA value for the Tottori beach sand was 51 and for the Noto Peninsula was 58.

Weathering effects can also be evaluated using the Chemical Index of Weathering (CIW) in molecular proportions) identical to the CIA, from the formula: CIW = Al_2O_3 / $(Al_2O_3+CaO^*+Na_2O)$] × 100. This equation is more appropriate in understanding the extent of plagioclase alteration alone since K_2O is subtracted from Al_2O_3 in the numerator and denominator of the CIA equation. However, it should be noted that Fedo *et al.* (1995) argued that the CIW was, in fact, inappropriate as a means of quantifying the intensity of chemical weathering. His reasoning was that there was potential for misinterpretation of the resultant CIW values. For example, the CIW of 80 for unweathered potassic granite was similar to that of residual products of smectite, while the CIW of 100 for clay minerals

(for example, gibbsite, illite and kaolinite) was similar to that of residual products of the same minerals. Interpretation of CIA and CIW is similar; a value of 50 represents unweathered UCC and a value in the order of 100 represents highly weathered materials with complete removal of alkali and alkaline-earth elements (McLennan *et al.*, 1983; McLennan, 1993; Mongelli *et al.*, 1996).

The CIW values of beach sands from the Eastern San' in coast, Tango Peninsula and Wakasa Bay ranged from 51 to 72, 56 to 71 and from 62 to 82 respectively. The average CIW values for Shimane, the Tango Peninsula and the Noto Peninsula sands (60, 62, and 67 respectively) were slightly higher that those of the northern Kyushu, Yamaguchi and Yamaguchi coasts (Table 8). The CIW index values are higher than CIA values for the analysed samples, on account of the exclusion of K_2O from the index. A low to moderate weathering of the beach sands collected from the six coastal regions of interest in South West Japan was determined from the calculated CIW values.

4. 5. Palaeoweathering indices in A-CN-K and A-C-M diagrams

Nesbitt and Young (1984) and Fedo *et al.* (1995) used the ternary diagrams $Al_2O_3-(CaO+Na_2O)-K_2O$ (the A-CN-K diagram) and $Fe_2O_3*+MgO-(CaO+Na_2O+K_2O)-Al_2O_3$ (the A-CNK-FM diagram) to infer weathering profiles. These diagrams are typically plotted such that A (that is, Al_2O_3) is located at the top apex (see Figure 27), CN (CaO*+Na_2O) is located at the bottom left apex and K (K₂O) is located at the bottom right apex. Interpretation of these plots assists in the understanding of weathering patterns and mineralogical composition,

as described by Nesbitt and Young (1985, 1989). Plagioclase and K-feldspar plot at 50% Al_2O_3 on the left and right boundaries, respectively to form the feldspar join. The clay mineral groups, kaolin, chlorites and gibbsite plot at the A apex (100% Al_2O_3). The initial weathering trends of igneous rocks are subparallel to CN-A. Calcite plots at the CN apex. Illite and smectites plot on the diagram at 70% and 85% Al_2O_3 . As weathering progresses, clay minerals are produced at the expense of feldspars and bulk composition of soil/sediments samples evolve up the diagram towards the A apex, along the weathering trend. Therefore, the samples that have been weathered most heavily will be dominated by aluminous clay minerals, which will be reflected in the A-CNK-FM diagram by positions closest to the A apex. The weathering trend intersects the A-K boundary once all plagioclase is weathered and then is redirected towards the A apex because K is extracted from the residues in preference to Al (Nesbitt *et al.*, 1996).

As shown in the ternary A-CN-K plot in Figure 24a, the CIA ratios of the investigated beach sands from Shimane, Tottori and the Tango and Noto Peninsulas were generally low (less than 60), indicating minimal weathering. In the A-CN-K ternary plot (Figure 24a), the majority of the investigated beach sand from northern Kyushu, Shimane and the Tango and Noto Peninsulas occupied the central part of the triangle, generally closer the A-CN line and plotted close to the plagioclase and K-feldspar lines, suggesting poor weathering conditions. Most of the beach sands from this group were around the weathering trends of granites and felsic volcanic rock near the plagioclase-K-feldspar line. Some beach sand samples from northern Kyushu, Yamaguchi, and the Tango

Peninsula scattered below the feldspar line, but close to the A-CN side of the diagram, confirming that due to their high CaO content they possess low CIA values.

Weathering trends may, as described by Nesbitt and Young (1984, 1989) also be observed in the molar proportions of Al_2O_3 (A), $CaO^*+Na_2O+K_2O$ (CNK) and FeO*+MgO (FM). In these diagrams, the upper (A) apex represents Al_2O_3 , the lower left (CNK) apex CaO*, Na₂O and K₂O and the lower right (FM) apex FeO* (total iron as FeO*) and MgO (Figure 24b). Plagioclase plus K-feldspar (Fel) plot on the left-hand boundary at 50% Al_2O_3 , illite plots on the left boundary at approximately 75% and greater, while Al_2O_3 , kaolin and gibbsite plot at the A apex. Biotite plots three-quarters of the way along the line between feldspars and the FM apex and chlorite plots on the right-hand boundary as a solid solution ranging from approximately 15% to 25% Al_2O_3 .

In this study, the majority of the beach sands investigated were positioned on a line connecting the FM apex with the Fel point on the A-CNK boundary of the A-CN-K diagram. Moreover, these sands were located close to the feldspar composition (that is, close to the Fel point). The average UCC and UCJA also plotted in a similar position. The exceptions to this were the sands from Yamaguchi and the tango Peninsula. Overall, on both the A-CN-K diagram (Figure 24a), and the A-CNK-FM diagram (Figure 24b), all investigated beach sands from the coasts of South West Japan, on the coastlines of Northern Kyushu, Yamaguchi, Shimane, Tottori and the Tango and Noto Peninsulas displayed an incipient weathering history.

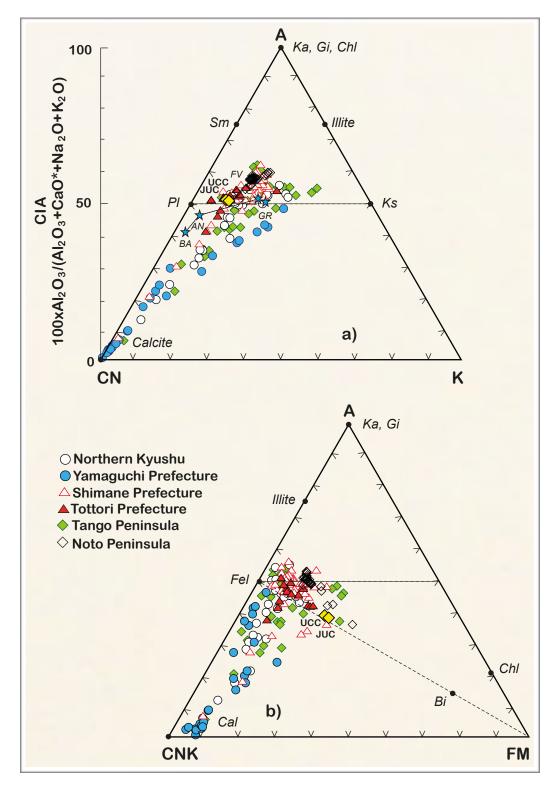


Figure 27. (a) A-CN-K and (b) A-CNK-FM (after Nesbitt and Young, 1984; Fedo *et al.*, 1995) and CIA showing weathering trends for investigated beach sands from the coasts of South West Japan, on the coastline of Northern Kyushu, Yamaguchi, Shimane, Tottori, Tango Peninsula, and Noto Peninsula. Ka= kaolinite; Chl= chlorite; Gi= gibbsite; Sm= smectite; Pl= plagioclase; Ks= K-feldspar; Fel= feldspar; Bi= biotite. Dotted line linking stars is the compositional trend in pristine average Phanerozoic-Cenozoic igneous rocks (Condie, 1993). Stars: BA= basalt, AN= andesite, FV= felsic volcanic rock, GR= granite. A= Al₂O₃; CN= CaO*+Na₂O; K= K₂O; CNK=CaO*+Na₂O; FM= FeO*+MgO.

Chapter Five

5. CONCLUSIONS

In this study, beach sands from the coastlines of the northern Kyushu, Yamaguchi, Shimane, Tottori and the Tango and Noto Peninsulas regions of South West Japan were investigated. Geochemical classification schemes of $\log(Na_2O/K_2O)$ versus $\log(SiO_2/Al_2O_3)$, and $\log(Fe_2O_3^*/K_2O)$ against $\log(SiO_2/Al_2O_3)$ showed the sands to be bracketed by arkose and subarkose. Furthermore, diminishing trends towards sublitharenite were seen, reflecting increasing abundances of both quartz and feldspar. Sands from Yamaguchi, Shimane, Kotogahama and Kotobikihama, which are rich in silica, were found to be highly mature geochemically, with SiO₂ content over 85wt%, diluting all other components. Sands from Tango Peninsula as well as Wakasa Bay were very similar showing moderate geochemical maturation. Beach sands from Tottori and the Noto Peninsula have lower SiO_2/Al_2O_3 values, reflecting the abundance of feldspar, suggesting geochemical immaturity. Those beach sands are derived from the local granite bedrock so is composed primarily of quartz, feldspar, plagioclase and trace iron-rich minerals.

After normalization to both the UCC and the UCJA, the compositions of the sands from northern Kyushu, Shimane, Tottori and the Tango and Noto Peninsulas demonstrated very similar shapes. A moderate degree of geochemical maturation was suggested for these five sites by the relative depletion of all elements other than SiO_2 in comparison with respect to the UCC and UCJA. The majority of the sediments formed under arid or semi-arid conditions tended towards increasing chemical maturity, suggesting that the beach sands are from multiple sources. The Chemical Index of Alteration, Plagioclase Index of Alteration and Chemical Index of Weathering suggested below average to moderate weathering conditions in the source area as well as immature to moderately mature beach sand sediment. This may reflect cold and/or arid climate conditions, encouraging an increase in chemical maturity in the source area.

Sands of northern Kyushu displayed lower carbonate contents than might have been expected given the warm-water currents there. This was attributed to the low water quality (Type B), which would reduce biogenic CaCO₃ productivity. By contrast, the Yamaguchi sands exhibited high to moderately low carbonate contents due to the abundance of warm-water species and good water quality (Type AA). The contents of local river and near-shore marine sediments differed significantly from those at Yamaguchi, suggesting that the inputs of existing river or marine sediment to the beach from currents or storm events were minimal.

Silica is not only maturity, if less derivation of clastics, carbonate productivity dominated, therefore, carbonate sand maturity could be one new word. In relation to Global warming, carbonate productivity could be accelerated, so the data of Yamaguchi beach will be a background for evaluation of the climate change.

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