

Article

## Major and trace element abundances in Gondwana Group sediments from GDH-45, Khalaspir basin, Bangladesh

Dhiman Kumer Roy\* and Barry P. Roser\*

### Abstract

Subsurface Permo-Carboniferous sediments of the Gondwana Group occur in the Khalaspir basin of Bangladesh. Similar successions crop out in South America, Australia, South Africa and India. This report presents whole-rock major and trace element X-ray fluorescence analyses of 93 sandstones and mudrocks from drill hole GDH-45 in the Khalaspir basin. For convenience the succession is divided into three units (Units 1 to 3, from oldest to youngest), based on lithofacies. Geochemical contrasts occur with the succession, between both units and lithotypes. Average  $\text{SiO}_2$  contents fall in Unit 3, and hence  $\text{Al}_2\text{O}_3$  contents increase. Average  $\text{CaO}$  contents are less than the upper continental crust, suggesting minimal effect of secondary calcite cements.  $\text{Na}_2\text{O}$  and  $\text{K}_2\text{O}$  and most of the other major elements generally decrease upward. In all units average trace element concentrations are greater in mudrocks than in sandstones, indicating control of their abundances by clay mineral content. Average large ion lithophile element (Ba, Rb, Sr) contents decrease from Unit 1 to 3. Other trace elements (Nb, Zr, Y, Cr and Ni) show little systematic variation between units in mudrocks, but some contrasts in sandstones. The data will be used in future work to unravel the role that climatic, tectonic and other factors play in the composition of the Gondwana succession in Bangladesh, and for comparison and correlation with other Gondwana basins elsewhere.

**Key words:** Gondwana Group, geochemistry, sandstones, mudrocks, Bangladesh

### Introduction

Gondwana Group sediments of Permo-carboniferous age on several continental blocks (South America, Australia, South Africa and India) record Gondwanaland continental breakup, widespread glaciations, and dramatic climatic, environmental and biosphere changes (Suttner and Dutta, 1986; Martini, 1997; Scheffler *et al.* 2003; Catuneanu *et al.* 2005; Stephenson *et al.* 2007). The Permo-Carboniferous Gondwana basin sediments of northwestern Bangladesh (Fig. 1A) have characteristics similar to the other Gondwana basins, and have been correlated based on spore-pollen study (Akthar, 2001). The Bangladesh Gondwana Group sediments have not as yet been correlated with specific formations in the other regional Gondwana basins in India, South America, South Africa and Australia. The Bangladesh Gondwana succession consists of sediments deposited in lacustrine and braided to meandering river depositional environments created by changes in tectonic and climatic conditions (Islam *et al.*, 1992; Akthar, 2001). These sediments thus contain potentially important information on the tectonic and climatic evolution of the region at the time. The total thickness of the Gondwana Group in Bangladesh is estimated to be about 1400+ m (Khan, 1991).

Geochemical compositions of sediments can be used to infer the influence of tectonic and climatic factors, and assist with basin-wide correlation. Geochemical characteristics

of sediments in relation to provenance (Roser and Korsch, 1988; McLennan *et al.* 1993), weathering and climate in the source (Nesbitt and Young, 1982; Rieu *et al.*, 2007; Clift *et al.*, 2008), hydrodynamic sorting during transportation (Ohta, 2008), tectonic setting of the source area and depositional basin (Bhatia, 1983; Roser and Korsch, 1986), and post-depositional K-metasomatism (Fedo *et al.*, 1995) have been well documented in the literature. Numerous processes thus influence the composition of sediments from source to sink, but they also leave characteristic geochemical imprints. Therefore, the geochemical compositions of sediments give the opportunity to scrutinize the role of these factors, and aid the reconstruction of past sedimentary environments.

The Gondwana Group sediments of Bangladesh have been investigated from several viewpoints, including general description and characterization (Islam *et al.*, 1992), palynomorphs (Akthar, 2001), provenance based on modal composition (Rahman and Ahmed, 1995; Hossain *et al.*, 2000), depositional environment (Islam *et al.*, 1992; Hossain *et al.*, 2002), and geochemical composition (Islam *et al.*, 2004). The study of Islam *et al.* (2004) included modal, major and trace element, and rare-earth element (REE) data for samples from the nearby Barapukuria basin. Although these data are valuable, the number of samples analyzed was relatively small ( $n=19$ ), and all were either medium or fine-grained sandstones.

This present account reports whole-rock major and trace element analyses of Gondwana Group sandstones and mudrocks (siltstones and mudstones) from drill hole GDH-

\* Dept. of Geoscience, Shimane University, 1060 Nishikawatsu, Matsue 690-8504, Japan

45 from the Khalaspir basin. This data for the Bangladesh succession adds to a growing database for Gondwana basins in adjacent India and in other continental blocks. The data contained in this report will be interpreted and discussed in a future publication, in the light of climatic fluctuation during the Carboniferous-Permian period and its significance.

### Geological Outline

The Permo-Carboniferous Gondwana sediments analyzed here were intersected in drill hole GDH-45 in the Khalaspir basin (Islam *et al.*, 1992). Gondwana sediments are not exposed at the surface in Bangladesh. GDH-45 provides a continuous 808 m (290-1098 m) section through the Gondwana succession (Fig. 1B). Although the Gondwana Group in Bangladesh has not been divided into formal formations, we have informally divided the succession into three units (Unit 1, 2, and 3, in ascending order), based on facies and depositional environment change (Fig. 2).

The Khalaspir sediments were deposited in an isolated asymmetric and fault-bounded intracratonic half-graben basin floored by Paleoproterozoic basement (Islam *et al.*, 1992; Hossain *et al.*, 2007; Frielingsdorf *et al.*, 2008). Details of lithology and depositional environment are presented by Islam *et al.* (1992). The lowermost Unit 1 (U1, 753-1098 m) is composed of feldspathic sandstones, sandstones, siltstones, and mudstones, and was deposited in a lacustrine or back swamp environment (Islam *et al.*, 1992). Unit 2 (U2, 640-753 m) consists mainly of conglomerates and sandstones that were deposited in a braided river environment, whereas the overlying Unit 3 (U3, 290-640 m) is composed mainly of sandstones and mudstones, along with coals in its uppermost part (Islam *et al.*, 1992; Hossain *et*

*al.*, 2002). Unit 3 was deposited from braided to sinuous streams (Hossain *et al.*, 2002).

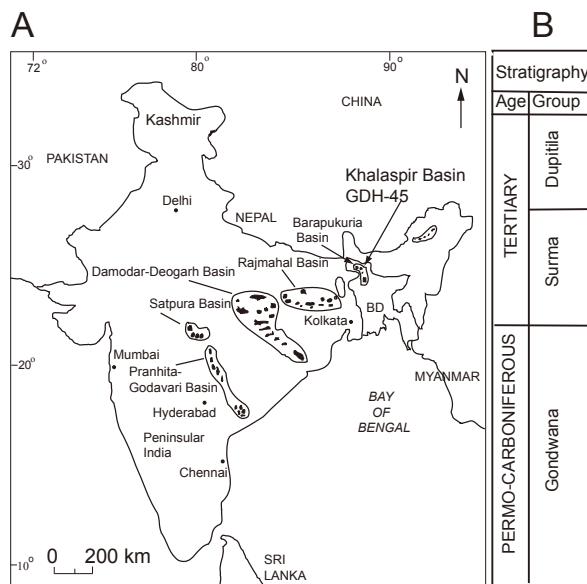
### Sampling and Analysis

#### Sampling

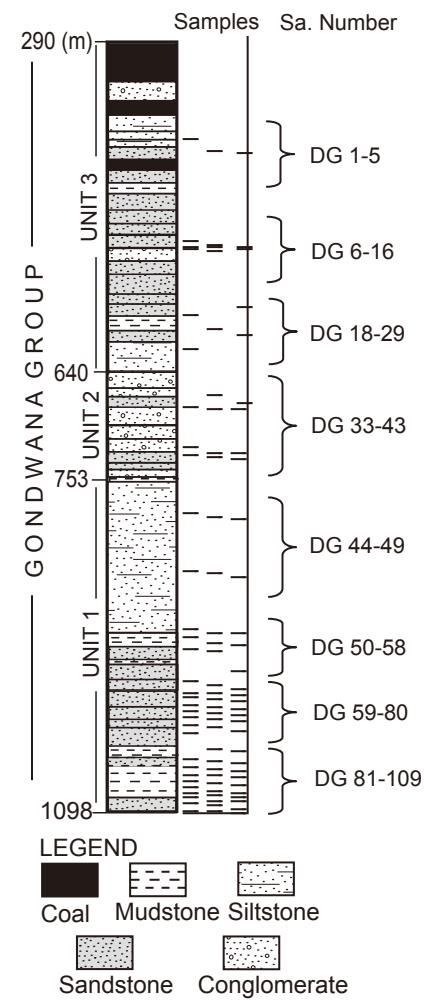
Depths of the sandstone and mudstone samples within the drillcore are listed in Table 1, and illustrated in Fig. 2. The number of sample collected in each unit was based on the availability from drill core. Ninety-three samples (41 sandstones and 52 mudrocks) were collected from GDH-45, with 66 taken from Unit 1, 11 from Unit 2, and 16 from Unit 3. Unit 1 is thus somewhat over-represented with respect to its thickness, and Unit 3 somewhat under-represented.

#### Analytical procedures

The drill core samples were broken manually into small chips (<1 cm), washed with deionized water to remove surface dust, and then dried in an oven at 110°C. The dried samples were then powdered in a Rocklabs tungsten carbide ring mill, with mills times generally <30



**Fig. 1.**(A). Location of Gondwana basins in the Indian subcontinent (after Singh and Singh, 2004; Islam, 2009). Arrow indicates the location of GDH-45 in the Khalaspir basin. BD=Bangladesh. (B) Stratigraphic succession in GDH-45.



**Fig. 2.** Stratigraphy, lithology and sample positions in drill hole GDH-45, Khalaspir basin, Bangladesh.

**Table 1.** Whole rock major (wt%) and trace element (ppm) analyses of Gondwana Group sandstones and mudrocks from GDH-45, Khalaspir basin, Bangladesh.

Index: SanNr-sample number, Lith-lithology (see footnote for codes), total iron as Fe<sub>2</sub>O<sub>3</sub>, LOI-Loss on ignition. Note that data are tabulated on an *anhydrous normalized basis*. The normalizing factors were applied to both the major and trace element data. Original anhydrous analysis is total (Total\*) and LOI data are included to allow recalculation to a hydrous basis if desired.

SanNr	Depth (m)	Lith	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Fe <sub>2</sub> O <sub>3</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	Ba	Ce	Cr	Ga	Nb	Ni	Pb	Rb	Sc	Sr	Th	V	Y	Zr	LOI	Total*	
<b>Gondwana Group</b>																													
<b>Unit 3</b>																													
DG-1	396	fst	69.76	1.00	22.16	5.73	0.05	0.33	0.12	0.13	0.68	0.03	142	88	131	21	20	37	16	29	28.7	38	7.3	154	41	306	9.31	99.61	
DG-4	409	fst	59.83	0.73	24.90	13.30	0.12	0.29	0.21	0.11	0.37	0.14	161	150	83	25	18	34	12	26	23.8	107	5.9	138	43	156	12.21	99.37	
DG-5	410	mst	62.36	0.85	23.49	11.05	0.12	0.65	0.26	0.16	0.21	0.09	142	89	118	26	20	34	10	36	16.0	72	10.7	130	26	192	12.02	99.73	
DG-6	502	mst	62.23	1.16	31.80	1.83	0.00	0.48	0.13	0.13	2.21	0.03	465	110	129	38	29	35	9	35	20	88	55	14.5	160	25	171	11.36	99.96
DG-7	507	mst	60.88	1.44	32.60	1.90	0.00	0.53	0.15	0.13	2.33	0.04	18	38	55	14	9	35	9	22	20.6	38	3.8	55	18	110	13.07	98.68	
DG-8	508	mst	62.23	1.06	30.37	3.35	0.02	0.67	0.26	0.10	1.89	0.05	262	140	139	36	25	60	16	79	90	29.8	90	12.6	170	53	291	13.01	99.77
DG-10	509	fst	73.66	0.65	20.43	2.37	0.02	0.55	0.20	0.19	1.49	0.05	125	83	98	20	15	24	13	52	8.5	61	7.8	84	19	334	8.44	99.90	
DG-11	509	fst	70.26	0.83	23.29	2.59	0.02	0.55	0.68	0.13	1.59	0.06	152	102	120	26	19	30	14	60	10.6	77	8.7	117	23	358	8.04	99.47	
DG-12	510	mst	54.37	0.56	15.34	9.87	0.17	5.38	12.44	0.20	1.61	0.06	145	56	82	17	14	30	8	44	17.5	99	5.6	71	35	263	19.66	98.62	
DG-13	510	mst	53.96	0.52	14.67	10.10	0.18	5.67	13.21	0.12	1.54	0.05	148	52	81	16	13	41	6	42	19.5	98	5.5	70	38	240	20.05	98.69	
DG-14	510	mst	61.35	0.43	12.82	9.48	0.15	4.25	9.82	0.12	1.56	0.03	130	37	66	14	10	30	8	41	24.4	83	4.6	68	33	243	16.02	98.96	
DG-16	512	mst	59.09	1.32	32.84	4.31	0.01	1.13	0.18	0.19	0.86	0.07	218	109	99	37	24	68	9	61	23.4	52	17.8	197	30	231	17.38	98.76	
DG-17	572	mst	63.21	1.53	30.86	2.66	0.01	0.55	0.14	0.10	0.91	0.04	185	165	136	37	36	43	6	63	29.0	47	26.5	190	49	306	21.90	99.85	
DG-19	579	mst	65.55	1.19	24.24	5.42	0.01	1.31	0.14	0.15	1.96	0.04	578	96	119	31	27	38	2	86	20.0	58	12.5	132	29	494	19.31	99.08	
DG-24	595	fst	63.22	0.90	28.05	4.37	0.00	1.00	0.18	0.12	2.08	0.07	529	117	105	32	21	23	28	81	22.1	62	11.5	170	30	188	10.58	100.48	
DG-25	601	mst	56.08	1.30	31.59	6.17	0.02	1.74	0.20	0.15	2.72	0.03	584	116	253	37	32	58	4	132	30.9	63	31.3	196	39	249	18.68	99.98	
DG-29	616	mst	56.08	1.30	31.59	6.17	0.02	1.74	0.20	0.15	2.72	0.03	584	116	253	37	32	58	4	132	30.9	63	31.3	196	39	249	18.68	99.98	
<b>Unit 2</b>																													
DG-33	663	csst	79.38	0.47	11.78	2.05	0.04	0.55	2.85	0.09	2.78	0.02	231	70	32	15	11	15	96	12.4	98	14.8	59	29	207	4.23	99.60		
DG-34	671	csst	81.17	0.43	12.22	1.98	0.03	0.46	1.55	0.08	2.07	0.02	176	64	21	12	13	14	17	68	12.0	85	12.1	52	21	231	3.62	99.71	
DG-35	675	zst	82.02	1.15	28.59	4.14	0.01	0.86	0.36	0.17	2.66	0.04	379	85	148	35	25	43	31	100	27.4	88	28.2	170	19	245	10.66	100.22	
DG-36	677	csst	81.24	0.46	12.18	1.86	0.03	0.46	1.53	0.17	2.05	0.02	168	69	22	12	13	11	43	68	9.6	82	16.4	53	20	208	3.72	99.20	
DG-37	678	csst	78.32	1.26	11.11	3.52	0.07	0.81	3.09	0.20	1.57	0.03	125	44	13	31	41	39	54	10.0	87	41.7	90	40	458	5.17	98.74		
DG-38	716	csst	74.17	0.38	13.47	4.55	0.05	1.62	1.49	0.17	2.33	0.06	1067	22	73	15	9	28	19	74	15.5	216	4.8	100	11	84	2.53	99.08	
DG-39	723	csst	72.64	0.59	14.79	2.87	0.04	1.05	1.25	0.24	4.25	0.07	1622	113	43	16	40	29	25	108	12.0	280	8.6	66	21	880	1.59	99.66	
DG-40	724	csst	73.68	0.25	13.43	2.49	0.04	3.18	2.74	0.21	0.06	0.06	2082	2	29	10	7	18	20	82	9.4	294	3.9	35	84	2.74	98.31		
DG-41	726	csst	72.88	0.56	13.50	3.14	0.06	1.09	3.30	2.74	2.66	0.07	1757	67	46	13	14	17	19	78	12.0	277	7.7	69	24	567	3.18	98.92	
DG-42	727	csst	76.17	0.32	11.98	3.96	0.06	1.35	2.06	2.28	1.78	0.06	818	23	39	10	7	21	15	60	12.1	215	6.5	44	12	78	2.62	100.33	
DG-43	729	csst	75.66	0.35	12.15	3.93	0.06	1.32	2.37	1.85	0.06	2.27	0.06	849	20	54	11	9	21	17	60	12.3	227	5.0	45	15	83	2.91	99.63
<b>Unit 1</b>																													
DG-44	785	zst	64.84	0.82	16.15	6.77	0.10	2.49	2.76	0.21	3.19	0.21	1192	114	93	22	23	35	23	122	16.0	374	19.7	111	42	386	3.52	99.55	
DG-45	790	mst	65.17	0.87	15.85	6.77	0.10	2.85	2.65	0.21	3.03	0.21	1106	122	97	22	20	35	24	116	17.6	369	19.9	118	44	409	3.36	99.41	
DG-46	793	mst	65.30	0.81	15.78	6.71	0.10	2.43	2.87	0.21	3.03	0.21	1117	111	98	20	23	34	25	112	14.8	371	19.6	110	42	398	3.31	99.49	
DG-47	846	zst	60.82	1.20	17.95	8.39	0.12	3.48	1.49	1.58	4.73	0.23	1052	147	135	31	34	47	22	193	19.7	217	30.4	185	47	379	4.93	99.47	
DG-48	849	mst	75.22	0.41	12.72	3.22	0.07	1.10	3.51	1.73	0.08	710	59	29	12	21	19	63	10.0	439	6.9	47	34	355	3.39	99.89			
DG-49	853	fst	74.11	0.33	13.34	3.68	0.06	1.31	3.17	1.90	2.02	0.08	721	49	27	13	9	14	22	67	8.6	458	5.7	33	21	292	3.66	99.96	
DG-50	907	mst	59.50	0.99	17.72	9.44	0.12	3.79	1.74	1.69	4.76	0.25	1041	117	133	27	28	58	29	235	16.8	235	31.0	168	41	409	3.36	99.41	
DG-51	910	mst	75.62	0.30	12.53	2.61	0.05	0.93	3.62	2.44	1.83	0.08	701	47	17	13	9	10	21	59	10.6	324	4.8	34	20	231	3.47	98.70	
DG-52	912	mst	59.24	1.00	17.64	9.19	0.12	3.69	1.99	1.91	4.84	0.38	1051	133	118	27	29	57	28	236	16.7	242	30.0	164	43	251	4.50	99.16	
DG-53	914	mst	60.13	1.00	17.71	8.86	0.11	3.57	1.77	1.78	4.85	0.22	1052	124	12														

Table 1 (ctd). Whole rock major (wt%) and trace element (ppm) analyses of Gondwana Group sandstones and mudrocks from GDH-45, Khalaspir basin, Bangladesh.

SaNr	Depth (m)	Lith	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Fe <sub>2</sub> O <sub>3</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	Ba	Ce	Cr	Ga	Nb	Ni	Pb	Rb	Sc	Sr	Th	V	Y	Zr	LoI	Total
<b>Unit 1 (ctd)</b>																												
DG-64	976	mst	61.75	0.98	15.16	9.48	0.13	3.60	2.49	3.03	3.15	0.24	91.1	136	123	24	29	54	27	97	18.8	281	23.6	157	45	303	3.31	99.93
DG-65	977	msst	75.03	0.36	12.22	3.07	0.06	1.09	2.90	2.84	2.33	0.11	63.4	64	33	14	13	24	77	9.4	220	7.0	40	25	293	2.29	99.32	
DG-66	979	zst	63.04	0.88	15.38	8.00	0.13	3.21	2.63	2.92	3.55	0.27	99.9	118	115	21	28	47	28	116	16.7	297	20.2	129	47	315	2.91	100.16
DG-67	982	fsst	70.83	0.50	13.76	4.31	0.07	1.54	3.85	2.84	2.13	0.17	57.9	83	57	17	15	22	25	68	12.0	305	10.9	67	33	393	3.27	99.60
DG-68	985	zst	61.25	1.22	14.53	9.73	0.15	3.58	2.84	2.79	3.49	0.42	92.7	186	156	27	38	62	37	104	20.6	249	31.7	173	67	454	3.04	99.59
DG-69	988	mst	55.79	0.77	14.84	6.70	0.11	2.51	3.34	2.84	2.85	0.23	78.7	120	92	21	21	37	30	89	14.2	321	18.4	115	42	351	3.41	99.58
DG-70	989	fsst	72.93	0.31	14.31	2.51	0.04	4.95	4.03	2.78	2.02	0.11	52.9	56	28	14	9	13	19	67	8.4	351	6.0	34	22	296	2.54	99.38
DG-71	991	mst	63.95	0.79	15.58	7.10	0.11	2.68	3.19	2.68	3.72	0.20	94.2	106	94	21	24	46	29	134	15.7	337	19.6	109	38	266	3.15	99.62
DG-72	994	zst	67.07	0.73	14.51	6.05	0.10	2.25	3.35	3.00	2.71	0.23	69.3	125	76	19	23	32	26	87	13.4	317	18.2	94	42	436	3.22	99.57
DG-73	997	mst	58.36	1.23	16.84	9.57	0.16	3.58	2.47	2.72	4.83	0.25	112.3	150	156	28	36	70	31	209	20.7	274	41.6	167	49	153	2.99	99.75
DG-74	998	fsst	72.93	0.41	13.22	3.72	0.07	1.32	2.79	2.91	2.49	0.14	66.4	73	44	15	13	18	26	81	9.6	239	8.8	11	27	297	2.49	99.20
DG-75	1000	mst	61.00	1.05	15.69	8.89	0.13	3.42	2.60	2.77	4.18	0.27	104.9	135	131	23	31	57	30	153	17.3	284	30.2	147	27	256	2.71	99.64
DG-76	1003	msst	73.23	0.73	11.67	5.25	0.14	1.55	2.82	2.37	2.11	0.12	59.1	121	63	14	20	15	21	69	13.6	188	14.6	84	49	647	1.91	99.24
DG-77	1006	fsst	67.56	0.56	15.47	4.89	0.08	1.85	4.67	2.25	5.01	0.18	67.0	82	58	18	16	28	24	85	12.7	418	13.3	75	32	307	3.22	99.57
DG-78	1008	zst	58.32	1.12	15.95	10.68	0.16	4.06	2.95	3.72	3.97	0.39	99.8	221	110	26	30	52	31	122	20.2	275	44.1	176	61	227	3.33	99.57
DG-79	1012	mst	54.45	1.60	16.48	12.76	0.20	4.69	2.09	3.08	4.36	0.30	114.3	206	125	29	35	58	41	131	26.8	259	73.2	226	68	253	3.89	99.96
DG-80	1015	zst	63.99	0.66	16.13	6.82	0.11	2.44	4.09	2.60	2.93	0.23	85.6	96	43	19	14	24	33	88	14.7	396	22.6	95	44	417	4.12	99.57
DG-81	1032	msst	69.12	0.53	14.77	4.51	0.07	1.60	4.25	2.57	2.40	0.19	61.4	83	55	15	16	33	22	79	11.5	357	11.5	68	31	313	3.69	99.61
DG-82	1034	mst	70.68	0.54	13.18	4.79	0.07	1.53	3.90	2.71	2.42	0.20	60.7	83	65	15	15	26	25	78	13.9	262	12.1	76	34	286	3.17	99.00
DG-83	1041	mst	71.98	0.32	13.52	3.28	0.05	0.97	5.38	2.38	2.03	0.10	52.1	47	35	12	10	24	67	9.0	341	6.0	39	21	222	4.64	98.97	
DG-84	1043	msst	71.19	0.54	12.51	4.72	0.09	1.77	3.89	2.72	2.39	0.17	58.2	81	52	14	16	24	23	78	13.1	202	10.5	62	37	344	3.14	99.39
DG-85	1046	mst	57.14	1.18	17.00	10.25	0.14	4.53	2.35	2.55	4.64	0.21	93.3	150	181	25	39	106	29	243	16.9	249	35.5	150	48	117	3.17	99.73
DG-86	1049	mst	55.70	1.21	17.23	11.15	0.15	4.94	2.28	2.49	4.69	0.19	98.4	147	194	26	38	110	28	240	17.8	239	37.0	162	47	97	4.07	99.84
DG-87	1052	mst	56.78	1.16	16.76	10.34	0.14	4.45	2.81	2.45	4.91	0.20	89.9	146	179	24	37	102	31	288	23.6	237	35.1	154	46	105	4.04	99.83
DG-88	1055	mst	59.17	1.05	16.43	9.45	0.13	4.11	2.43	2.70	4.33	0.20	86.7	135	153	22	34	86	29	236	16.7	244	30.5	146	44	137	3.20	99.87
DG-89	1058	mst	60.63	1.16	16.07	8.63	0.12	3.65	2.65	2.90	3.97	0.23	93.6	141	162	22	38	89	32	163	17.8	258	34.3	145	51	188	2.83	99.61
DG-90	1059	mst	59.22	1.00	16.54	9.52	0.13	4.16	2.14	2.88	4.19	0.21	90.2	123	142	21	32	84	23	208	13.1	202	10.5	62	37	344	3.14	99.39
DG-91	1061	mst	54.81	1.27	17.61	11.21	0.14	4.92	2.09	2.26	5.45	0.18	97.3	123	194	25	39	112	25	341	19.9	212	34.9	167	42	56	3.99	100.11
DG-92	1064	mst	56.62	1.22	17.25	10.44	0.14	4.64	2.18	2.57	4.75	0.20	96.9	117	193	24	40	114	30	254	20.1	231	33.6	155	43	76	3.55	99.71
DG-93	1067	mst	54.73	1.45	17.61	11.12	0.16	4.82	2.48	2.56	4.84	0.23	108.4	214	25	44	116	24	233	22.8	241	38.5	187	51	124	3.85	99.89	
DG-94	1070	fsst	66.93	0.68	12.72	5.25	0.10	1.65	8.15	2.26	2.07	0.18	49.3	112	56	13	18	20	22	67	16.6	277	17.3	81	43	629	6.56	99.03
DG-95	1070	fsst	61.83	0.69	16.83	8.15	0.10	2.79	5.78	2.55	2.36	0.21	72.2	90	75	18	48	23	38	14.4	445	13.3	97	31	235	4.82	99.35	
DG-96	1076	mst	56.92	1.12	16.97	10.61	0.14	4.49	2.39	2.57	4.57	0.20	102.1	127	167	24	34	98	34	264	17.4	278	32.0	156	43	110	3.72	99.84
DG-97	1077	mst	24.98	0.76	7.58	6.59	0.73	2.65	5.26	3.05	2.66	0.31	47.7	111	90	11	20	58	8	128	15.0	227	21.2	60	51	49	28.85	100.44
DG-98	1078	zst	62.61	0.77	16.69	6.47	0.11	2.65	4.64	2.75	3.01	0.21	89.7	101	101	17	24	30	30	103	14.6	392	17.5	86	38	258	4.18	99.20
DG-99	1079	mst	59.20	1.05	16.36	9.44	0.13	4.16	2.39	3.31	3.73	0.23	105.6	116	157	24	33	87	26	125	19.3	263	25.0	154	44	114	3.05	99.79
DG-100	1081	mst	58.45	1.06	16.67	9.72	0.16	4.28	3.03	3.87	3.07	0.21	108.2	110	145	23	33	84	33	137	17.0	268	27.0	157	43	129	3.34	100.12
DG-101	1082	mst	61.88	0.98	15.68	8.14	0.13	3.46	2.36	3.63	3.52	0.22	98.8	108	126	22	30	64	29	111	18.6	252	22.5	147	43	143	2.97	99.86
DG-102	1085	mst	56.80	1.11	17.00	10.51	0.16	4.68	2.30	3.17	4.08	0.21	114.6	100	166	25	33	98	28	253	27.9	173	41	91	3.51	99.88		
DG-103	1086	zst	61.10	1.03	15.73	8.39	0.14	3.50	3.03	3.31	3.55	0.23	100.2	113	154	23	33	70	27	113	17.0	282	21.9	143	45	183	4.36	99.40
DG-104	1088	mst	64.43	0.73	15.34	6.76	0.11	2.78	3.74	3.03	2.88	0.20	82.0	96	93	19	23	38	27	94	14.7	322</td						

seconds for mudrocks and 40-45 seconds for sandstones. Additional description of the crushing procedures used is given by Roser *et al.* (1998). Total loss on ignition (LOI) was determined gravimetrically on 8-10 g splits of oven-dried powder (>24 hours at 110°C), from net weight loss after ignition in a muffle furnace for at least 2 h at 1020°C. The ignited samples from the LOI determinations were then hand-ground in an agate pestle and mortar and returned to a 110°C oven before preparation of glass fusion beads for X-ray fluorescence (XRF) analysis.

Abundances of major elements and 14 trace elements (Ba, Ce, Cr, Ga, Nb, Ni, Pb, Rb, Sc, Sr, Th, V, Y, Zr) were determined at Shimane University. Analyses were made on glass fusion beads prepared using an alkali flux (80% lithium tetraborate ( $\text{Li}_2\text{B}_4\text{O}_7$ ) and 20% lithium metaborate ( $\text{LiBO}_2$ )), with a flux to sample ratio of 2:1 (Kimura and Yamada, 1996). The instrument used was a Rigaku RIX 2000 XRF spectrometer, equipped with a Rh-anode X-ray tube. Calibration and correction for spectral interferences followed the methodology of Kimura and Yamada (1996). Internal correction for intra-run drift was made using secondary calibration against ten Geological Survey of Japan rock standards spanning the compositional range from gabbro (JGb-1) through to granite (JG-2).

## Results

Major and trace element analyses of the Gondwana Group sediments are listed in Table 1, on an anhydrous normalized basis. Each unit shows variations in their elemental composition. LOI values in Unit 1 and 2 are generally low (<5 wt%), whereas all except two samples in Unit 3 have considerably higher values (>10 wt%), and a maximum of 21.9 wt%. There is no consistent association of the higher LOI values in Unit 3 with a single element. Some samples with high LOI are also enriched in CaO, whereas others have low CaO, but are enriched in  $\text{Al}_2\text{O}_3$ . The higher LOI values in Unit 3 are thus most likely due to either high  $\text{CaCO}_3$  or clay contents, or a combination of both. However, enrichment in organic matter may also be a factor. Nevertheless, when the anhydrous normalized data are considered, some chemical contrasts and variations are observed between units and between lithotypes.

Geochemical compositions of sediments are mainly controlled by source rock composition, weathering in the source area, hydrodynamic fractionation of constituent minerals during transport and deposition, and by post-depositional diagenesis. Especially, the system is strongly influenced by relative proportions of quartz plus feldspar versus clays i.e. compositions of the end products are influenced by the dilution effect of quartz ( $\text{SiO}_2$ ) or the abundance of clays ( $\text{Al}_2\text{O}_3$ ).

In all units  $\text{SiO}_2$  contents are characteristically greater in sandstones than in companion mudrocks; conversely contents of  $\text{Al}_2\text{O}_3$ ,  $\text{TiO}_2$  and most other elements are greater

in the mudrocks than in the sandstones. Average  $\text{SiO}_2$  contents in sandstones increase from Unit 1 (71.34 wt%) to Unit 2 (76.53 wt%), and subsequently decrease in Unit 3 (63.55 wt%). However, there is no significant variation of  $\text{SiO}_2$  in the mudrocks of Unit 1 (60.05 wt%), Unit 2 (62.02 wt%) and Unit 3 (60.87 wt%). Average  $\text{Al}_2\text{O}_3$  contents of the mudrocks increases significantly from Unit 1 (15.98 wt%) through Unit 2 (28.59 wt%) to Unit 3 (30.51 wt%). Average  $\text{Al}_2\text{O}_3$  content of Unit 3 sandstones (20.65 wt%) is also greater than equivalents in Unit 1 (13.48 wt%). Average  $\text{CaO}$  contents are similar in all units (<5 wt%), and are generally lower than in average upper continental crust (Taylor and McLennan, 1985). Some samples in Unit 1 contain anomalously high  $\text{CaO}$  (maximum, 52.68 wt%). Average  $\text{Na}_2\text{O}$  content decreases significantly from Unit 1 through Unit 2 and Unit 3, in both sandstones and mudrocks. Contents of  $\text{K}_2\text{O}$  also decrease upward from Unit 1 to Unit 3. Overall, most of the major elements except  $\text{Al}_2\text{O}_3$  decrease stratigraphically upward, from Unit 1 to Unit 3.

Barium is the most abundant trace element in both mudrocks (M) and sandstones (S) in Unit 1 (M=970 ppm, S=640 ppm), Unit 2 (M=379 ppm, S=889 ppm) and Unit 3 (M=268 ppm, S=234 ppm). However, average Ba contents are lower in the younger units. Average Rb contents in Unit 1 (M=157 ppm, S=73 ppm) are greater than in Unit 2 (M=100 ppm, S=75 ppm) and Unit 3 (M=69 ppm, S=51 ppm), and thus Rb also decreases to the younger units, similar to Ba. Strontium contents also show similar behavior. Average Sr content in Unit 1 (M=282 ppm, S=315 ppm) decreases in Unit 2 (M=88 ppm, S=186 ppm), and is lower still in Unit 3 (M=60 ppm, S=76 ppm). Average contents of all these large ion lithophile elements thus decrease from the oldest to the youngest unit.

Average contents of Nb, Y, Th and Zr do not show any great variation between the units, especially in the mudrocks. However, average Zr concentration in Unit 1 sandstone (359 ppm) is greater than in Unit 2 (288 ppm) and Unit 3 (287 ppm). Average Ga concentration increases from Unit 1 (M=23 ppm, S=15 ppm) through Unit 2 (M=35 ppm, S=13 ppm) and Unit 3 (M=32 ppm, S=22 ppm). Average Ga concentration thus increases from the oldest to youngest unit, paralleling the trend in  $\text{Al}_2\text{O}_3$ .

Among the ferromagnesian elements, average Cr and Ni contents in the mudrocks show relatively little contrast between units. However, average Cr content in Unit 3 sandstones (98 ppm) is significantly greater than in Unit 1 and 2 equivalents (40 and 47 ppm, respectively). Similarly, average Ni content is higher in Unit 3 sandstone (S=32 ppm) than in Unit 1 and 2 sandstones (both averaging 21 ppm). Average Sc concentration increases from the oldest to youngest, from Unit 1 (M=17 ppm, S=12 ppm) to Unit 2 (M=27 ppm, S=12 ppm) and Unit 3 (M=25 ppm, S=19 ppm). Average V content also increases slightly from oldest to youngest: Unit 1 (M=139 pp, S=58 ppm); Unit 2 (M=170 ppm, S=61 ppm); Unit 3 (M=157 ppm, S=111 ppm).

These broad variations in major and trace element abundances between units and lithotypes thus record the signatures of differing factors and processes.

### Conclusions

Average geochemical compositions of sandstones and mudrocks in the three units of the Gondwana Group show significant variations. These contrasts represent the imprint of the factors that influence the composition of the sediments. Gondwana sediments on other continental blocks (India, Australia, South America, and South Africa) record the signature of dramatic climate change. The variations in the geochemical composition of the sediments of the Gondwana Group in Bangladesh will be interpreted in detail in a future publication, with the emphasis on climatic influence and possible correlation with the Gondwana successions on other continental blocks.

### Acknowledgements

We thank the Geological Survey of Bangladesh (GSB) for access to drill core GDH-45, with special thanks to Dr. Nehal Uddin, Director of the GSB, for his kind permission and suggestions during sample collection. We also thank Drs Syed Samsuddin Ahmed, Sultan-Ul-Islam, Mrinal Kanti Roy and Ismail Hossain for their encouragement and suggestions during sample collection. This work was supported by a Monbukagakusho scholarship to DKR and internal research funding to BPR.

### References

- Akther, A. 2001. Gondwana sediment of Bangladesh and its correlation with those of other regions of the world on the basis of spore-pollen. *Gondwana Research*, **4**, 135-136.
- Bhatia, M. R. 1983. Plate tectonics and geochemical composition of sandstones. *Journal of Geology*, **91**, 611-627.
- Catuneanu, O., Wopfner, H., Eriksson, P. G., Cairncross, B., Rubidge, B. S., Smith, R. M. H. and Hancock, P. J. 2005. The Karoo basins of south-central Africa. *Journal of African Earth Science*, **43**, 211-253.
- Clift, P. D., Hodges, K. V., Heslop, D., Hannigan, R., Long, H. V. and Claves, G. 2008. Correlation of Himalayan exhumation rates and Asian monsoon intensity. *Nature Geosciences*, **1**, 875-880.
- Fedo, C. M., Nesbitt, H. W. and Young, G. M. 1995. Unraveling the effects of potassium metasomatism in sedimentary rocks and paleosols, with implications for paleoweathering conditions and provenance. *Geology*, **23**, 921-924.
- Frielingsdorf, J., Islam, Sk.A., Block, M., Rahman, M. M. and Rabbani, M. G. 2008. Tectonic subsidence modeling and Gondwana source rock hydrocarbon potential, Northwest Bangladesh modeling of Kuchma, Singra and Hajipur wells. *Marine and Petroleum Geology*, **25**, 553-564.
- Hossain, H. M. Z., Islam, M. S. U., Ahmed, S. S. and Hossain, I. 2002. Analysis of sedimentary facies and depositional environments of the Permian Gondwana sequence in borehole GDH-45, Khalaspur Basin, Bangladesh. *Geoscience Journal*, **6**, 227-236.
- Hossain, I., Ahmed, S. S. and Islam, M. S. 2000. Provenance of Gondwana sandstones in the borehole GDH-40, Barapukuria basin, Dinajpur, Bangladesh. *Bangladesh Journal of Geology*, **19**, 35-43.
- Hossain, I., Tsunogae, T., Rajesh, H. M., Chen, B. and Arakawa, Y. 2007. Palaeoproterozoic U-Pb SHRIMP zircon age from basement rocks in Bangladesh: A possible remnant of the Columbia supercontinent. *Comptes Rendus Geosciences*, **339**, 679-686.
- Islam, M. N., Uddin, M. N., Resan, S. A., Islam, M. S. and Ali, M. W. 1992. Geology of the Khalaspur Coal Basin, Pirganj, Rangpur, Bangladesh. *Records of the Geological Survey of Bangladesh*, **6**, 1-76.
- Islam, M. R., 2009. Geo-environmental hazards associated with multi-slice longwall mining of the Gondwana Barapukuria coal basin, NW Bangladesh: Constraints from numerical simulation (Ph.D. thesis), University of the Ryukyus, 183 p.
- Islam, M. S. U., Chowdhury, K. R. and Ishiga, H. 2004. Geochemistry of the Gondwana sedimentary rocks from the Barapukuria Basin, Bangladesh. *Bangladesh Geoscience Journal*, **10**, 83-102.
- Khan, F. H., 1991. *Geology of Bangladesh*. The University Press Limited, Dhaka, 207 p.
- Kimura, J.-I. and Yamada, Y., 1996. Evaluation of major and trace element XRF analyses using a flux to sample ratio of two to one glass beads. *Journal of Mineralogy, Petrology and Economic Geology*, **91**, 62-72.
- Martini, I. P. 1997. *Late glacial and postglacial environmental changes, Quaternary, Carboniferous-Permian, and Paleozoic*. Oxford, Oxford University Press, 343 p.
- McLennan, S. M., Hemming, S., McDaniel, D. K. and Hanson, G. N. 1993. Geochemical approaches to sedimentation, provenance and tectonics. In: Johnsson, M. J., and Basu, A., eds, *Processes Controlling the Composition of Clastic Sediments*. Geological Society of America Special Paper, **284**, 1-19.
- Nesbitt, H. W. and Young, G. M. 1982. Early Proterozoic climates and plate motions inferred from major element chemistry of lutites. *Nature*, **299**, 715-717.
- Ohta, T. 2008. Measuring and adjusting the weathering and hydraulic sorting effects for rigorous provenance analysis of sedimentary rocks: a case study from the Jurassic Ashikita Group, south-west Japan. *Sedimentology*, **55**, 1687-1701.
- Rahman, M. H. and Ahmed, F. 1995. Petrographic analysis of the Gondwana clastic sedimentary rocks encountered at Barapukuria area, Dinajpur, Bangladesh. *Bangladesh Geoscience Journal*, **1**, 1-20.
- Rieu, R., Allen, P. A., Plötze, M. and Pettke, T. 2007. Climatic cycles during a Neoproterozoic "snowball" glacial epoch. *Geology*, **35**, 299-302.
- Roser, B. P. and Korsch, R. J. 1986. Determination of tectonic setting of sandstone mudstone suites using SiO<sub>2</sub> and K<sub>2</sub>O/Na<sub>2</sub>O ratio. *Journal of Geology*, **94**, 635-650.
- Roser, B. P. and Korsch, R. J. 1988. Provenance signatures of sandstone-mudstone suites determined using discrimination function analysis of major-element data. *Chemical Geology*, **67**, 119-139.
- Roser, B. P., Sawada, Y. and Kabeto, K. 1998. Crushing performance and contamination trials of a tungsten carbide ring mill compared to agate grinding. *Geoscience Reports of Shimane University*, **17**, 1-11.
- Scheffler, K., Hoernes, S. and Schwark, L. 2003. Global changes during Carboniferous-Permian glaciations of Gondwana: Linking polar and equatorial climate evolution by geochemical proxies. *Geology*, **31**, 605-608.
- Singh, B. D. and Singh, A. 2004. Observations on Indian Permian Gondwana coals under fluorescence microscopy: An overview. *Gondwana Research*, **7**, 143-151.
- Stephenson, M. H., Angiolini, L. and Leng, M. J. 2007. The Early Permian fossil record of Gondwana and its relationship to deglaciation: a review. In: M. Williams, A. Heywood, F. J. Gregory and D. N. Schmidt, eds, *Deep-Time perspective on Climate Change: Marrying the signal from Computer Models and Biological Proxies*. The Micropalaeontological Society Special Publications, The Geological Society, London, 169-189.
- Suttner, L. J. and Dutta, P. K. 1986. Alluvial sandstone composition and paleoclimate, I. Framework mineralogy. *Journal of Sedimentary Petrology*, **56**, 329-345.
- Taylor, S. R. and McLennan, S. M. 1985. *The continental crust: its composition and evolution*. Oxford, Blackwell Scientific Publications, 312 p.

(Received: Oct. 3, 2011, Accepted: Oct. 28, 2011)