

## Major and trace element abundances in the <180 $\mu$ m and sand fractions of stream sediments from the Hino River, Tottori, Japan

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### Abstract

X-ray fluorescence analyses for major elements and 14 trace elements were made of both the <180 $\mu$ m and sand fractions of 103 stream sediment samples collected from active channels of the Hino River, west Tottori Prefecture. Combined histograms of elemental abundances show clear dependence of composition on grain size, with the fine fractions containing greater abundances of all elements except SiO<sub>2</sub>, K<sub>2</sub>O, Ba, Pb, and Rb. Distributions are representative of fractions taken from river sediments, with mainly positively skewed patterns. Bimodal or polymodal distributions in the sand fractions for a number of elements (Sr, SiO<sub>2</sub>, K<sub>2</sub>O, Al<sub>2</sub>O<sub>3</sub>, CaO, Na<sub>2</sub>O, and Rb) reflect the influence of local source rocks on composition, as one of the modes is related to samples with catchments draining Mount Daisen. Variation diagrams also clearly display fractionation between the fine and sand fractions. The results show that samples originating from Mount Daisen are the most distinctive, especially with respect to high Sr abundances and relatively low Y, which reflect the adakitic nature of the source volcanics. Intermediate abundances in samples from the lower main channel reflect mixing and homogenization of detritus from two areas with contrasting lithologies, felsic granitoids and volcanics to the south, and volcanic products from Daisen to the northeast. Values of Cr are high in the upper reaches of the main channel, reflecting a strong provenance fingerprint from ultrabasic rocks. Abundances of this element then decrease steadily downstream due to dilution from Cr-poor lithotypes.

**Key words:** Geochemistry, stream sediments, <180 $\mu$ m fraction, sand fraction, Daisen products, granitoids, Hino River, Tottori.

### Introduction

The geochemical compositions of modern stream sediments mainly reflect the chemical characteristics of the source lithotypes and derived soils contained in their catchments, and the influence of weathering, transport, sorting, and post-depositional processes. Studies of stream sediments have potential application in varied geological and geoenviromental research. This article presents the results of X-ray fluorescence (XRF) analysis of both fine (<180 $\mu$ m) and sand fractions (180-2000 $\mu$ m) of 103 stream sediments collected from the Hino River in the northern San-in District.

The Hino River catchment drains an area of about 860 km<sup>2</sup>, and is located in west Tottori Prefecture, where its watershed constitutes the boundary with Shimane Prefecture. Study of the Hino River catchment and characterization of the sediments within it is of direct relevance to other work examining development of Yumigahama Peninsula related to past tataru mining, and of dispersion paths of sediments in Miho Bay.

The object of this report is to present the raw data obtained by XRF analysis, and to outline some of the general relations between abundances of elements in fractions of the bulk sediments from the Hino River. More detailed interpretation of the results will be published later. The dataset reported in this paper constitutes another step

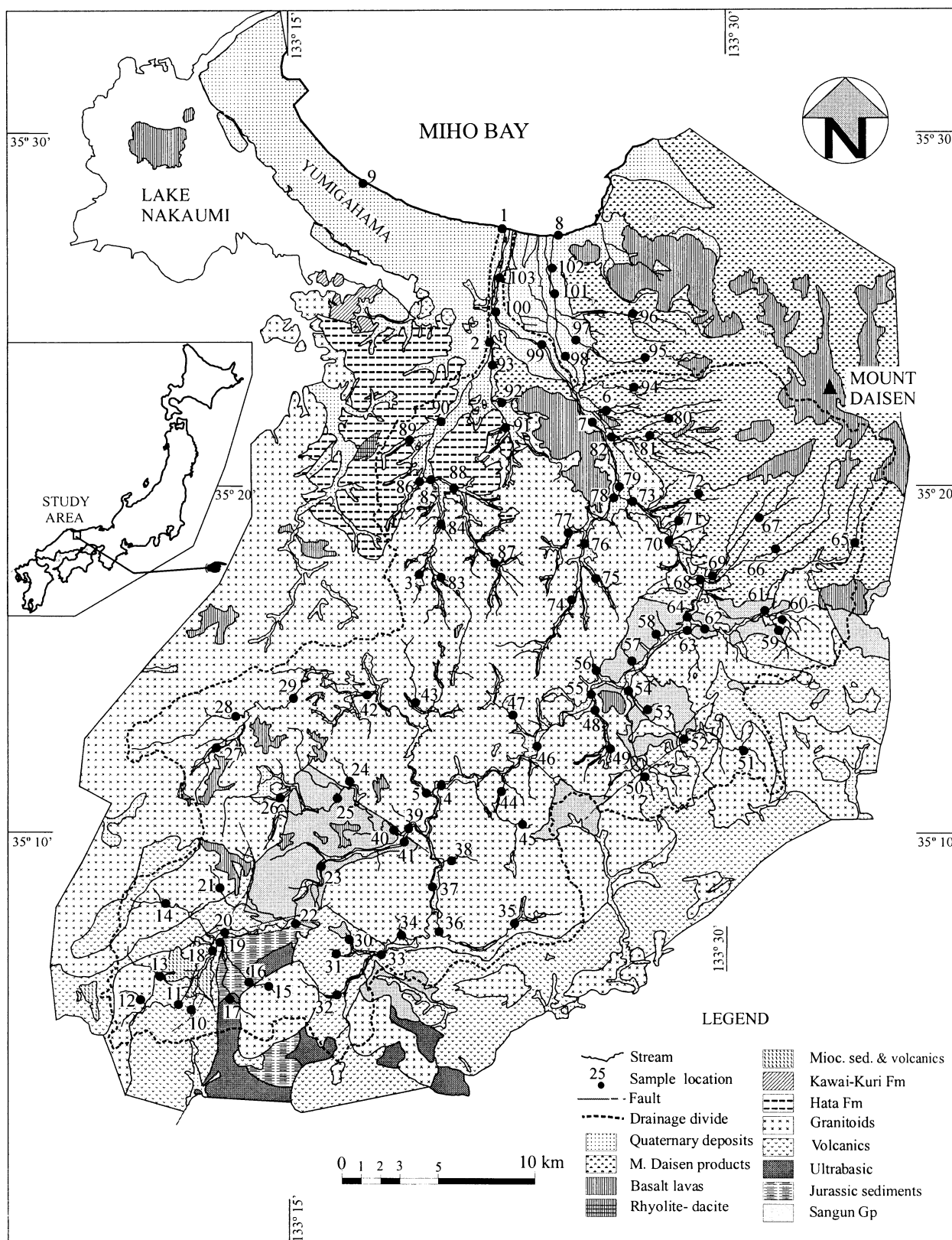
toward achieving a regional geochemical database for the northern San-in area. This work began with study of the Kando River (Ortiz and Roser, 2003; Ortiz and Roser, *this volume*), and is intended to extend to the Hii River.

### Geological Outline

Most of the catchment area of the Hino River is occupied by felsic geological units, mainly Cretaceous to Paleogene granitoids and coeval volcanic rocks (Fig. 1). The granitoids are distributed throughout the central and southern regions of the watershed, and comprise granites, granite porphyries, diorites, granodiorites and gabbros. In contrast, the coeval volcanic rocks crop out mainly on the edges of the watershed, especially towards the southwest (Japan Institute of Construction Engineering "JICE", 1984; EBGMS, 1997). Lithotypes include rhyolite, quartz andesite, andesite, and pyroclastics. Outcrops of several other less extensive units are scattered across the area (Fig. 1). These include Sangun Group ultrabasic-basic, calcareous, siliceous, and psammitic schists; Miocene-Pliocene augite basalts, olivine andesite lavas and pyroclastics; and Miocene to Jurassic sedimentary rocks (JICE, 1984).

Most of the rocks and units outlined above have been examined in varying detail during regional mapping or more specific studies of aspects such as granitoid petrogenesis (e.g., Ota, 1962; Hattori and Katada, 1964; Hashimoto, 1973; 1990; Hattori and Shibata, 1974; Murakami, 1974; Iizumi *et al.*, 1984; 2000; Research Group

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**Fig. 1.** Map showing the generalized distribution of lithotypes in the catchment of the Hino River and location of sample sites. Geology based on the 1:200,000 geological maps of Chugoku Area No 3 (Japan Institute of Construction Engineering, 1984) and Shimane Prefecture (Editorial Board of the Geological map of Shimane Prefecture, 1997).

for the Batholith in the San'in Zone, 1982; Nagao *et al.*, 1990).

The geology in the downstream reaches consists of different lithotypes. Rhyolites, andesites, and andesitic pyroclastic rocks of the Hata Formation (EBGMSP, 1997) crop out in the northwest. Conversely, volcanic products associated with Mount Daisen dominate the northeastern corner of the study area. Daisen products include andesite to quartz andesite pyroclastics, websterites, amphibolite andesites, tuff-breccias, lapilli and volcanic ash deposits, among others. Several studies have examined the distribution of Daisen rocks and their characteristics (e.g., Ota, 1959; Miura *et al.*, 1991; Morris, 1995; Kimura *et al.*, 1999, 2003).

Considerable contrasts in source rock compositions thus exist within the catchment of the Hino River. For example, about one third of the total length of the main channel (corresponding to the lower reaches) receives material almost exclusively derived from the western flank of Mount Daisen. The rest of the main channel is primarily fed from felsic source rocks. Any modification of the drainage (e.g., dams, artificial channels) at specific locations could thus lead to changes in the chemical composition of detritus reaching Miho Bay. Such modification has occurred in the past, with massive influx of granitoid waste from historic tatara mining leading to accelerated growth of Yumigahama Peninsula.

### Sampling and Sample Preparation

#### Field sampling

Field work to collect the stream sediments was carried on nine days in October and early November, 2003. Weather conditions were generally fine, and streams clear. Composite samples were collected with a plastic water scoop from active channels, and stored in plastic zip-top bags. Where possible, samples were taken from both sides of the streams over a channel length of 10-30 m.

An array of sampling sites was selected with the aim of establishing a uniform distribution that would cover all source lithologies present. Sites were identified on 1:50,000 topographic maps, according to the lithotypes present from the 1:200,000 geological maps of western Tottori and Shimane Prefectures (JICE, 1984; EBGMS, 1997).

Samples were collected from 103 sites in the Hino catchment (Fig. 1), numbered in order of collection. Some sampling sites had to be displaced from the location originally planned due to unfavorable conditions, including lack of access, disturbance by current construction projects, or simply because of paucity of sand-sized sediment. The overall sampling density was one sample per 8.3 km<sup>2</sup>.

#### Sample preparation

Bulk samples were dried at 80-90°C the day after collection. Dried bulk samples were then passed through an

8.6 mesh sieve to remove material coarser than 2 mm. Average weight of the <2 mm fractions was 1342 g. These were then split into manageable portions using a simple aluminium chute, and these later put through an 83 mesh stainless steel sieve to separate the <180 μm (fine fraction) and the 180-2000 μm (sand) fraction. For most samples, sieving of an eighth or quarter split was sufficient to provide enough <180 μm material for the XRF analysis. For six samples, half the bulk <2 mm fraction had to be used, and for one particularly well-sorted sample three quarters of the bulk material was sieved. On average, the fine fractions comprised 12.1 wt% of the <2 mm bulk fraction. Only in ten samples did the fine fraction exceed 30 wt%.

The fine and sand fractions were then separately crushed in tungsten carbide mills. The fine fractions were crushed in a small puck mill for about 10-15 seconds, in loads of 12-20 g, whereas the sand fractions were crushed in a larger ring mill for 30 seconds, in loads ranging between 50 and 150 g. Subsamples (7-10 g) of the crushed material were then stored in glass vials and dried at 110°C for at least 24 hours before determination of loss on ignition (LOI).

### Analytical Methods

Gravimetric LOI was determined for each sample from the net weight loss after ignition in a muffle furnace at 1000 °C for two hours. The ignited materials were then manually disaggregated in an agate pestle and mortar, and returned to a 110°C oven for at least 24 hours prior to the preparation of fusion beads for X-ray fluorescence analysis (anhydrous basis).

Major elements and 14 trace elements (Ba, Ce, Cr, Ga, Nb, Ni, Pb, Rb, Sc, Sr, Th, V, Y, Zr) were determined using a Rigaku RIX-2000 XRF at Shimane University. All analyses were carried out on fused glass beads prepared with an alkali flux comprising 80% lithium tetraborate and 20% lithium metaborate. The flux to sample ratio was 2:1. The analytical methods, instrumental conditions, and calibrations used were those described by Kimura and Yamada (1996). Analyses were monitored using new beads for seven GSJ and USGS standards. More detailed descriptions of the methods utilized in this study are given by Roser *et al.* (2000, 2001) and Ortiz and Roser (2003).

### Results and Discussion

Major and trace element results for the fine fractions and the sand fractions are given in Tables 1 and 2 respectively, reported on a hydrous basis. LOI values in the fine fractions were generally <10 wt%, but occasional higher values (>10%) suggest that some samples are enriched in organic matter. In contrast, LOI values for the sand fractions are considerably lower, with most <3 wt%. Only six samples have higher values, ranging up to 5.7 wt%.

Summary statistics for the fine and sand fractions are



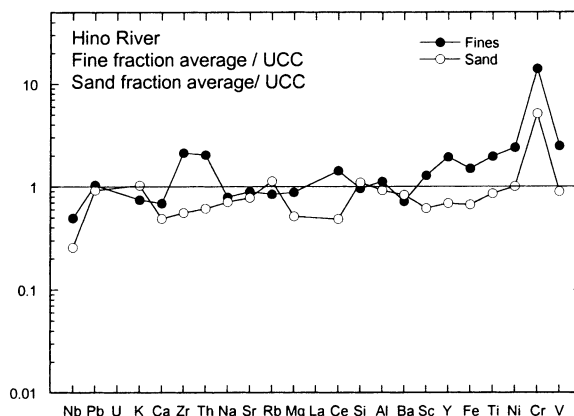


**Table 3.** Summary statistics for all <180 $\mu\text{m}$  and sand fraction samples (anhydrous normalized data).

Element	Fine Fraction N= 101				Sand Fraction N= 102			
	Mean	Min	Max	SDp	Mean	Min	Max	SDp
<i>Major elements (wt%)</i>								
SiO <sub>2</sub>	63.55	52.22	73.74	3.59	72.60	62.04	82.45	4.88
TiO <sub>2</sub>	0.98	0.36	3.18	0.42	0.43	0.13	1.43	0.22
Al <sub>2</sub> O <sub>3</sub>	17.14	12.13	22.05	1.75	14.04	8.84	21.14	2.46
Fe <sub>2</sub> O <sub>3</sub> T	7.52	2.06	22.02	2.84	3.35	0.79	10.30	1.74
MnO	0.20	0.09	0.48	0.08	0.08	0.03	0.22	0.03
MgO	1.94	0.62	9.57	1.10	1.13	0.33	5.72	0.74
CaO	2.89	1.18	5.99	1.23	2.06	0.40	6.18	1.47
Na <sub>2</sub> O	3.09	0.65	4.82	0.81	2.76	0.69	4.47	0.83
K <sub>2</sub> O	2.54	1.19	5.16	0.76	3.50	1.41	6.67	1.09
P <sub>2</sub> O <sub>5</sub>	0.15	0.04	0.58	0.07	0.06	0.01	0.14	0.03
<i>Trace elements (ppm)</i>								
Ba	396.5	287.9	574.8	63.9	456.0	293.6	647.5	77.8
Ce	91.8	10.7	340.8	49.5	31.1	6.0	61.0	9.6
Cr	493.7	13.4	8487.9	1234.1	179.6	0.4	5471.7	645.0
Ga	19.0	12.5	24.7	1.9	15.7	9.5	22.3	2.8
Nb	12.4	5.5	30.9	4.8	6.4	3.1	15.6	1.7
Ni	48.2	7.5	743.5	86.6	20.1	0.7	361.3	43.0
Pb	20.7	5.2	418.2	40.4	18.4	3.9	185.2	18.8
Rb	94.4	3.5	212.3	37.7	126.7	34.1	277.3	49.9
Sc	14.2	5.8	26.3	4.1	6.8	0.05	21.0	3.6
Sr	313.1	118.6	850.0	204.3	272.7	63.3	908.9	227.8
Th	21.9	2.7	118.8	18.7	6.6	2.5	13.9	2.5
V	149.7	33.4	696.0	88.4	53.8	0.5	304.8	46.1
Y	42.8	8.3	179.1	26.3	15.2	3.8	32.7	5.2
Zr	406.5	105.5	1072.7	203.0	106.2	50.4	203.5	31.9
Min	Minimum							
Max	Maximum							
SDp	Population standard deviation							

given in Table 3, calculated from analyses normalized to 100% to negate the effects of varying LOI. Average elemental abundances show considerable contrast between the fractions, especially for some trace elements (e.g. Cr, Zr). The average compositions of both fractions over the entire suite broadly correspond with that of average Upper Continental Crust (UCC), although some differences are also evident (Fig. 2). The fine fraction is slightly depleted in Nb, K, Ca and Ba with respect to UCC, whereas Zr, Th, and Ce abundances are a little greater. The greatest divergence from UCC in the fine fraction average is observed in the segment Sc-V, where all elements are progressively enriched, with a marked peak in Cr of more than 10 times UCC abundance. All of this group of elements are associated with mafic components. The pattern for the fine fraction average is generally linear, and is inclined from left to right, similar to the composition of average Mesozoic-Cenozoic greywacke (Condie, 1993; Dinnelli *et al.*, 1999). In contrast, the pattern of the sand fraction average is almost flat, and most elements are slightly depleted compared to UCC, especially for Nb, Ca-Sr, and Mg-Ce. Abundances also tend to increase in the segment Sc-V, as in the fine fraction, but generally remain less than or equal to UCC levels. Only the content of Cr is significantly enriched relative to UCC, although to a lesser degree than in the fine fraction.

Combined histograms (bar charts) of anhydrous-normalized elements for the fine and sand fractions display clear dependence on grain size, although overlap is considerable. The fine fractions tend to have greater abundances of most elements, except for SiO<sub>2</sub>, K<sub>2</sub>O, and to a lesser extent Ba, Pb, and Rb, which are enriched in the sand fraction (Fig. 3 and 4). Na<sub>2</sub>O (Fig. 3), Sr, and Pb show



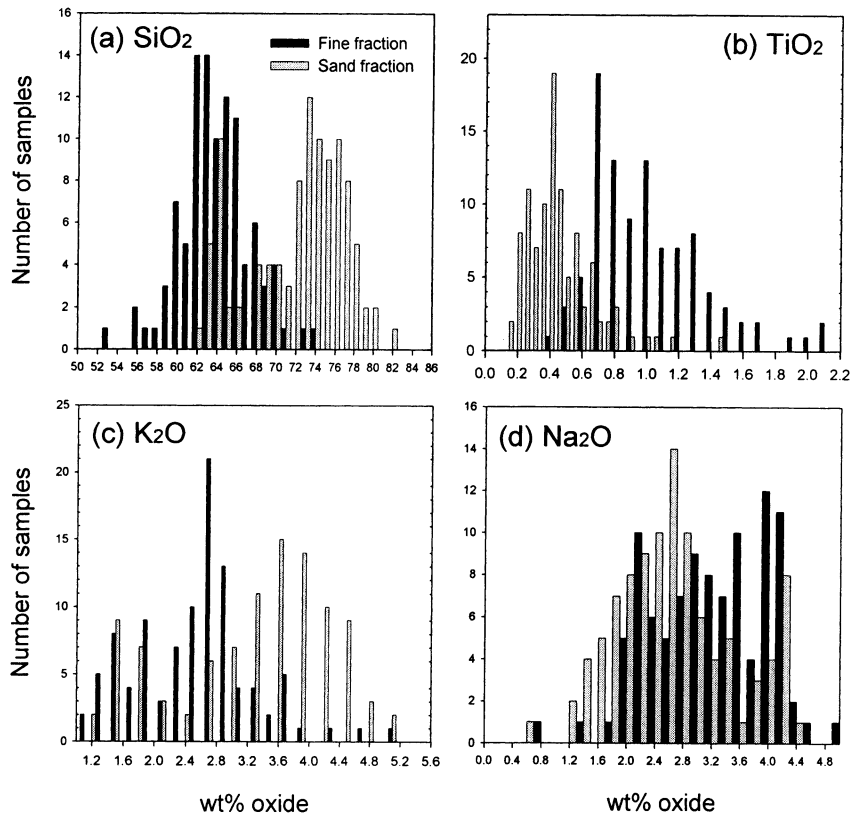
**Fig. 2.** Multi-element plot showing the average composition (anhydrous normalized) of the fine (<180 $\mu\text{m}$ ) and sand fractions from the Hino River (data from Tables 1 and 2) normalized against the Upper Continental Crust (UCC) average of Taylor and McLennan (1985). Elements are arranged from left to right in order of increasing normalized abundance in average Mesozoic-Cenozoic greywacke (Condie, 1993) relative to UCC, following the methodology of Dinelli *et al.*, (1999). The major elements are normalized as oxides.

relatively little contrast in the ranges of their distributions in the fractions. However, statistical tests suggest that the means of all elements except Sr and Pb differ significantly (95% confidence level) between the two fractions.

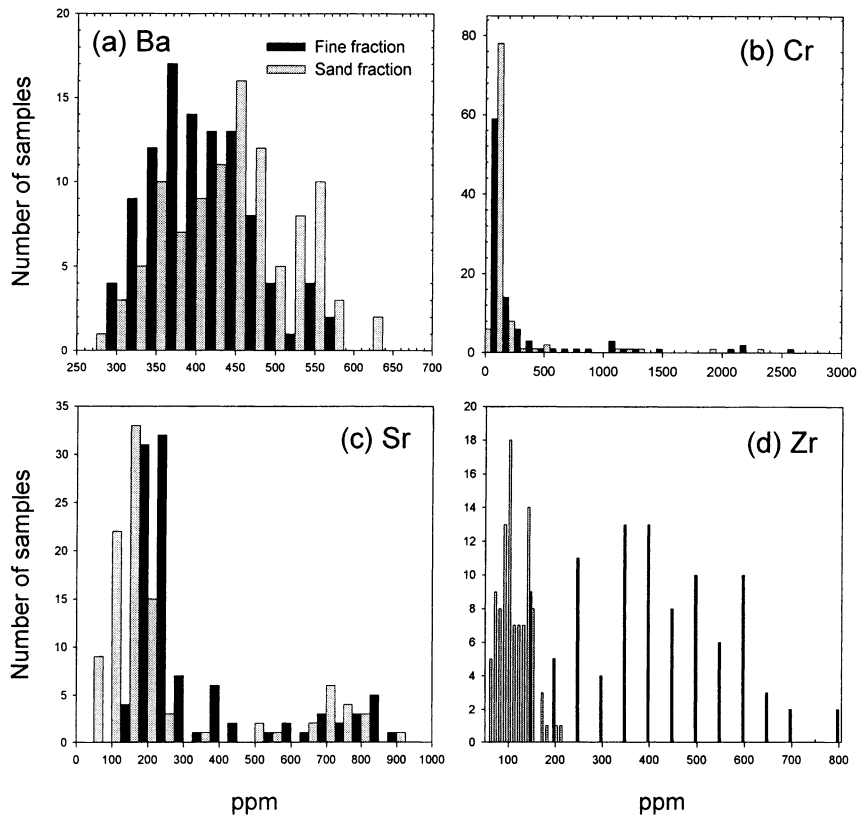
#### Fine Fraction

Most of the major elements have variable distributions and contain anomalous values. Distribution of SiO<sub>2</sub> (Fig. 3a) is relatively normal except for lower values in several samples, coupled with higher abundances of other elements. This is especially marked in sample HN 23, which has the lowest SiO<sub>2</sub> content (52.2%) due to enrichment in Fe<sub>2</sub>O<sub>3</sub> (22.02%), TiO<sub>2</sub> (3.18%), and MnO (0.38%), probably as a result of Fe-Ti-Mn oxide heavy mineral concentration or authigenic Fe-Mn-O crust material. TiO<sub>2</sub> (Fig. 3b), Fe<sub>2</sub>O<sub>3</sub>T and MnO are markedly skewed to higher values. MnO and P<sub>2</sub>O<sub>5</sub> show the same characteristic, but to a lesser extent. The remaining distributions are more irregular, with K<sub>2</sub>O (Fig. 3c) and Na<sub>2</sub>O (Fig. 3d) possibly bimodal, whereas Al<sub>2</sub>O<sub>3</sub> and CaO are polymodal.

Trace elements have also variable abundances. Ba (Fig. 4a) and Ga resemble normal distributions, with few anomalous values. Cr (Fig. 4b), Ce, Ni, Pb, Th, V, and Y are strongly skewed to higher values, with some of the most anomalous values recorded in sample HN 23 (e.g. Cr 6629 ppm; V 696 ppm; Zr 578 ppm). Several other samples also have extreme values, such as HN 17 (Cr 8488 ppm; Ni 744 ppm), HN 87 (Zr 1073 ppm), and HN 25 (Zr 1048 ppm). The distribution of Sr (Fig. 4c) is clearly bimodal, with a marked contrast between one group with relatively low values (100 to 400 ppm) and another with high values (500-900 ppm). The latter group is associated with samples collected from streams draining Mount Daisen. The



**Fig. 3.** Examples of combined histograms of major element abundances (anhydrous normalized data) in the  $<180\ \mu\text{m}$  and sand fractions, Hino River. (a)  $\text{SiO}_2$  - normal distribution; (b)  $\text{TiO}_2$  - skewed to higher values; one sample with  $>2.2$  wt% not plotted; (c)  $\text{K}_2\text{O}$  - bimodal, two samples  $>5.6$  wt% not plotted; (d)  $\text{Na}_2\text{O}$  - distinctly polymodal.



**Fig. 4.** Examples of combined histograms of trace element abundances (anhydrous normalized data) in the  $<180\ \mu\text{m}$  and sand fractions, Hino River. (a) Ba - normal distribution; (b) Cr - strongly skewed to higher values, five samples with 3400-8500 ppm not plotted; (c) Sr - clear bimodal distributions and (d) Zr - polymodal, five samples with 845-1080 ppm not plotted.

remaining trace elements analyzed (Zr (Fig. 4d), Nb, Rb, and Sc) have comparatively wide ranges and polymodal distributions.

#### *Sand fraction*

The most notable feature of the patterns among the major elements is the relatively high number of bimodal distributions (five in total), all of which identify a group of samples from streams that drain catchments consisting almost exclusively of products from Mount Daisen. For SiO<sub>2</sub> (Fig. 3a) and K<sub>2</sub>O (Fig. 3c) this mode is found at lower values, whereas for Na<sub>2</sub>O (Fig. 3d), Al<sub>2</sub>O<sub>3</sub>, and CaO it occurs at higher values. In all cases, contrast between the two major modes is evident, and reflects the strong geochemical signature that volcanic products from Mount Daisen imprint to the sediments. The remaining major elements (TiO<sub>2</sub> (Fig. 3b), Fe<sub>2</sub>O<sub>3</sub>, MnO, MgO, and P<sub>2</sub>O<sub>5</sub>) exhibit variable and strongly right-skewed distributions. As in the fine fraction, several samples have lower SiO<sub>2</sub> values due to enrichment in Fe<sub>2</sub>O<sub>3</sub>, TiO<sub>2</sub>, and MnO, and again this is most marked in sample HN 23.

Among the trace elements, a number of elements (Ba (Fig. 4a), Ce, Ga, and Sc) have normal distributions with relatively few anomalous values. In contrast Cr (Fig. 4b), Nb, Ni, Pb, and V are moderately to strongly skewed to higher values. Sr (Fig. 4c), Rb and to a lesser extent Y display bimodal distributions, with one of the modes associated with Daisen products. The contrast between the modes is particularly marked for Sr. Finally, abundances of Zr (Fig. 4d) and Th vary considerably, and distributions are polymodal.

#### *Anomalous values*

Elemental abundances in both fractions generally have positively skewed distributions. Most of the elements analyzed are not prone to disturbance from human activity, but several elements have elevated values (>2 s.d.) that could be considered exceptional. Anomalous values are most commonly observed for Cr, from samples (e.g., HN 17-20, 22, 23) containing basic and ultrabasic rocks in their catchments. These values may have been increased by activity related to chromium mining in the area. Some of this group of samples also contain elevated levels of Ni. With one exception, abundances of the potentially environmentally sensitive element P<sub>2</sub>O<sub>5</sub>, are low and within the range expected for the source lithotypes. There is thus no clear evidence in the sediments for anthropogenic inputs (e.g. via fertilizers). The compositions observed suggest clear association with the characteristics of source rock lithotypes and derived products. Greater values for a group of elements including Fe, Mn, Ti, Cr, Ni, Ce, Th, V, Y, and Zr are very likely related to concentrations of high density accessory minerals (see Ortiz & Roser, 2003). This will be verified by future work.

The bimodal and polymodal distributions observed

reflect the control of rock source in the composition of derived sediments. As already stated, this is particularly distinctive in the sand fraction, including the bimodal patterns of SiO<sub>2</sub>, K<sub>2</sub>O, Al<sub>2</sub>O<sub>3</sub>, CaO, Na<sub>2</sub>O, and Sr, elements for which one of the modes is clearly associated with samples collected from catchments draining Mount Daisen.

#### *Classification of sample sites by source categories*

As described above, the geology of the Hino river catchment differs spatially, with felsic igneous and volcanic rocks dominating the central and southern parts, volcanic products from Mount Daisen in the northeast, and volcanic rocks of the Hata Formation in the northwest. Following the methodology adopted by Ortiz and Roser (2003) for the Kando River, sample localities from the Hino watershed were divided into four categories according to the characteristics of their main source rocks. The sand fractions include an additional category consisting of two samples (HN 8, 9) collected from the beach in Miho Bay. The <180  $\mu$ m fraction in these two samples was too small to be analyzed.

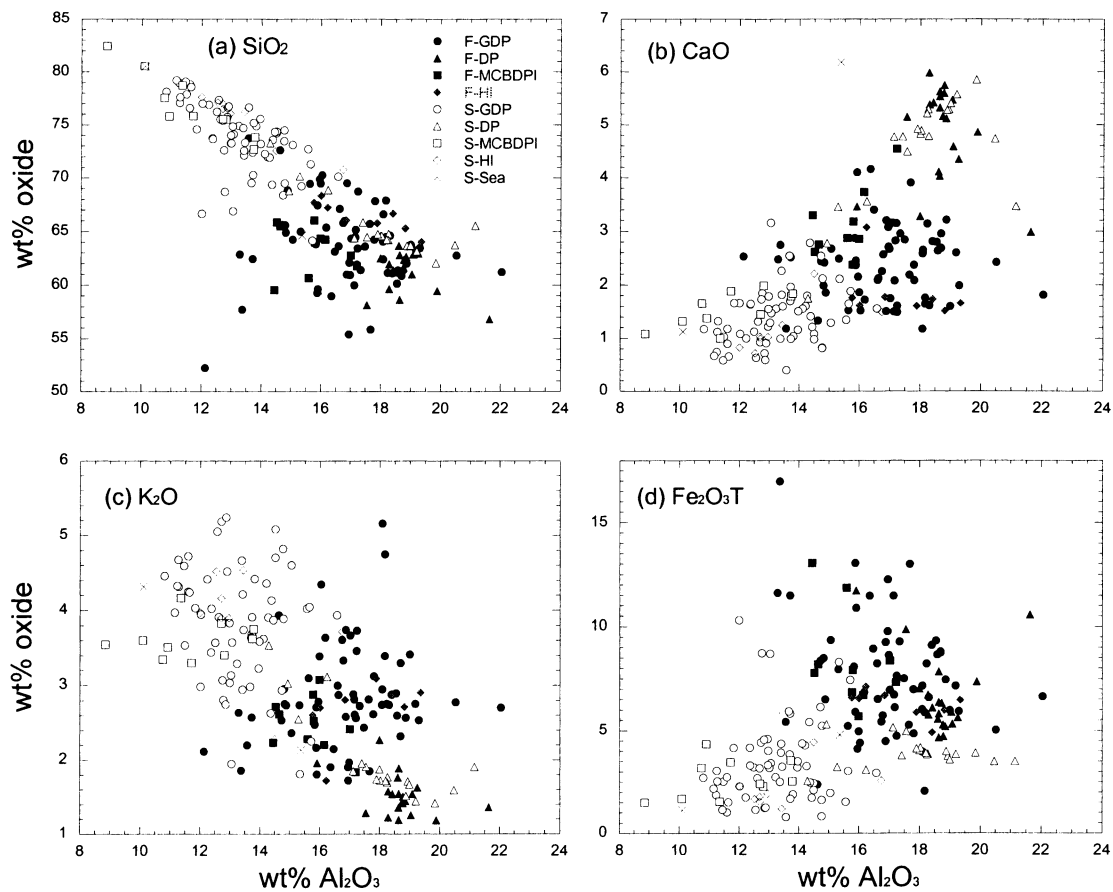
The main categories are:

- (1). Granitoid-dominated products (GDP) and felsic volcanics, covering a vast area and including more than 61% of the total sample sites;
- (2). Daisen products (DP), consisting of samples from catchments dominated by volcanic products from Mount Daisen, although sites HN 61, 62, 69, and 71 also contain other lithologies (e.g., granitoids, psammitic schists);
- (3). Main channel below Daisen product input (MCBDPI); namely HN 1, 7, 64, 68, 70, 73, 79, 99, 100, and 103;
- (4). Hata Formation input (HI); sites HN 2, 85, 86, 88, 89, 90, 91, and 93.

Simple element-Al<sub>2</sub>O<sub>3</sub> variation diagrams constitute a useful tool to illustrate differences between abundances of elements according to the above categories. The first feature observed is that although there is significant overlap and scatter, the distributions of elements show broad linear trends from the sand fractions (concentrated at lower abundances of Al<sub>2</sub>O<sub>3</sub>) toward the fine fraction. These linear trends are especially clear for SiO<sub>2</sub> (Fig. 5) and Ga (Fig. 6). Distribution of the fractions with respect to Al<sub>2</sub>O<sub>3</sub> corresponds with the histograms of elemental abundances (Figs. 3 and 4), and illustrate the dependence of chemical composition on grain size as reported by several authors (e.g., Fralick & Kronberg, 1997; Vital & Stattegger, 2000).

The distributions of the major elements generally overlap, but samples derived from Daisen products tend to be distinguishable. This is especially clear for CaO (Fig. 5b), and K<sub>2</sub>O (Fig. 5c), for which samples in both fractions from that category have higher and lower values (respectively) compared to the other categories. Fe<sub>2</sub>O<sub>3</sub> (Fig. 5d), TiO<sub>2</sub>, and MnO abundances in the fine fractions tend to be greater





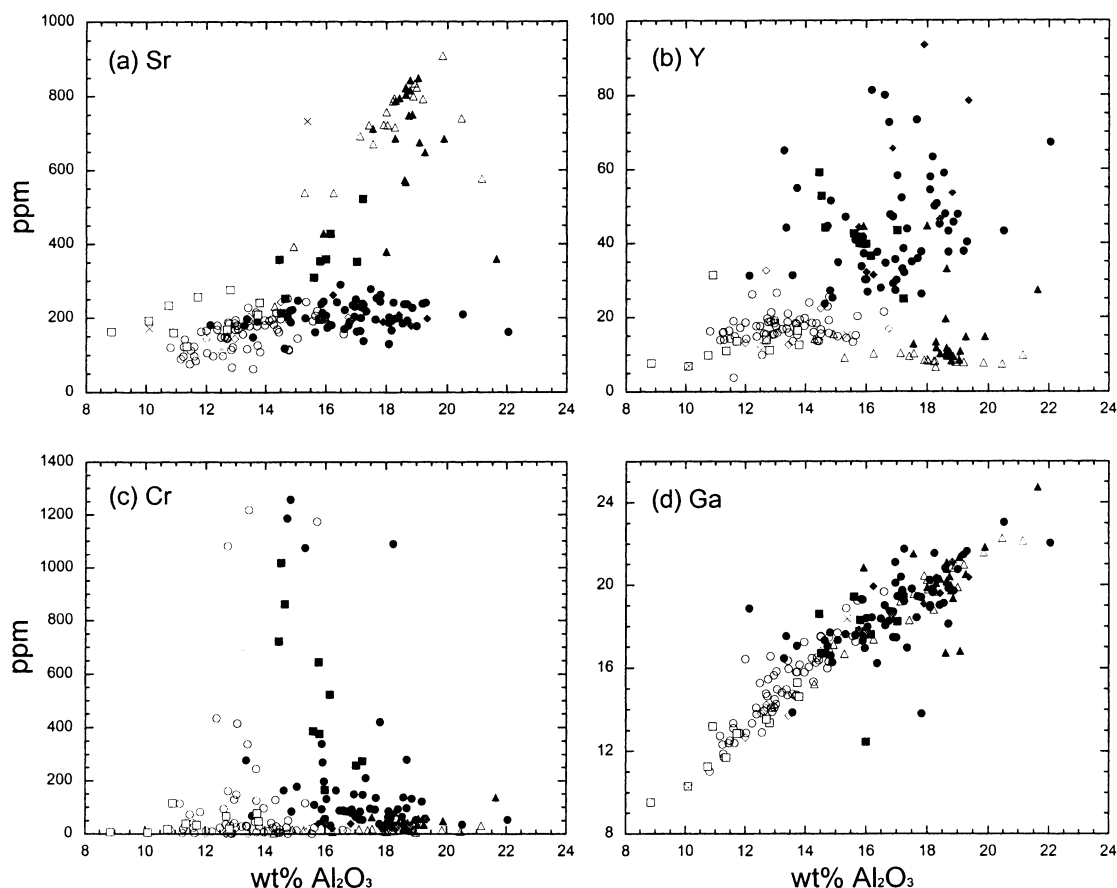
**Fig. 5.** Examples of major element- $\text{Al}_2\text{O}_3$  variations in the fine and sand fractions (anhydrous normalized data), Hino River, according to main source lithotypes. F -  $<180\ \mu\text{m}$  (fine) fraction; S - sand fraction; - GDP = granitoid-dominated, - DP= Daisen products; - MCB DPI = main channel below Daisen product input - HI= Hata input; - sea = beach sands. (a)  $\text{SiO}_2$  - linear trend between sand and fine fractions, lower abundances in samples derived from Daisen products; (b) CaO - greater abundances in samples from Daisen products; (c)  $\text{K}_2\text{O}$  - lower abundances in samples from Daisen products, two samples with  $>6$  wt% not plotted; and (d)  $\text{Fe}_2\text{O}_3\text{T}$  - abundances in the fine fractions greater than the sands in each category, one sample with  $>18$  wt% not plotted.

than the sands in each category. Conversely, for  $\text{Na}_2\text{O}$ ,  $\text{MgO}$ , and  $\text{P}_2\text{O}_5$  only sand fractions derived from Daisen products have distinctively higher values, whereas for  $\text{SiO}_2$  (Fig. 5a) the same samples have lower values. Samples derived from Hata Formation rocks generally show little contrast with samples from granitoid-dominated sites, and tend to plot towards lower values only for CaO (Fig. 5b) and MgO, especially the fine fraction. Although scatter is considerable, samples from the main channel below Daisen product input tend to have values intermediate between Daisen products and the granitoid-dominated group, reflecting mixing of detritus in the lower reaches of the river. Beach sample HN 8, collected from the eastern part of Miho Bay, has high CaO (Fig. 5b) in its sand fraction, possibly reflecting abundance of Daisen detritus at that site.

Trace elements behave in a similar way to the major elements; and hence the most distinctive features are related to the distribution of samples derived from Daisen products. Both fractions of samples from this category plot well apart from the remaining groups, especially for Sr (Fig. 6a), with

higher values; and for Y (Fig. 6b), Rb, Th, and to a lesser extent Ba and Pb, all of which have lower values than the other groups. The higher values for Sr and lower for Y reflect the adakitic nature of the Daisen volcanic products, as described in a number of studies (e.g., Morris, 1995; Kimura *et al.*, 2003). For Cr (Fig. 6c), Ce, Nb, Zr, and to a lesser extent for V, only the fine fractions of samples derived from Daisen products have distinctive lower values. Daisen-derived sand fractions overlap with the remaining categories. Conversely, for Ga (Fig. 6d) and to some degree for Sc, the Daisen sand fractions tend to have higher values, and equivalent fine fractions show little contrast with the other categories. Samples derived from Hata sources are again scattered among the granitoid-dominated data, except for Cr (Fig. 6c) and Ni, for which concentrations tend to be lower in the fine fractions. This suggests the geochemical signatures of Hata inputs do not differ greatly from those derived from granitoid-dominated catchments.

Although overlap and scatter are significant, trace element abundances in samples from the lower main



**Fig. 6.** Examples of trace element- $\text{Al}_2\text{O}_3$  variations in the fine and sand fractions (anhydrous normalized data), Hino River, according to main source lithotypes. Symbols as in Fig. 5. (a) Sr - markedly greater abundances in samples from Daisen products; (b) Y - lower abundances in samples from Daisen products, four samples with >100 ppm offscale; (c) Cr - extreme abundances in some samples; 12 samples with >1400 ppm offscale (d) Ga - linear trend between fine and sand fractions.

channel tend to be intermediate between those of the granitoid- and Daisen-derived products. This represents mixing in the lower reaches of the main channel of Hino River of detritus originating from these two contrasting areas, the largest of which is mainly felsic in composition. Consequently, the primary source signatures of these two sources are obscured.

Abundances of Cr in the upper main channel are greater than those in the secondary drainages in all categories, excepting sites directly draining ultrabasic rocks (16, 17, 19). Levels in the main channel decrease relatively regularly downstream, reflecting dilution of Cr-rich detritus from the ultrabasics with Cr-poor detritus from all other lithotypes. This is most marked for the <math><180\mu\text{m}</math> fraction, but is also evident in the sands.

### Conclusions

The chemical compositions of sediments from the Hino River system are representative of the nature of their source lithotypes. Dependence of composition on grain size is displayed by contrasting elemental distributions of the <math><180\mu\text{m}</math> and sand fractions. The distributions reflect the

influence of different lithotypes, with especially distinctive contrasts produced in sediments mainly derived from Mount Daisen. Sediments originating from Daisen record adakitic signatures from such products. Elevated concentrations for a number of elements (e.g. Zr, Cr) are likely related to local concentrations of heavy minerals. The intermediate chemistry of sediments collected from the main channel below the first input of Daisen products reflects mixing and homogenization of detritus in the lower reaches of the Hino River.

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#### (要 旨)

E. Ortiz · B.P. Roser, 2004, 鳥取県、日野川の河川堆積物の 180  $\mu\text{m}$  以下と砂部分の堆積物の主成分ならびに微量成分組成. 鳥根大学地球資源環境学報告, **23**, 27-37

鳥取県、日野川の現河床から採取した 103 の試料を 180  $\mu\text{m}$  より細粒な堆積物と砂部分に分け、それぞれについて XRF 分析による主成分元素と 14 の微量成分元素の解析を行った。元素含有量のヒストグラムからは明らかに組成が粒径に依存していることが示された。とくに細粒な堆積物では  $\text{SiO}_2$ ,  $\text{K}_2\text{O}$ , Ba, Pb, Rb を除くすべての元素がより多く含まれていた。細粒堆積物、粗粒堆積物それぞれの分布のパターンは主に正の歪みを示し、それぞれにはっきりとした違いが見られた。多くの元素(Sr,  $\text{SiO}_2$ ,  $\text{K}_2\text{O}$ ,  $\text{Al}_2\text{O}_3$ , CaO,  $\text{N}_2\text{O}$ , Rb)について、砂部分は二峰性または多峰性のパターンを示す。これは局所的な供給源の影響を反映している。とくにモードのうちの 1 つが大山を集水域にもつサンプルと関係している。組成の変化を表したダイアグラムは細粒部分と砂部分で別が起きていることをはっきりと示している。この結果から、大山からのサンプルは特に Sr 含有量が高く、Y 含有量がかなり低い、起源火山岩のアダカイト質の特徴を反映していることが示された。下流の流路堆積物からは中間的な組成が得られた。そのことは岩相の対照的な 2 つの地域、すなわちフェルシクな花崗岩類と火山岩からなる南部と大山からの火山噴出物からなる北部からの碎屑物が混合、均質化していることを反映している。Cr の値が上流域で高く、このことは超塩基性岩からの供給を強く示唆している。この元素量は下流に向かって安定的に減少するが、それは Cr をあまり含まない碎屑物による希釈を表している。