

Data

Major and trace element analyses of sandstones and mudstones from the Siwalik Group, Bakiya Khola, central Nepal

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Abstract

This report contains whole-rock major and trace element analyses of 206 Middle Miocene to Pleistocene sandstones and mudstones from the Siwalik Group in the Bakiya Khola area of Central Nepal, as part of a continuing study of their geochemistry. Data are presented for the Rapti (n=85), Amlekhganj (n=90), Churia Khola (n=10) and Churia Mai Formations (n=21). All were analyzed by X-ray fluorescence for the major elements and 18 trace elements. Lithotype averages for each formation show SiO₂ contents generally increase up section, whereas all other major elements decrease. CaO abundances are enriched in the Rapti and Amlekhganj suites. Trace element averages also fall up section in response to the increase in SiO₂. Average trace element contents in the mudstones of each formation are generally higher than in their companion sandstones, suggesting the proportion of clay minerals present controls their abundances. Zirconium is an exception, reflecting zircon concentration in the sandstones. Normalization against average upper continental crust (UCC) shows most elements are present at crustal levels, especially in the sandstones. However, Na₂O, Sr and to a lesser extent CaO are strongly depleted, whereas Zr, Th, Ce, Y are enriched relative to UCC. Ferromagnesian elements (Sc, Fe, Ti, Ni, Cr, V) are also enriched in the mudstones relative to UCC. The causes of these anomalies will be examined in future work.

Key words: Geochemistry, major and trace elements, sedimentary rocks, Siwalik Group, Bakiya Khola, Nepal

Introduction

The Mid-Miocene to Pleistocene Siwalik Group of Nepal is situated between the Lesser Himalaya in the north and the Indo-Gangetic plain in the south, and was deposited in a foreland basin setting. Structurally it is bounded by the Main Boundary Thrust (MBT) in the north and the Main Frontal Thrust (MFT) in the south (Fig. 1). The Siwalik Group consists of a 6000m+ thickness of fluvial sediments derived from the erosion of the Himalaya and Tibetan Plateau (Ulak and Nakayama, 1998). The dominant rock types are sandstone, mudstone, shale and conglomerate. The Siwalik Group is informally subdivided into three subgroups, the lower, middle and upper Siwaliks (Auden, 1935; Gautam and Rösler, 1999). The lowermost subgroup is composed mainly of mudstones and fine to medium grained sandstones, and accumulated in a meandering river environment. The middle Siwaliks consists of medium to coarse grained, mica-rich sandstone deposited from sandy meandering to sandy braided fluvial systems, whereas the upper part is predominantly composed of pebble to boulder sized conglomerates. These were deposited in gravelly braided river and alluvial fan environments (Gautam and Rösler, 1999; Beek et al., 2006).

The Siwalik succession in Nepal has been investigated in many studies, and from a variety of viewpoints (e.g.

Dhital et al., 1995; Ulak and Nakayama, 1998; Gautam and Rösler, 1999; Nakayama and Ulak, 1999; Roser et al., 2002; Huyghe et al., 2005; Beek et al., 2006; Szulc et al., 2006 and many others). However, reports of whole-rock geochemical data are limited. The purpose of this publication is to record whole-rock major and trace element analyses of a stratigraphically controlled suite of 206 sandstones and mudstones spanning the Siwalik succession in the Bakiya Khola area. Preliminary interpretation of the major element data for these samples has already been given by Roser et al. (2002), who found that elemental abundances varied up section, due to decreasing carbonate content and possibly change in provenance. Intensification of source weathering and change in fluvial style also influenced compositions. Subsequent trace element analyses for the samples are also reported here. Interpretations of the trace element data will be given in future publications.

Sample suites

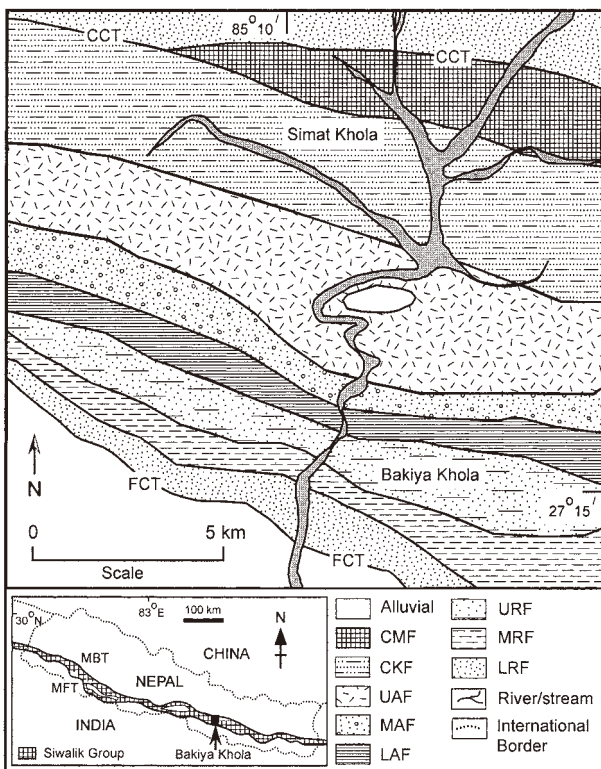
The sequence sampled is exposed mainly along the Bakiya Khola and Simat Khola areas of Central Nepal (Fig. 1). The lithostratigraphy of the Siwalik Group in the study area has been established by Sah et al. (1994) and Ulak and Nakayama (1998). The succession is divided into four formations, the Rapti, Amlekhganj, Churia Khola and Churia Mai Formations, in ascending order (Table 1). Ulak and Nakayama (1998) and Nakayama and Ulak (1999) described the lithostratigraphy and depositional environment of each

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Table 1. Stratigraphy of the Siwalik succession, Bakiya Khola area, Central Nepal (Ulak and Nakayama, 1998; Nakayama and Ulak, 1999; Ulak, 2006).

Formation	Member	Lithology	Depositional environments	Thickness (m)
Churia Mai		Poorly-sorted, clast-supported gravel and boulder	Debris flow dominated braided system	500+
Churia Khola		Well-sorted, pebble to cobble size, clast-supported conglomerate	Gravelly braided system	1100
Amlekhganj	Upper	Very coarse grained, pebbly sandstone	Shallow sandy braided system	2100
	Middle	Coarse to very coarse grained, 'pepper and salt' sandstone	Deep sandy braided system	600
	Lower	Coarse grained sandstone and gray mudstone	Sandy meandering system	340
Rapti	Upper	Fine to medium grained sandstone and variegated mudstone	Flood flow dominated meandering system	450
	Middle	Fine grained sandstone and variegated mudstone	Fine grained meandering system in a proximal setting	350
	Lower	Variegated mudstone	Fine grained meandering system in a distal setting	210+

**Fig. 1.** Generalized geological map of the Siwalik succession, Bakiya Khola area, Central Nepal (modified after Nakayama and Ulak, 1999). Abbreviations: MBT, Main Boundary Thrust; MFT, Main Frontal Thrust; CCT, Central Churia Thrust; FCT, Frontal Churia Thrust; LRF, Lower Rapti Formation; MRF, Middle Rapti Formation; URF, Upper Rapti Formation; LAF, Lower Amlekhganj Formation; MAF, Middle Amlekhganj Formation; UAF, Upper Amlekhganj Formation; CKF, Churia Khola Formation; CMF, Churia Mai Formation.

formation in the Bakiya Khola section.

The lowermost Rapti Formation has a thickness of about 1010+ m. It is well exposed in the Rapti River (Ulak, 2006), and where it is composed mainly of fine- to medium-grained gray sandstones, interbedded variegated mudstones, and siltstones. The Rapti Formation is further subdivided into lower, middle and upper members based on dominant lithotypes (Ulak and Nakayama, 1998). The lower member (210+ m thick) is mudstone-dominated, and was deposited by meandering rivers in a distal setting. The middle member (350 m thick) consists of approximately equal proportions of sandstone and mudstone deposited by meandering rivers in a more proximal setting. The upper member (450 m thick) consists of fine- to coarse-grained sandstone beds and associated dark gray to variegated mudstones; these were deposited by flood flow-dominated meandering river systems (Ulak and Nakayama, 1998; Nakayama and Ulak, 1999). Analyses were made of 85 Rapti Formation sandstones and mudstones.

The overlying Amlekhganj Formation (3040 m thick) consists of coarse to very coarse grained gray sandstones and interbedded dark gray mudstones, and is well exposed in the Dudhaura Khola near Amlekhganj village (Ulak and Nakayama, 1998). The Amlekhganj Formation is also subdivided into lower, middle and upper members. The lower member (340 m thick) consists mainly of coarse-grained "pepper and salt" sandstones with alternating gray mudstones. The "pepper and salt" appearance arises from abundance of quartz and biotite. The lower member was deposited by a flood flow-dominated sandy meandering system (Nakayama and Ulak, 1999; Roser et al., 2002). The middle member (600 m thick) is composed predominantly of coarse to very coarse-grained "pepper and salt" sandstones and gray mudstones. The depositional system in this

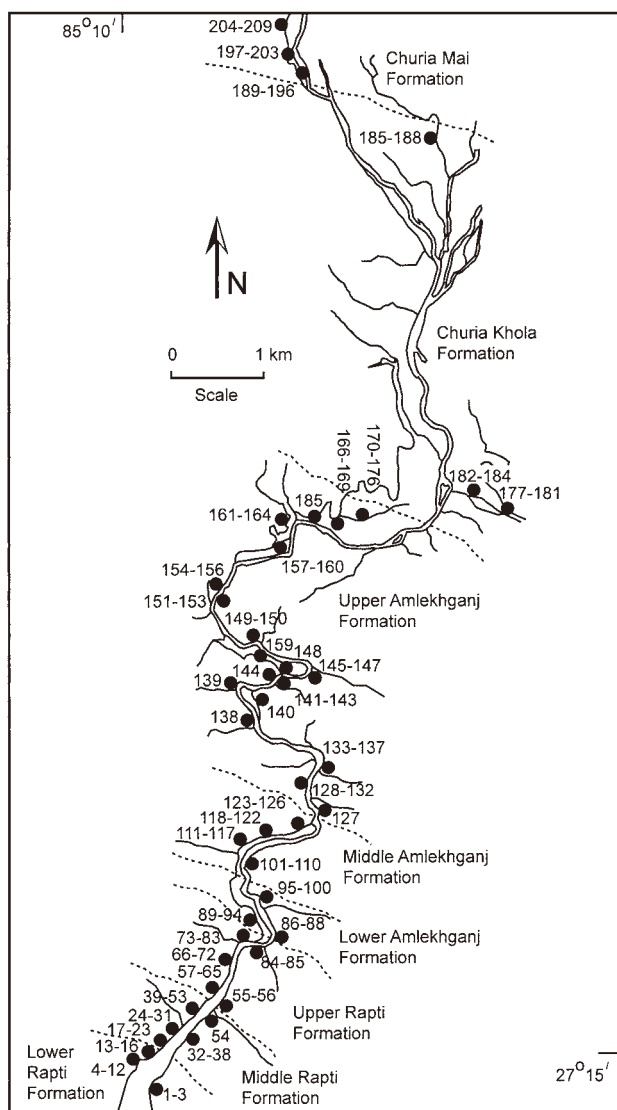


Fig. 2. Map showing location of sample sites in the Bakiya Khola area, Central Nepal, grouped by analysis number.

member was a deep and sandy braided river. The upper member (2100 m thick) of the Amlekhganj Formation is characterized by pebbly, very coarse-grained sandstones and gray mudstones, and was deposited by a shallow sandy braided system. Ninety Amlekhganj sandstones and mudstones were analyzed in this study.

The Churia Khola Formation (1100 m thick) is composed mainly of unconsolidated cobble to pebble-sized conglomerates with associated reddish-brown sandstones and gray mudstones, and accumulated from a gravelly braided river system (Ulak and Nakayama, 1998). Clasts of quartzite and limestone suggest derivation from the Lesser Himalaya. Analyses of 10 Churia Khola sandstones and mudstones are included here.

The Churia Mai Formation (500+ m thick) is composed of poorly sorted boulder conglomerates with interbedded cobbles and pebbles, along with subordinate gray sandstones and mudstones. Presence of boulder-sized Siwalik sandstone

clasts (mainly Rapti Formation) is typical. The Churia Mai Formation was deposited by a debris flow-dominated braided fluvial system (Nakayama and Ulak, 1998). Twenty-one Churia Mai Formation sandstones and mudstones were analyzed.

Methods

Samples ($n=206$) were collected from measured sections in the Bakiya Khola and Simat Khola areas of Central Nepal (Fig. 1), indexed to the paleomagnetic framework established for the area by Harrison et al. (1993). The sections were located between latitude $27^{\circ}13'N$ to $27^{\circ}28'N$ and longitude $85^{\circ}05'E$ to $85^{\circ}15'E$, southeast of Kathmandu, the capital of Nepal. Locations of the sample points are indicated in Fig. 2, grouped by analysis number. Sample preparation and analytical techniques are described in Roser et al. (2002), but are also briefly summarized below.

Samples weighing 70–100 g were manually disaggregated or chipped, and their grain size estimated visually using a grain size comparator. They were then rinsed in distilled water and dried at $110^{\circ}C$ prior to crushing. All samples were crushed in a tungsten carbide ring mill, with mills times generally of 30–45 seconds. Splits (8–10 g) of the resulting powders were stored in glass vials and dried at $110^{\circ}C$ for at least 24 h prior to determination of loss on ignition (LOI). The LOI values were determined by net weight loss after ignition in a muffle furnace at $1020^{\circ}C$ for at least 2 hours. The ignited samples from the LOI determinations were returned to glass vials and stored at $110^{\circ}C$ prior to preparation of glass fusion beads for X-ray fluorescence (XRF) analysis (ignited basis). The beads were prepared using an alkali flux (80% lithium tetraborate; 20% lithium metaborate) with a sample to flux ratio of 2:1, following the method of Kimura and Yamada (1996). The whole-rock major and trace element analyses were made using a Rigaku RIX 2000 XRF at Shimane University. Major elements and 14 trace elements (Ba through Zr in Table 2) were determined from the glass beads, using methodology and calibrations after Kimura and Yamada (1996). Four additional trace elements (La, As, Cu, Zn) were subsequently determined from pressed powder briquettes, using conventional peak over background method and calibration against eight Geological Survey of Japan rock standards.

Results

Major and trace element analyses (hydrous basis) of the Bakiya Khola samples are listed in Table 2, arranged in stratigraphic order. Average values for sandstones and mudstones in each formation are given in Table 3.

Amongst the major elements, the most conspicuous feature of the sandstone averages is an increase in SiO_2 content from the Rapti Formation (67.89 wt%) through to the Churia Khola Formation (83.07 wt%), and only slight

Table 3. Average sandstone (S) and mudstone (M) compositions, Siwalik Group, Bakiya Khola section. n = number of samples.

Fmtn Lithology n	Rapti		Amlekhganj		Churia Khola		Churia Mai	
	S	M	S	M	S	M	S	M
Major elements (wt%)								
SiO ₂	67.89	58.83	73.68	62.29	83.07	74.12	78.82	73.43
TiO ₂	0.54	0.67	0.47	0.69	0.43	0.54	0.56	0.66
Al ₂ O ₃	9.52	13.66	9.54	14.62	8.44	13.06	8.80	12.16
Fe ₂ O ₃ T	3.35	4.93	3.31	5.44	2.44	2.94	2.71	4.00
MnO	0.07	0.07	0.07	0.06	0.03	0.03	0.03	0.04
MgO	1.30	2.23	1.32	2.16	0.80	1.17	0.90	1.11
CaO	6.63	6.38	3.33	2.99	0.22	0.36	1.63	0.67
Na ₂ O	0.86	0.78	0.86	0.60	0.30	0.20	0.25	0.20
K ₂ O	2.47	3.76	2.41	4.05	2.13	3.11	2.18	2.94
P ₂ O ₅	0.06	0.08	0.07	0.08	0.05	0.03	0.05	0.04
LOI	7.19	8.24	4.41	5.97	1.99	3.71	3.57	3.70
Total	99.89	99.63	99.46	98.96	99.92	99.27	99.48	98.96
Trace elements (ppm)								
Ba	399	579	378	593	321	457	373	492
Ce	87	90	79	97	72	93	93	101
Cr	38	62	35	66	30	48	39	57
Ga	11	17	9	16	10	15	7	13
Nb	13	16	12	16	10	14	13	15
Ni	21	31	19	34	17	27	20	28
Pb	23	31	20	30	67	27	22	25
Rb	111	164	119	190	108	163	105	143
Sc	9	13	7	13	6	11	7	10
Sr	77	83	65	73	20	24	30	31
Th	17	20	18	23	14	19	19	21
V	56	88	55	96	54	69	58	81
Y	25	29	24	31	21	31	23	29
Zr	293	226	238	233	246	214	329	289
La	39	43	35	46	37	45	40	49
As	8	9	5	6	4	5	4	5
Pb	21	27	18	27	19	27	22	25
Zn	51	76	46	83	43	70	51	66
Cu	14	23	12	29	13	21	15	20

subsequent decrease in the Churia Mai (78.82 wt%). Average SiO₂ contents in the mudstones are also lower in the Rapti and Amlekhganj Formations (58.83 and 62.29 wt%, respectively) than in the Churia Khola and Churia Mai equivalents (74.12 and 73.43 wt%). As a consequence of these changes in SiO₂ content, average levels of most of the other major elements decrease stratigraphically upward, although not regularly (Table 3). Average CaO contents in the Rapti sandstones and mudstones are high (both >6 wt%), and remain significant (3 wt%) in Amlekhganj sandstones and mudstones before falling to <1 wt% in the Churia Khola Formation. A similar pattern is shown by the LOI values, reflecting the abundance of diagenetic and detrital carbonate.

Average abundances of most of the trace elements in the sandstones and mudstones tend to decrease from the Rapti to Churia Mai Formations in response to increasing average SiO₂ content (Table 3). Strontium is the exception, falling from average values of 77 and 83 ppm in Rapti sandstone and mudstone to 20 and 24 ppm in Churia Khola equivalents; levels remain low in the Churia Mai averages (30 and 31 ppm). This pattern and that of CaO and LOI suggests association of Sr with the carbonate fraction. The main feature of the trace element data is that average concentrations in the mudstones of each formation are generally greater than in their companion sandstones, suggesting the proportion of clay minerals present controls their abundances. Zirconium is an exception, being higher in the sandstone average in each formation than in companion mudstone, suggesting preferential concentration of the heavy mineral zircon in the former. Stratigraphic variation in the averages suggest that the trace element concentrations have also been affected by

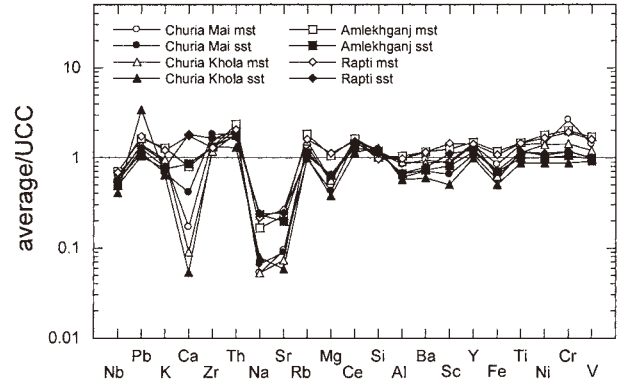


Fig. 3. Multi-element plot sandstone and mudstone averages for each formation normalized against the average upper continental crust (UCC) composition of Taylor and McLennan (1985). Elements are arranged from left to right in order of increasing abundance in average Mesozoic-Cenozoic greywacke (Condie, 1993) relative to UCC, following the methodology of Dinelli et al. (1999).

change in provenance and fluvial style, as have the major elements (Roser et al., 2002).

The sandstone and mudstone averages of each formation were normalized against the average upper continental crust (UCC) composition of Taylor and McLennan (1985) for comparative purposes. A spidergram of the elements determined from glass beads (Fig. 3; following the method of Dinelli et al., 1999) shows that many elements are present at levels at or near that of UCC, especially in the sandstones. However, there are some conspicuous departures, with marked depletion (0.05-0.5x) in Na₂O and Sr relative to UCC in all averages. Calcium is also depleted in all averages except those of the Rapti Formation, which are enriched relative to UCC. Four elements (Zr, Th, Ce, Y) are persistently enriched relative to UCC. The association suggests concentration of resistant heavy minerals, including zircon, monazite, and apatite. Although the patterns for average sandstone and mudstone in each formation are similar, the mudstones are generally enriched relative to the sandstones, especially in the "ferromagnesian" elements at the right of the figure (Sc to V), to levels above those in UCC. This reflects association with clays and the influence of sorting.

These patterns show that although the Siwalik sediments were derived from the crustal section of the Himalaya, their compositions have been significantly modified in the process. Local provenance variations, tectonism, varying intensity of weathering, sedimentary processes (e.g. sorting), differing transport times (e.g. fluvial storage) and change in fluvial style are all factors which may have contributed. The possible roles of these factors will be examined in future work.

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